

A perspective on “Radar Hydrology” based on space-time scales

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*... and colleagues + students at LTHE and elsewhere



Outline

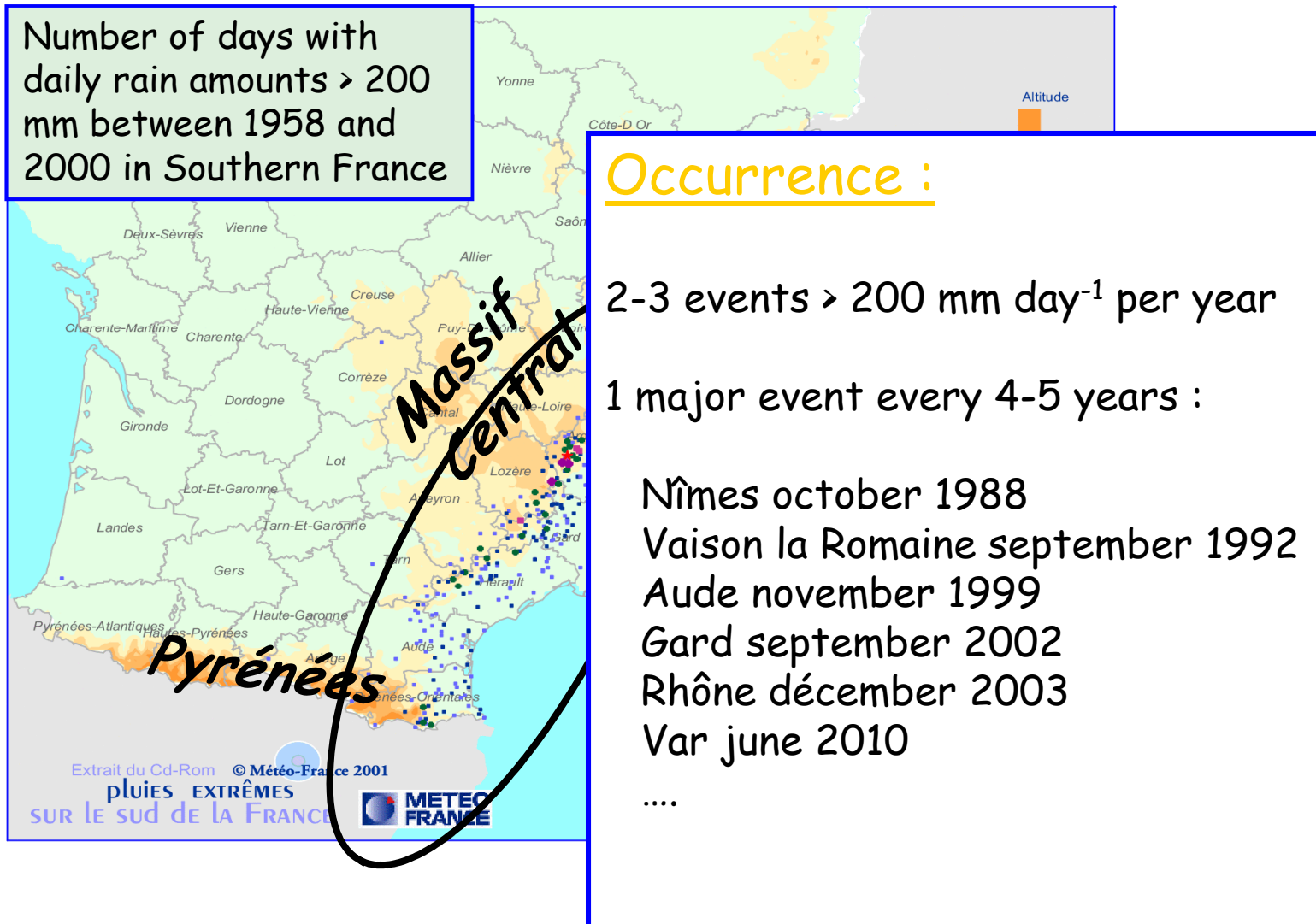
- Hydrometeorological risk prediction in the Mediterranean region
- Radar QPE: error models
- Hydrological observation and modelling

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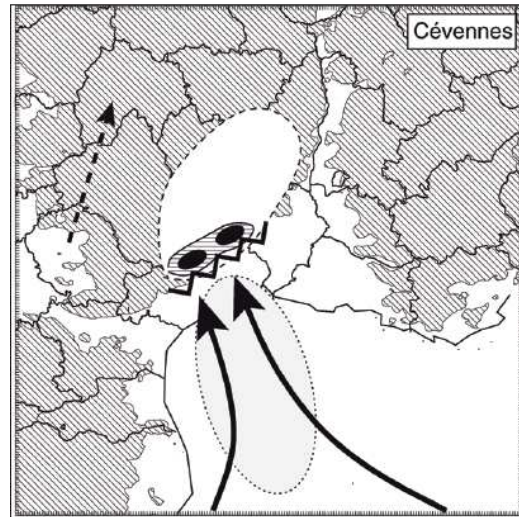
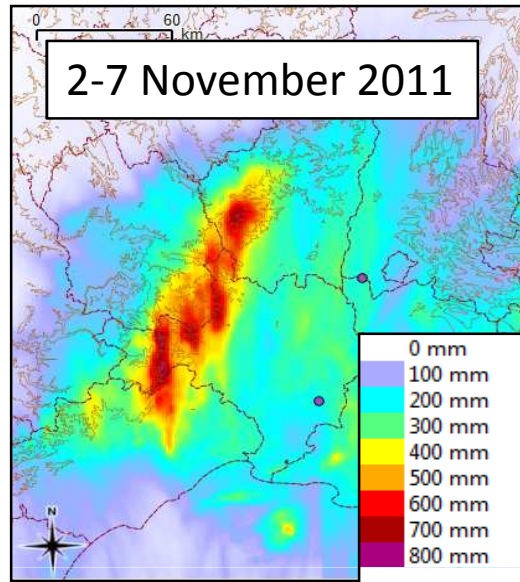
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Heavy rain events in France : location and occurrence

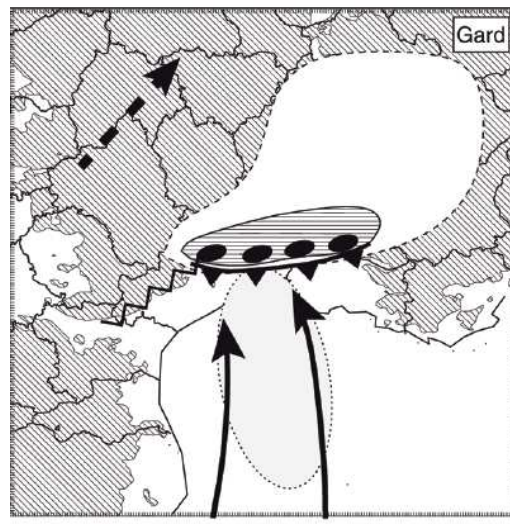
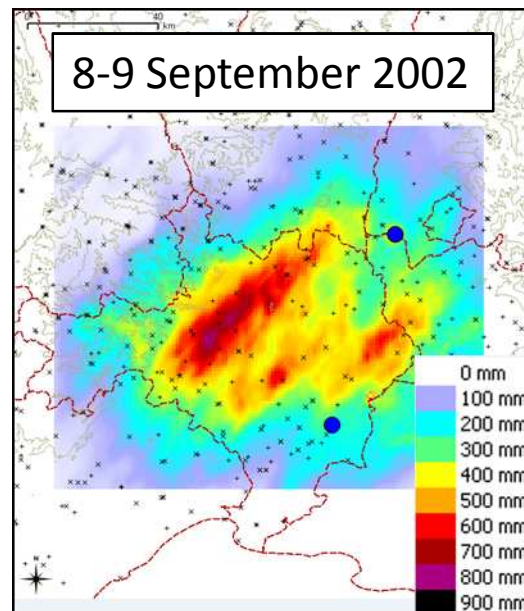
Météo France 2001 (CD-rom Pluies extrêmes dans le sud de la France)



