Project METRO:

Upper ocean stratification and response to

westerly wind bursts:

Links between intraseasonal variability and ENSO

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This study:

How are WWBs & warm pool advection related?

- Why do only some WWBs displace the warm pool eastward?
- Which processes are important?
- What are the feedbacks between WWBs and ENSO?

Strategy:

- ORCA025 model DRAKKAR run
 - Based on the NEMO OGCM
 - Forcing: shortwave fluxes & surface meteorological variables using the DFS3 bulk formula (based on ERA-40); radiation from ISCCP.
 - output resolution: 5-day, 0.25°x0.25°,
 - vertical resolution 6-m (surface) to 250-m (depth)
 - output fields (present study): temperature, salinity, zonal currents
- Define WWBs:
 - At each time-step, locate points for which: { wind stress > 0.02 N/m²
 - wind stress anomaly > 0.06 N/m²
 - WWBs defined as patches of having zonal extent > 10° & duration ≥ 5 days
 - Using 1980-2007 (28 years, ~120 WWBs)







Importance of *series* of WWBs?



Series of WWBs and warm pool displacement: relationship to ENSO

Nino-3.4 index & annual warm pool displacement 100 Annual sum of WWB-related atior 0.5 warm pool 50 displacement O (degrees lon.) -0.4 -12 0 6 12 6 Lag (months) Nino 3.4 index (°C) Nino-3.4 warming follows warm pool 1980 1990 displacement by 2000 2010 3–9 months

Lagged correlation between





What processes drive surface current anomalies? Simplified momentum balance:



Below WWB: wind stress dominates

At warm pool edge: zonal advection

Along Kelvin wave path: pressure gradient

Momentum balance (integrated) for all WWBs Below the WWB



dominates the zonal current acceleration

 t_o = start of WWB t_f = end of WWB

Momentum balance (integrated) for all WWBs At the warm pool edge =t ∂u \overline{x} /\11 \mathcal{U} dt \sim оĤ $\rho \partial x$ surface current wind forcing advection ndient forci acceleration R=0.49 R=0.35 ² R=0.48 •

At the edge of the warm pool, advection and pressure gradient are also important $t_o = \text{start of WWB}$ $t_f = \text{end of WWB}$

Summary:

- 1. Westerly wind bursts drive eastward zonal surface currents:
 - Wind stress dominates current acceleration locally.
 - At the edge of the warm pool, **momentum advection** and **zonal pressure gradient** are also critical.
- 2. Warm pool displacement is more sensitive to surface currents than to local temperature gradient.
- 3. WWB-related warm pool displacement leads El Niño-related warming by 3–9 months.

Next steps:

- Re-run model with higher time resolution and full momentum budget:
 - Diagnose processes and quantify impacts more carefully.
- Explore the impacts of temperature and salinity variations:
 - Pressure gradient forcing.
 - Effect of barrier layers.
- Explore the feedbacks between ENSO and WWBs:
 - How large-scale conditions related to ENSO impact WWB properties (occurrence, intensity, location) and upper ocean response, and how this can feed back onto ENSO via the warm pool.