A variety of convective systems

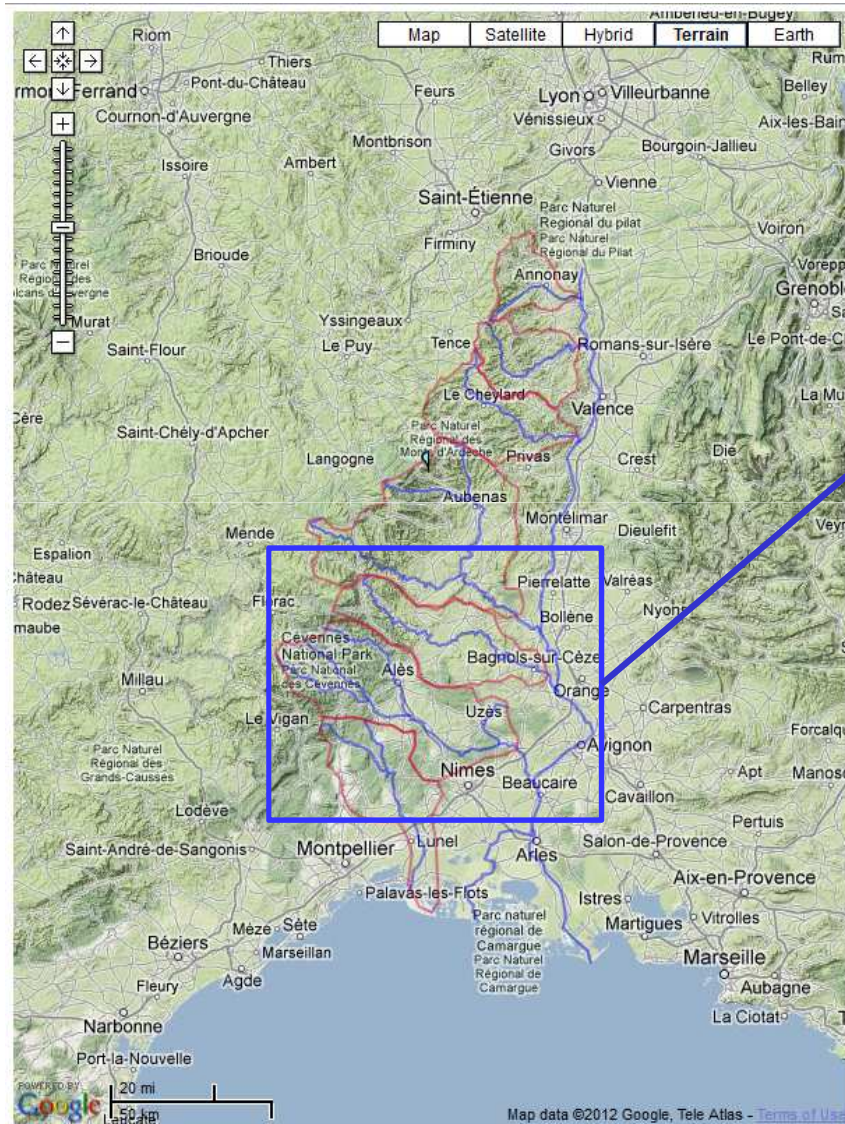


- Long lasting events
- Shallow / banded convection
- Orographic forcing



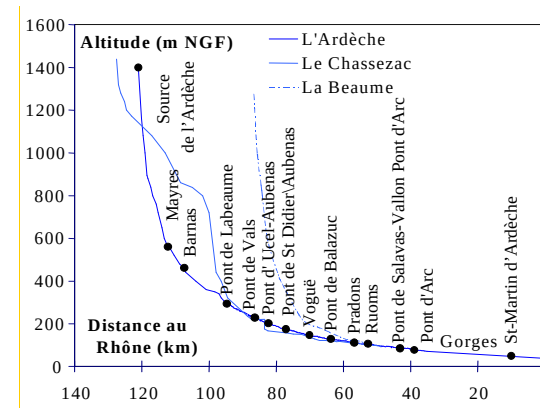
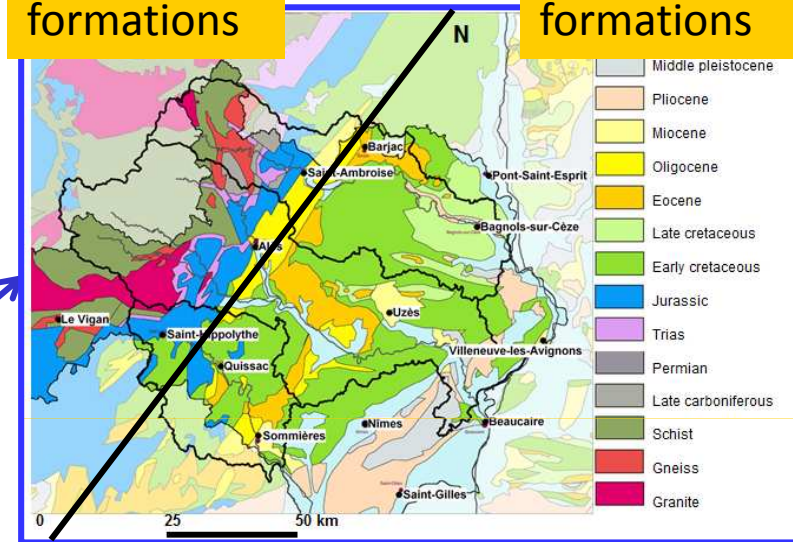
- V-shaped Mesoscale Convective Systems
- Deep convection
- Stationarity, dynamic forcing (cold pool)

Geomorphological features of the Cévennes watersheds



Primary Era
formations

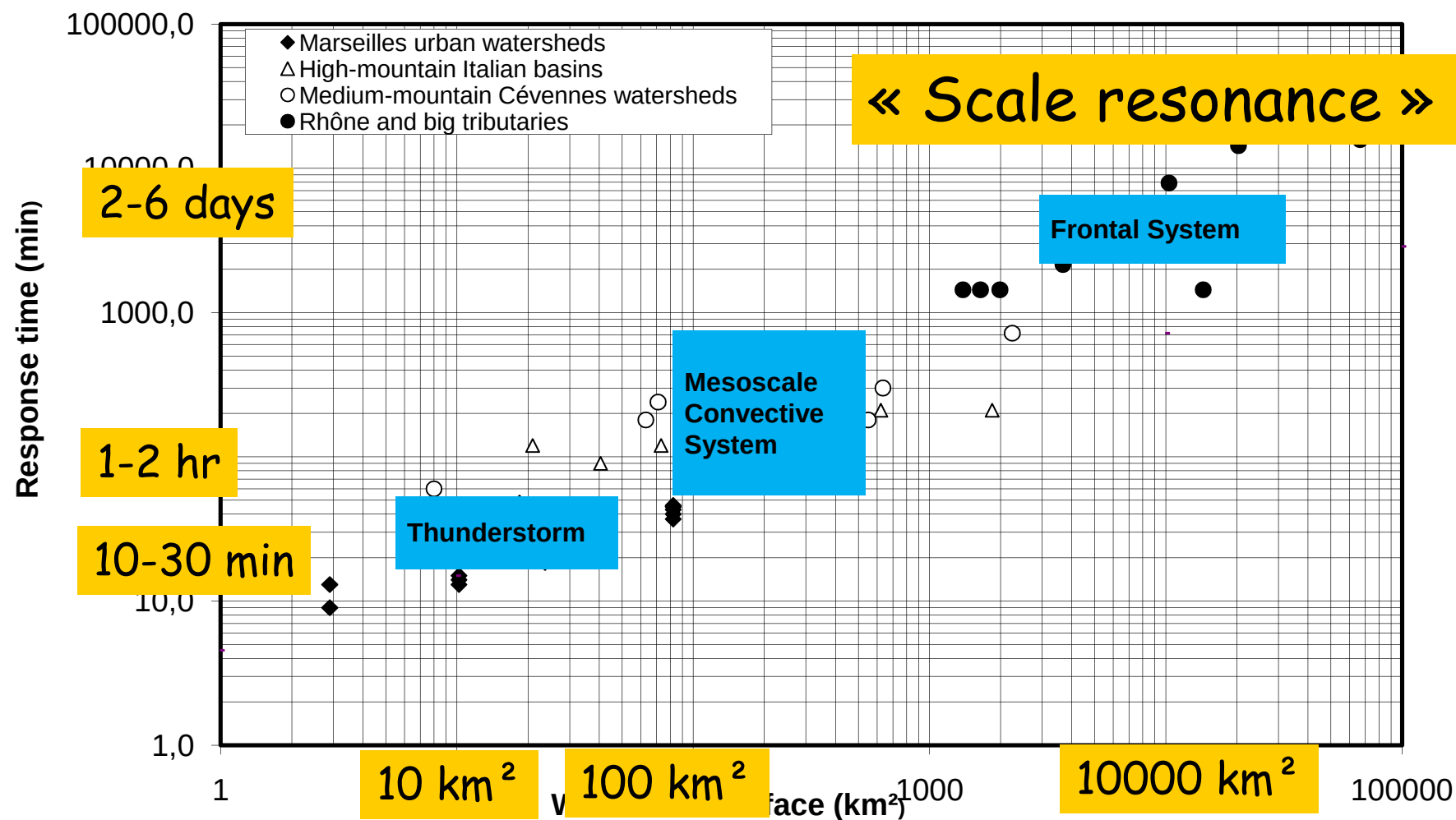
Sedimentary
formations



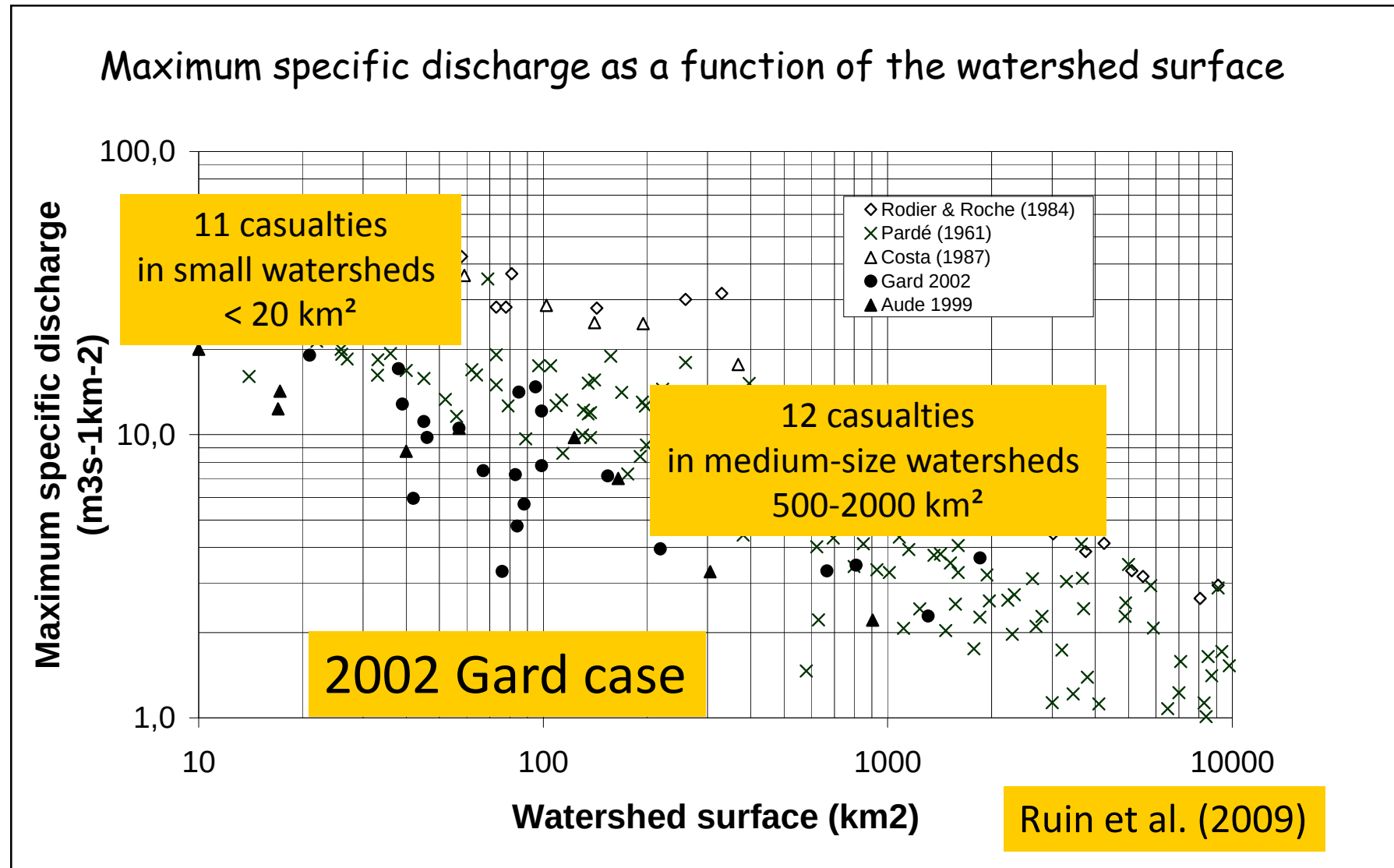
steep karst
slopes

Hydrological response: time to peak

Hydrologic response time of Mediterranean watersheds as a function of watershed surface



Severity of the hydrological response



We should be able to predict HPEs and the associated flash-floods and floods over a **wide range of spatial scales (local to regional)** and to communicate the warnings and alerts with **appropriate lead times** to the end-users.

« Prediction for Ungauged Basins »

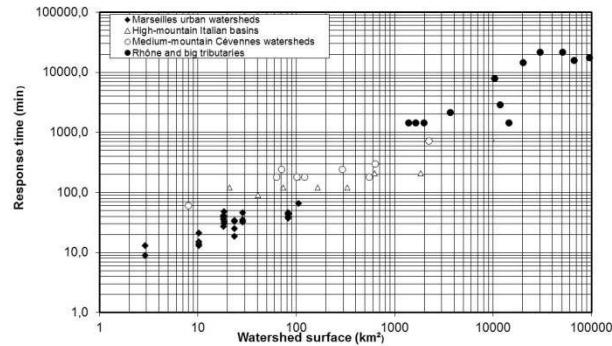
« Change of scale problem »

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Required space-time resolution of rainfall measurements

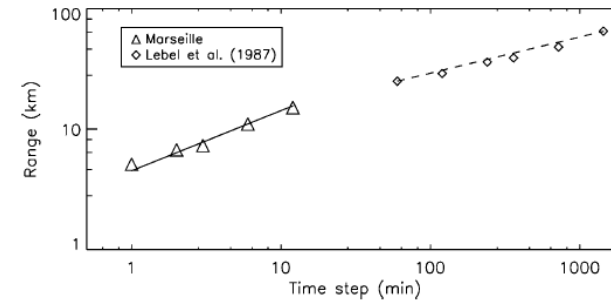
Time to peak



$$\Delta t = T_p / 3 = 1.1 A^{0.75}$$

(min) (km²)

Rainfall decorrelation distance



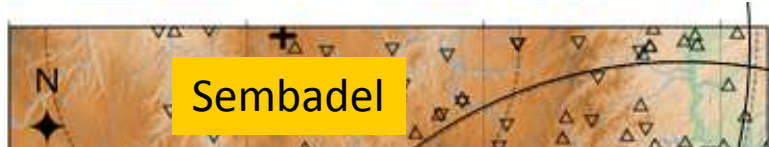
$$1 < \Delta t < 20 \text{ min}: \Delta r = \frac{R}{3} = 4.5 \Delta t^{0.5}$$

$$1 < \Delta t < 24 \text{ hrs}: \Delta r = \frac{R}{3} = 25 \Delta t^{0.3}$$

Watershed surface (km ²)	Time resolution	Spatial resolution
1	1 min	1.6 km
10	6 min	3,7 km
100	35 min	6,4 km
1000	3,3 hrs	15 km
10000	18 hrs	36 km

Lebel et al. (1987); Berne et al., (2004)

Rainfall observation system of the OHM-CV



Actual spatial resolution	Actual time resolution	Effective time resolution	Watershed scale properly sampled
11 km	1 h	2,6 h	735 km ²

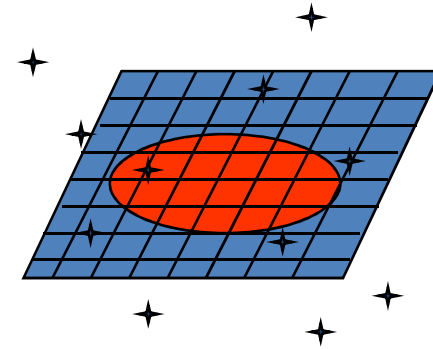
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32000 km²; 0-1800 m asl
 1 hour : 250 raingauges
 24 hours : 400 raingauges
 2 S-band radars
 2 C-band radars
 2 disdrometers
 11 meteo-GPS stations

Error model: reference rainfall derived from raingauge data

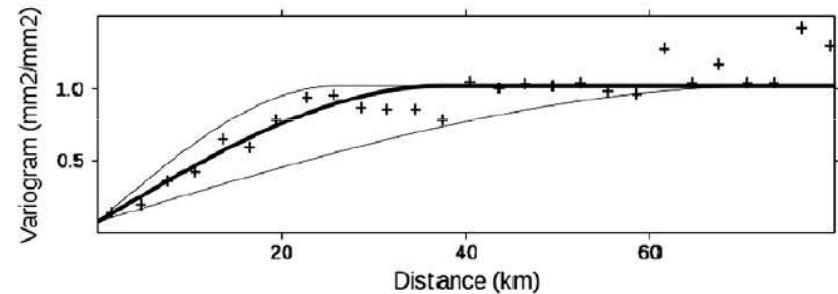
True unknown rainfall

$$R_{AT}(\underline{x}, t) = \frac{1}{A} \frac{1}{T} \iint_{A, T} R(\underline{u}, v) d\underline{u} dv$$



Radar QPE

$$R_{AT}^*(\underline{x}, t) = \frac{1}{N_A} \sum_{i=1}^{N_A} R_T^*(a_i, t)$$

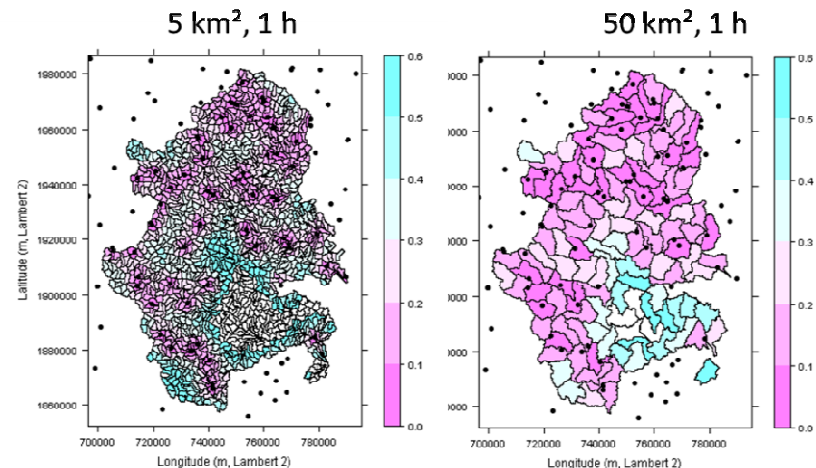


Reference rainfall:

Ordinary Kriging of raingauge values

Linear estimator: $R_{AT}^{ref}(\underline{x}, t) = \sum_{i=1}^{N_g} \lambda_i G(\underline{x}_i, t)$

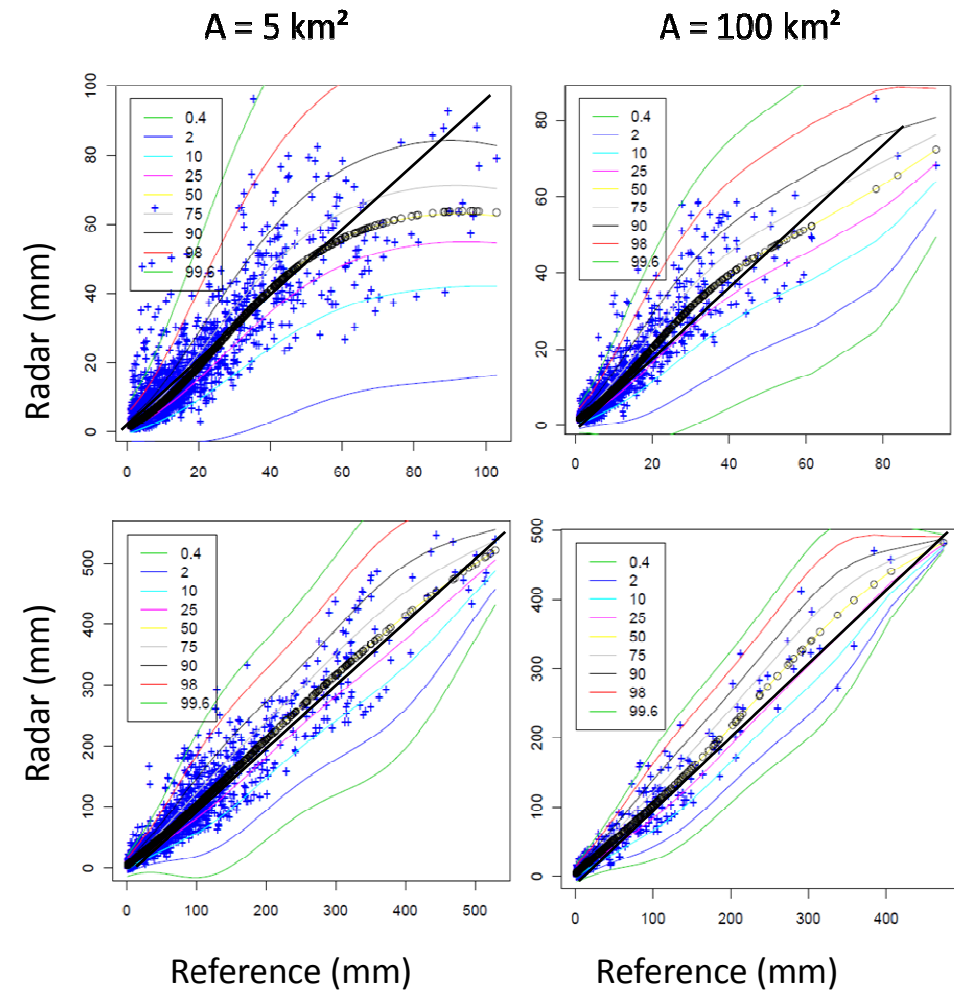
Estimation variance: $E(R_{AT}^{ref}(\underline{x}, t) - R_{AT}(\underline{x}, t))^2$



Assessment of the OHM-CV rainfall re-analyses

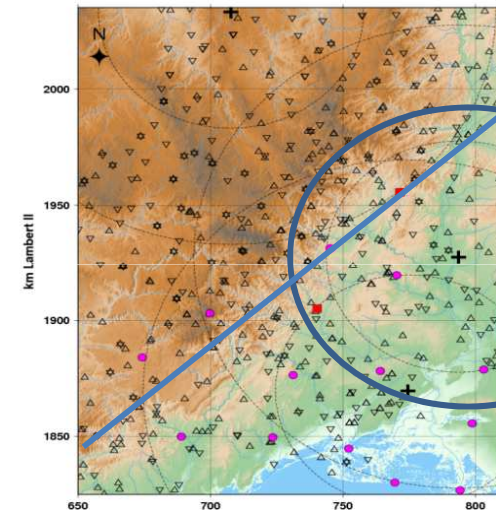
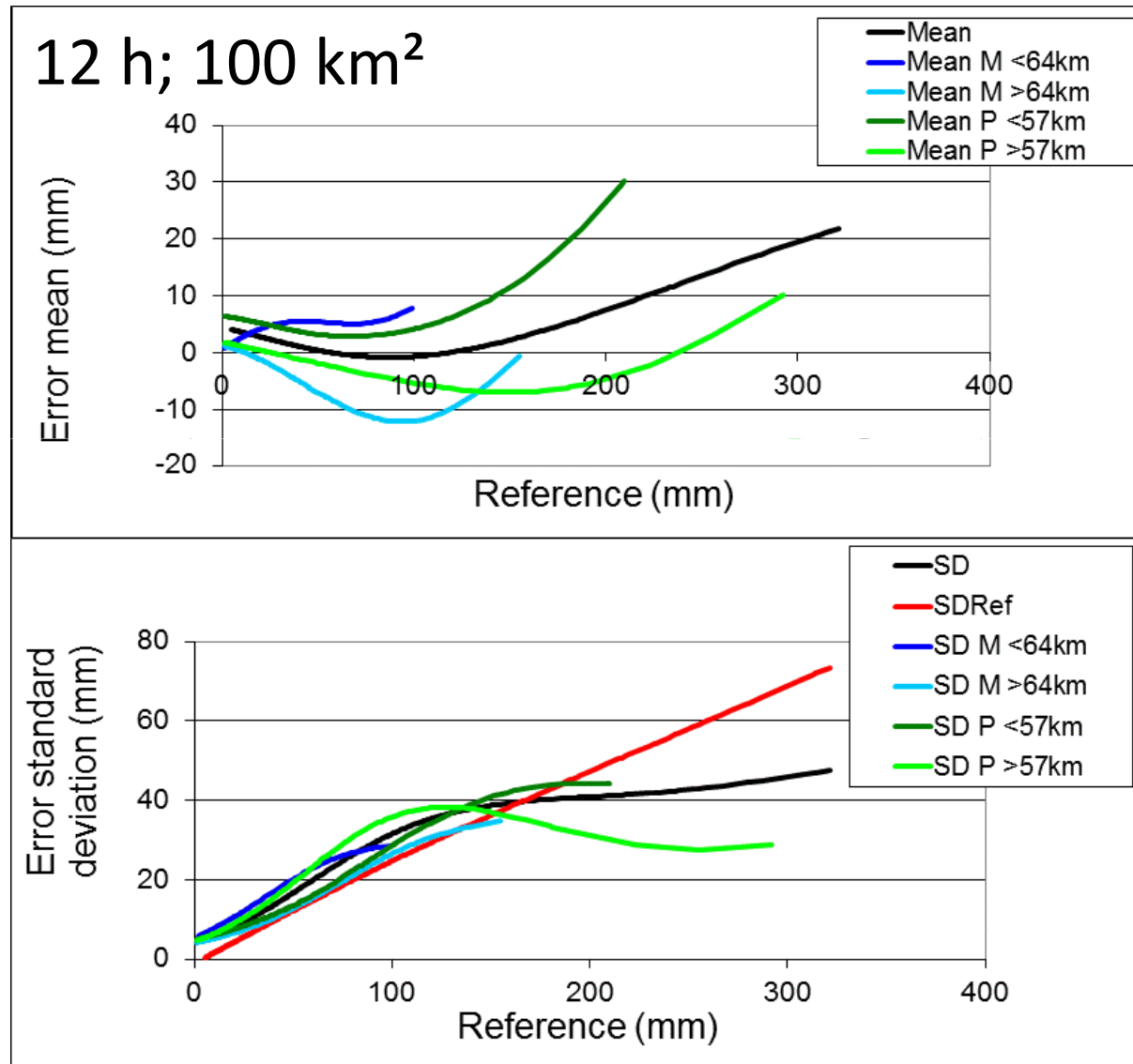
- Our best radar data processing:
4-radar mosaic/ GC / screening
and VPR...)
- Optimization of a single Z-R
relationship /event

$\Delta t = 1 \text{ hr}$



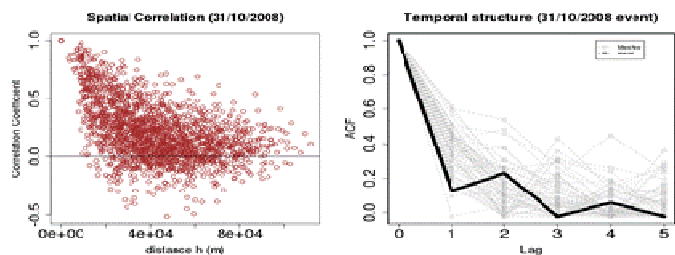
Conditional analyses
of the radar – reference rainfall
using the **GAMLSS framework**
(Rigby & Stanisopoulos 2004)

Conditional mean and standard deviation of the errors



Toward probabilistic radar QPEs

Spatial and temporal correlation

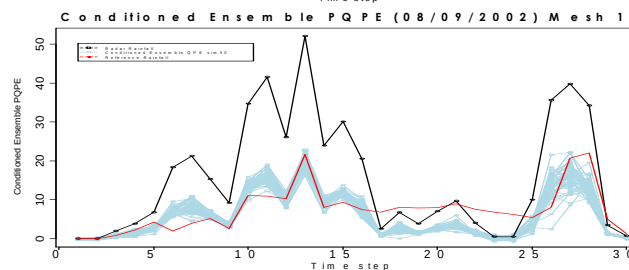
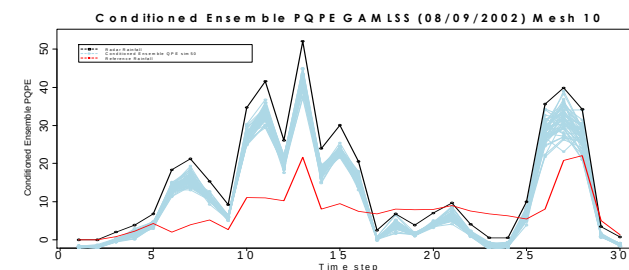
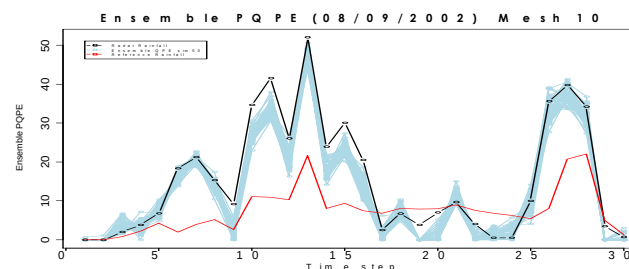


Conditioned by $\mu(\Delta)$
and $\sigma(\Delta)$ estimated globally
over the region of interest

Conditioned by
 $\mu(\Delta|R)$ and $\sigma(\Delta|R)$
estimated globally
over the region of interest

Conditioned by
 $\mu(\Delta|R)$ estimated
locally and $\sigma(\Delta|R)$
estimated globally

100 km², 1 hour



The radar error structure is complex

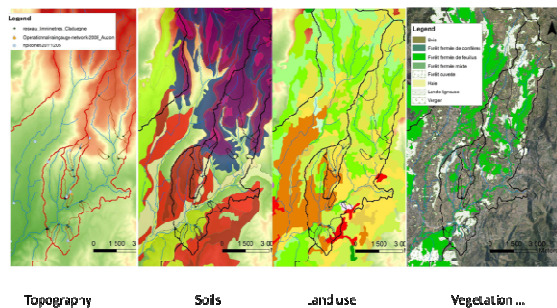
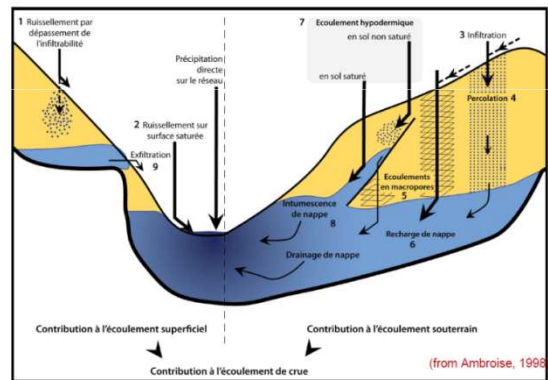
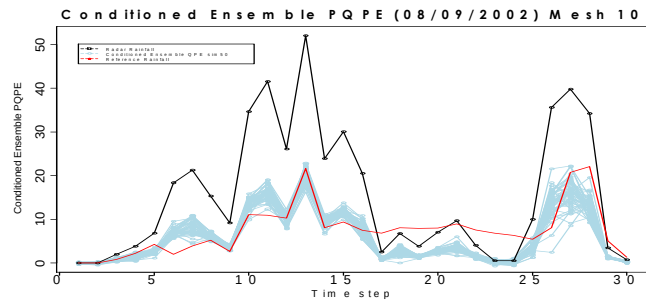
- rain-rate dependent
- space-time scales dependent
- location (range) dependent
- rain-type dependent

We are limited in our willing to provide radar error models useful in hydrology by the reference rainfall problem

- micronets, microwave links?
- work on estimation variance of merged products through geostatistical or Bayesian techniques; introduce some physics

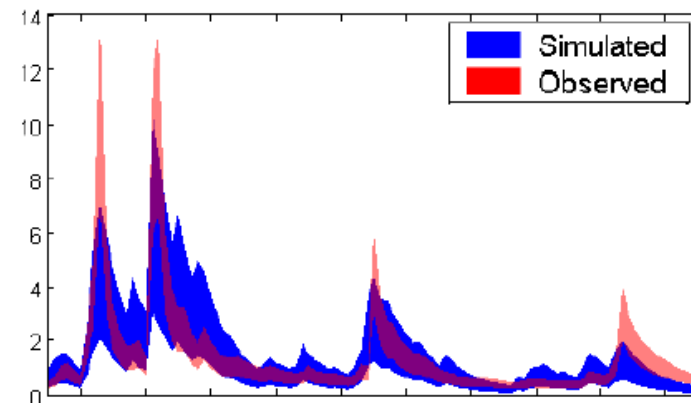
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Hydrological modelling:

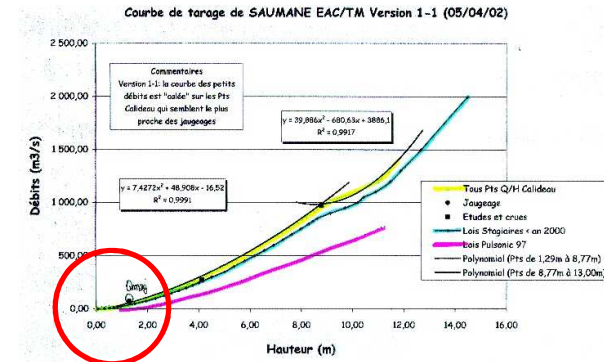
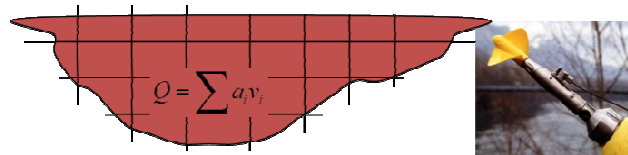
- Processes complexity
- Observability problem of the sub-surface
- Numerous landscape descriptors
- Spatial discretization
- processes representation (and selection?)
- Input-output uncertainties
- Structural and parametric uncertainties



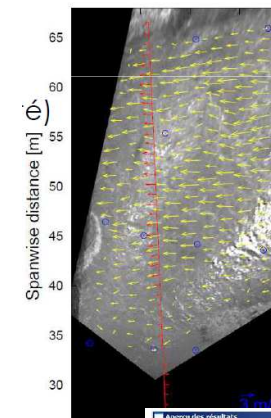
Monitoring of flooding rivers



Unfeasibility of conventional techniques during floods



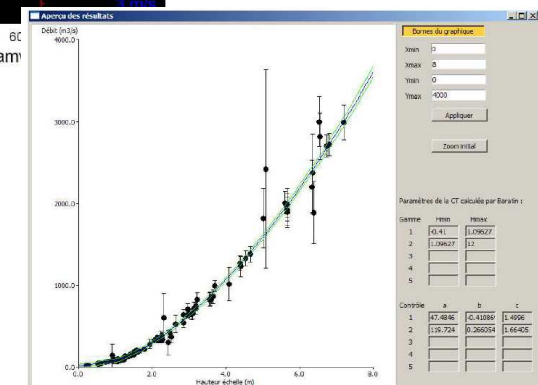
Need to develop « non-contact » gauging techniques



Need to improve rating curves (hydraulic modelling + gauging uncertainty accounting)

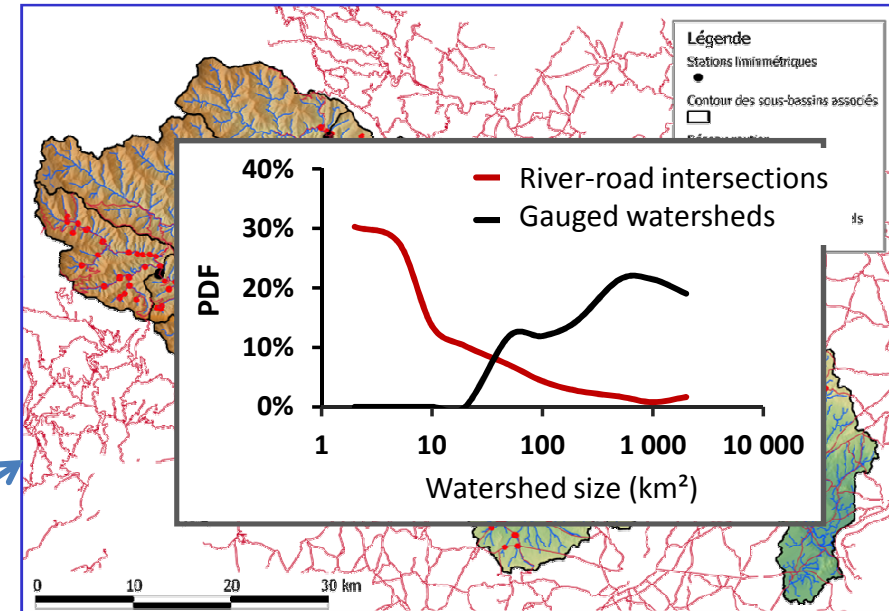
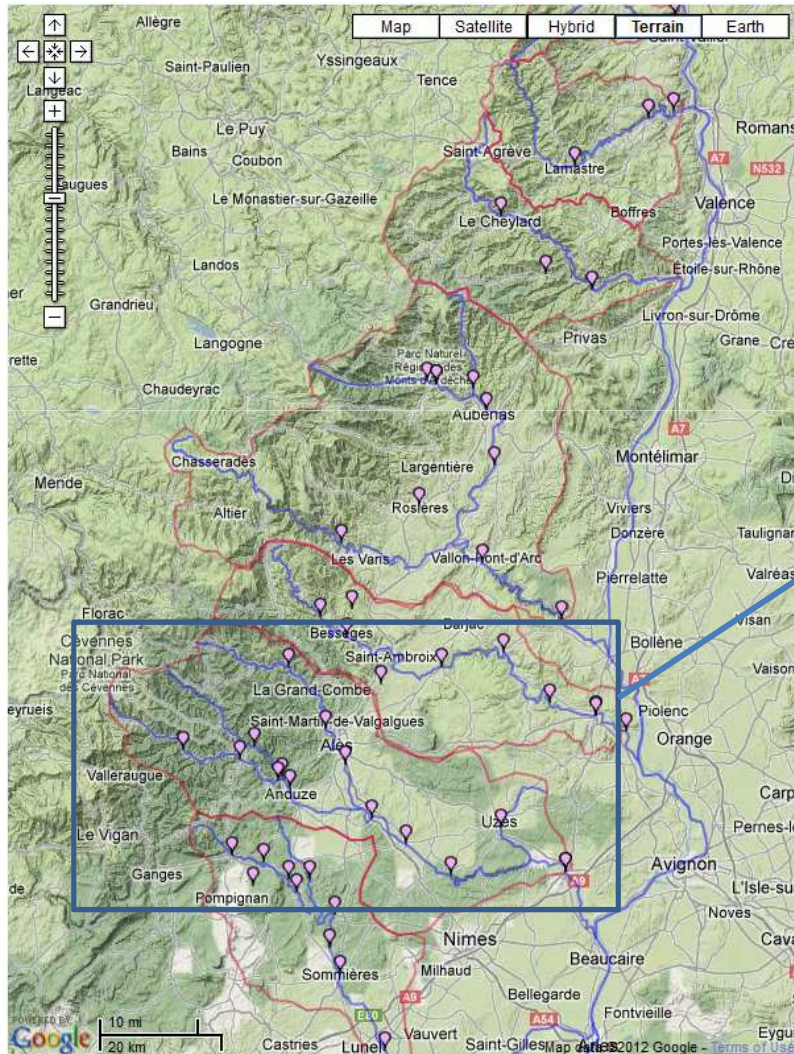


Le Coz et al.



Streamgauge network density

45 stream gauges in the OHM-CV



O. Vannier

- Post-event surveys
- A team/network of local observers?

Hydrological models, example 1: TOPMODEL

“topographic similarity”

Runoff production:

sub-surface lateral flow and saturation excess overland flow

Runoff transfer:

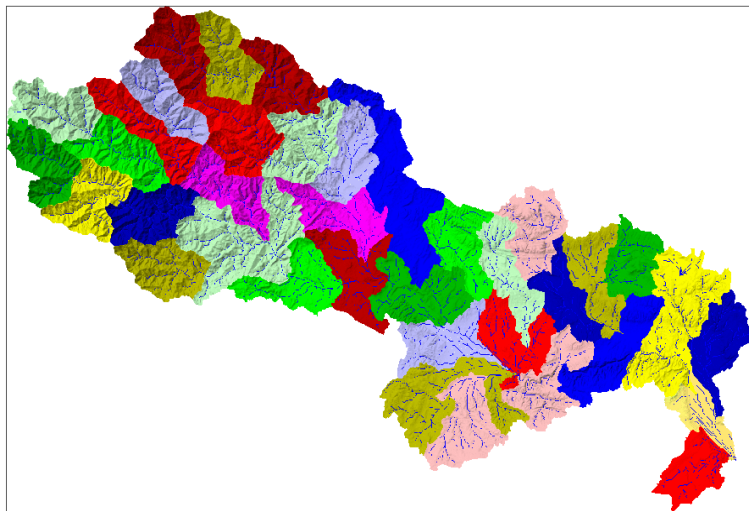
Geomorphological unit hydrograph

« Parcimony »: 4 parameters



1) nTopmodels

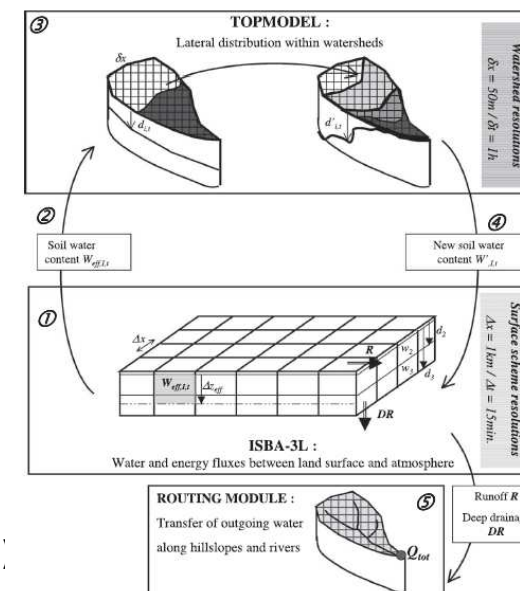
Hydrological meshes of ~constant area



Le Lay and Saulnier (2009)

2) ISBA-TOPMODEL

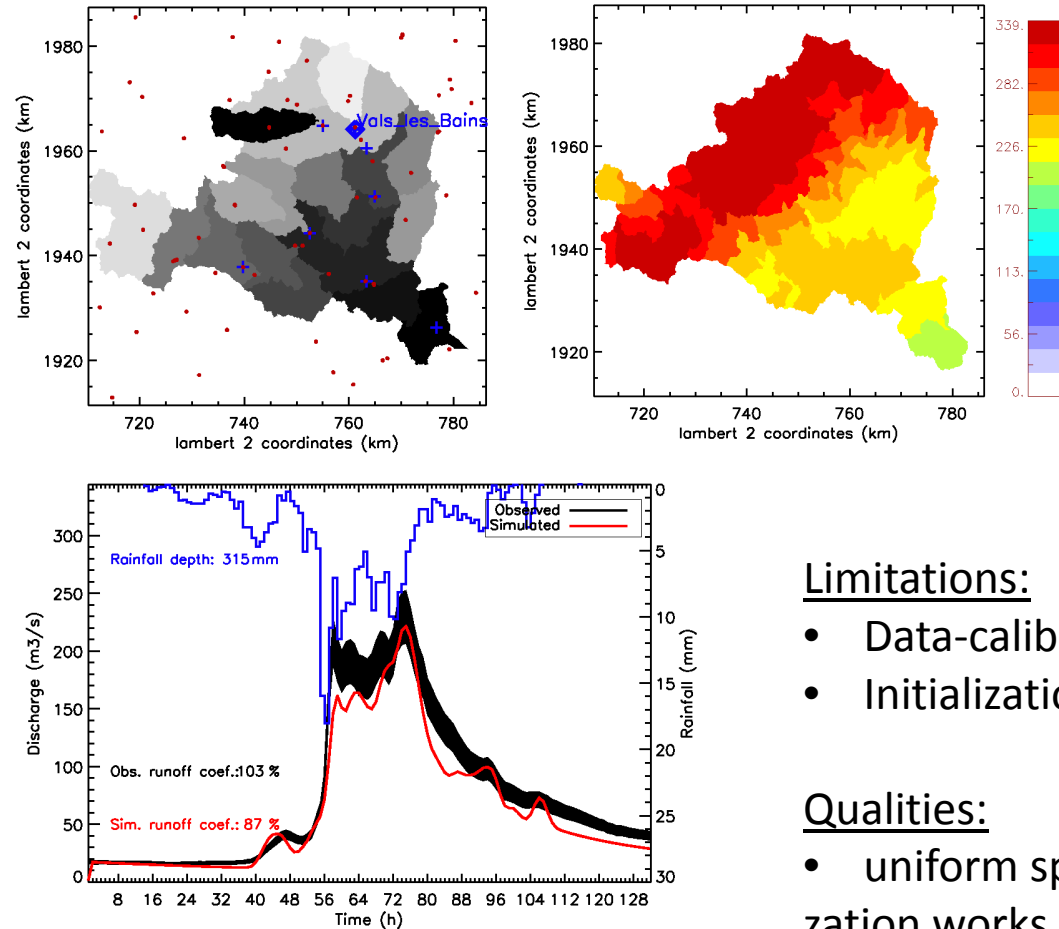
Coupled land surface – hydrological model



Vincendon et al. (2009)

Spatial assessment of n-Topmodels

The 2003 flood in the Ardèche catchment



Limitations:

- Data-calibrated
- Initialization?

Qualities:

- uniform spatial parameterization works reasonably well
- sensitivity to the rainfall spatial distribution

Hydrological models, example 2: CVN

“Landscape similarity”

Landscape discretization ($\sim 0.5 \text{ km}^2$)

Based on topography and soil maps



Spatialized forcings:

- Radar-raingauge rainfall
- Evapotranspiration (based on soil and vegetation maps)

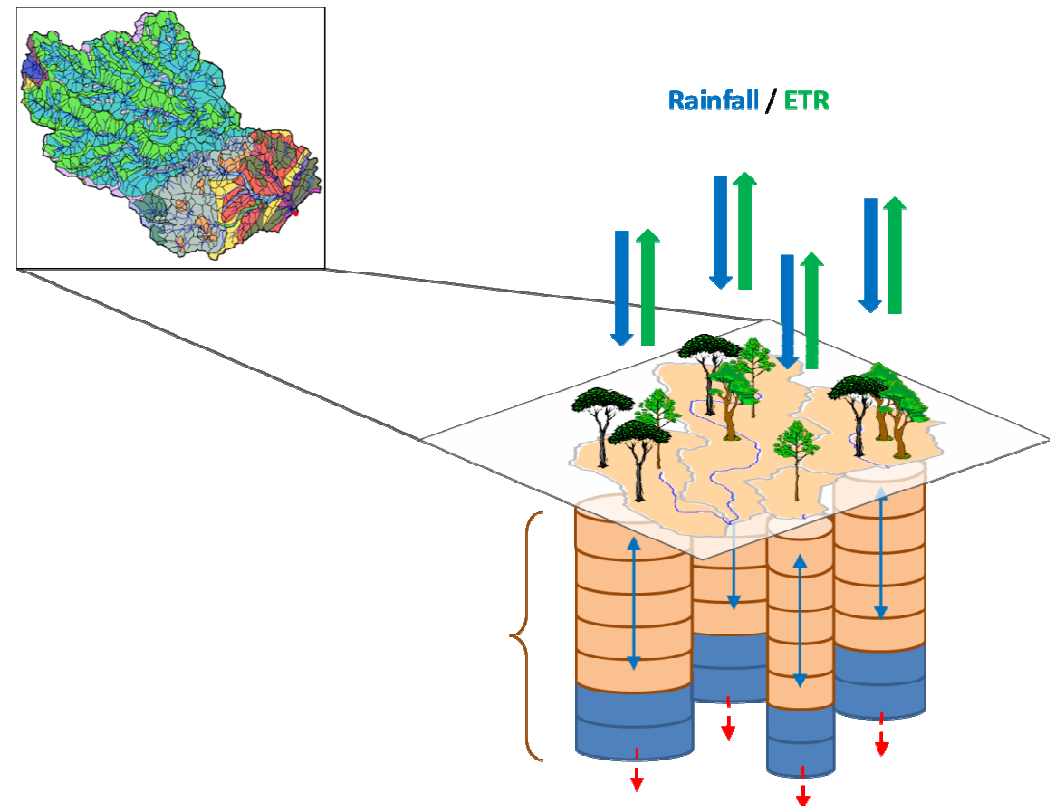
Runoff production:

Vertical transfer of water in the soil

- use Solve Richards 1D
- use $h(\Theta)$ et $K(\Theta)$ relations (Brooks & Corey) for the various layers

Runoff transfer:

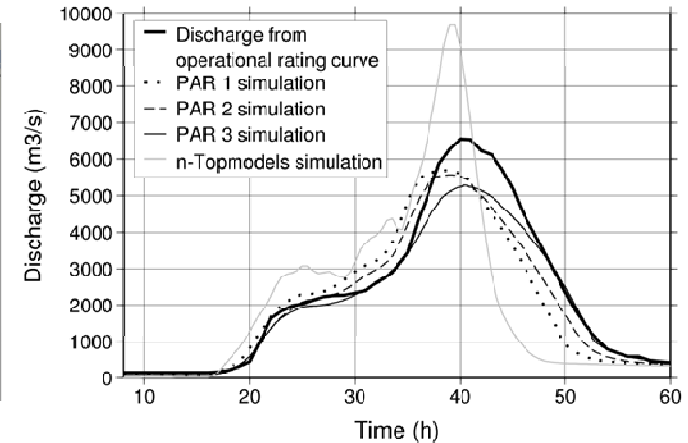
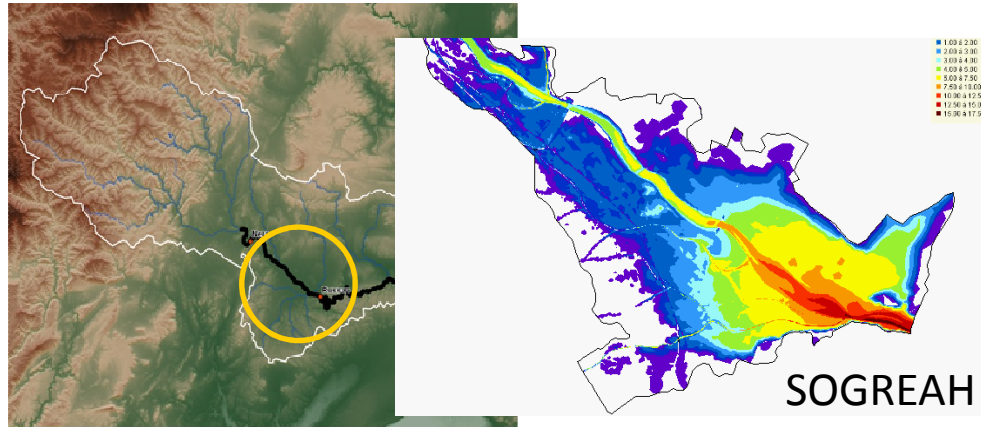
- Ponding sent directly to the river network
- Transfer with kinematic wave equation



Braud, Anquetin, Vannier et al.

Geomorphological controls:

Gardonnenque plain + Gardons gorges during the Gard 2002 flood

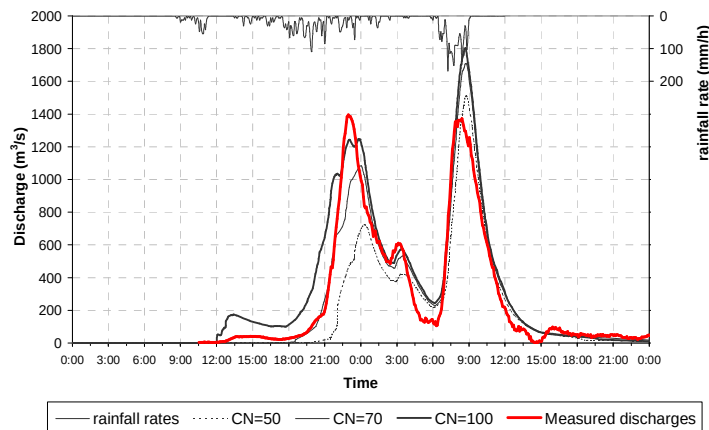


Bonnifait et al. (2009)

→ Coupling hydrologic and hydraulic models

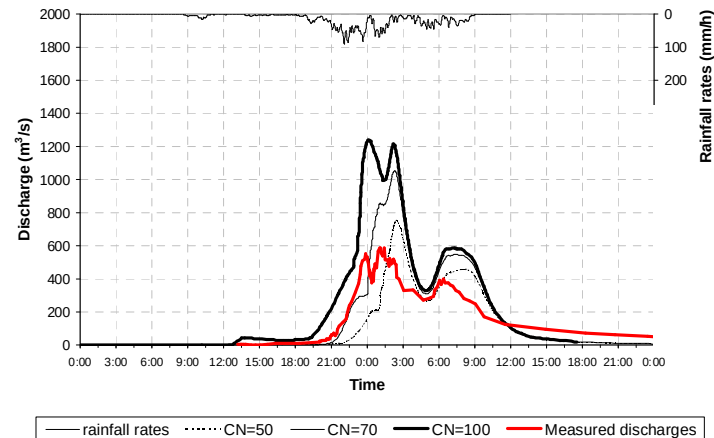
Geological controls:

Crieulon, Bge de la Rouvière (93 km²) :
Marls and non karstic limestones



→ Solution?

Vidourle, Bge de Conqueyrac (86 km²) :
Karst

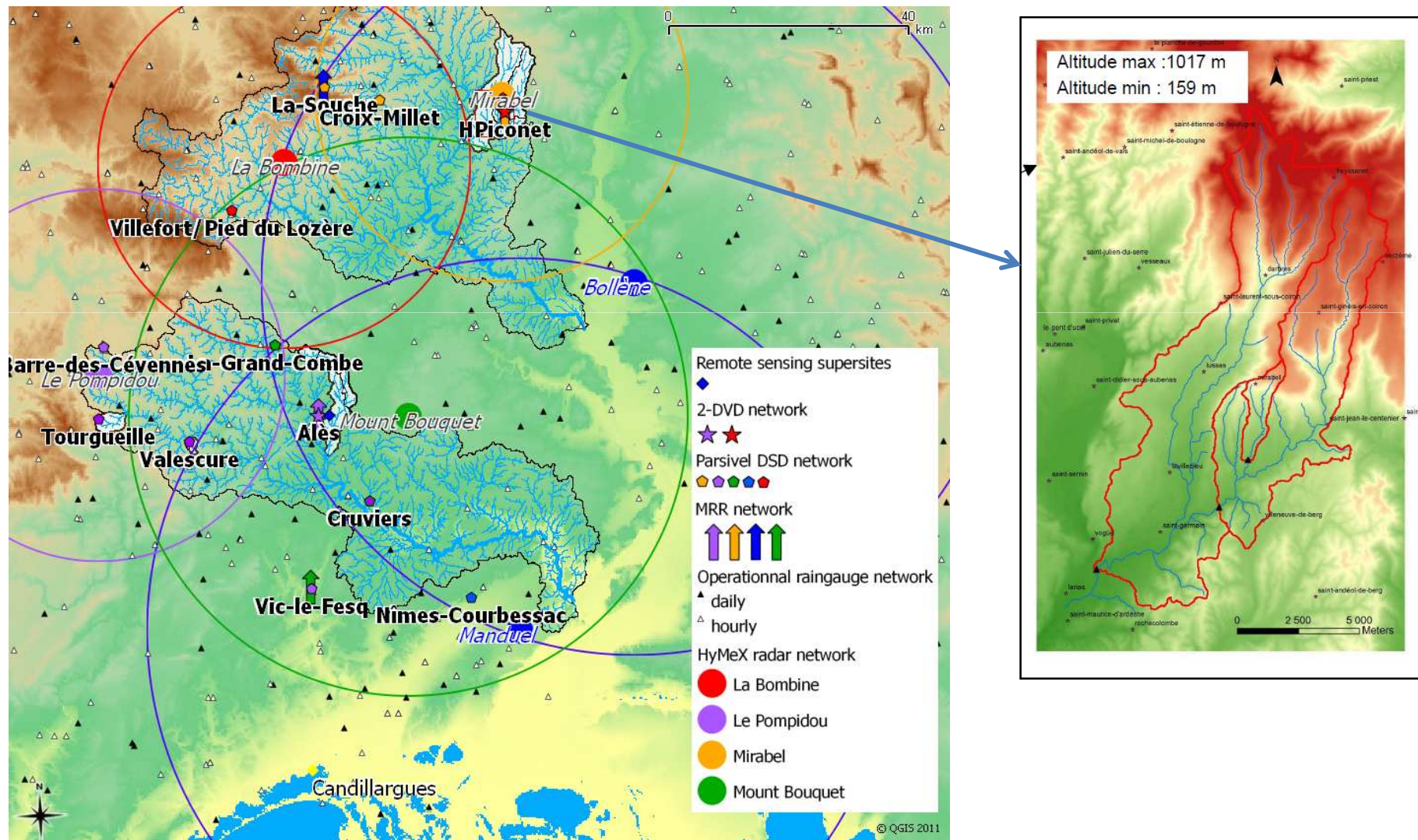


Gaume et al. (2004)

Coming work within HyMeX/FloodScale projects

Investigate the « missing scales »:

- Reinforce the rainfall observation system
- Study hydrological processes for a series of Cévennes landscapes



Thank you!