

Four-Dimensional Ensemble-Variational Data Assimilation 4DEnVar

Colloque National sur l'Assimilation de données

Andrew Lorenc, Toulouse France. 1-3 décembre 2014

© Crown copyright Met Office



4DEnVar: Topics in Talk

Design

Comparison with 4DVar

• Localisation & filtering of ensemble covariance

• Plans

Data Assimilation for Numerical Weather Prediction Met Office Design of a Modern System

Bayesian combination of observations and prior (background)

Incremental formulation – find best correction to prior

Cycled DA – each cycle assimilates a batch of observations spread over a time-window

Gaussian assumptions about Probability Distribution Functions

"Errors Of The Day" information in prior (background) PDF comes from an ensemble of forecasts.

Met Office Ours are incremental, with no outer-loop and no model error terms. Each gives a 4D best-fit to prior & observations in a 6 hour window. Use underline to denote 4D variables and operators:

Four-Dimensional Variational DA Methods

$$\mathbf{x}^b$$
background trajectory \mathbf{P} 4D error covariance of \mathbf{x}^b $\delta \mathbf{x}$ 4D analysis increment $\mathbf{y} = H(\mathbf{x}^b + \delta \mathbf{x})$ model estimate of obs $J(\delta \mathbf{x}) = \frac{1}{2}\delta \mathbf{x}^T \mathbf{P}^{-1}\delta \mathbf{x} + \frac{1}{2}(\mathbf{y} - \mathbf{y}^o)^T \mathbf{R}^{-1}(\mathbf{y} - \mathbf{y}^o)$ penalty function

4DEnVar uses a linear combination of perturbation trajectories $\delta \mathbf{x} = \Sigma_{k=1}^{N} \mathbf{\alpha}_{k} \circ \mathbf{x}'_{k}$ hybrid \Rightarrow a combination of climatological and localized ensemble covariances.

4DVar uses linear model $\delta \mathbf{x} = \mathbf{M} \delta \mathbf{x}_0$



Possible Methods

Met Office

• [Hybrid / Ensemble] - 4DVar Use ensemble to augment 3D covariance at beginning of window, then use linear & adjoint model for time-dimension.

• [Hybrid] - 4DEnVar

Use ensemble trajectories to determine 4D covariance directly.

• EnKF – e.g. 4D-LETKF

Localised Ensemble Transform Kalman Filter – transforms background ensemble perturbations to sample the analysed PDF from Kalman eqn.

All need to use [localisation / smoothing / hybridisation] to get usable covariances from a small ensemble!!



4DEnVar: using an ensemble of 4D trajectories which samples background errors

Ensemble trajectory matrix $\mathbf{X} = \begin{bmatrix} \mathbf{x}'_1 \cdots \mathbf{x}'_N \end{bmatrix}$ where $\mathbf{x}'_k = \frac{1}{\sqrt{N-1}} (\mathbf{x}_k - \bar{\mathbf{x}})$ Model 4D \mathbf{P} directly, $\mathbf{P} = \mathbf{C} \circ \mathbf{X} \mathbf{X}^T$

 $\underline{\mathbf{C}} = \underline{\mathbf{U}}^{\alpha} \underline{\mathbf{U}}^{\alpha T}$

$$\boldsymbol{\alpha}_{k} = \boldsymbol{\underline{U}}^{\alpha} \mathbf{v}_{k}^{\alpha}$$
$$\delta \boldsymbol{\underline{x}} = \boldsymbol{\Sigma}_{k=1}^{N} \boldsymbol{\underline{\alpha}}_{k} \circ \boldsymbol{\underline{x}}_{k}'$$
$$\mathbf{v}^{T} = \left[\mathbf{v}_{1}^{\alpha T} \cdots \mathbf{v}_{N}^{\alpha T} \right]$$
$$J(\mathbf{v}) = \frac{1}{2} \mathbf{v}^{T} \mathbf{v} + \frac{1}{2} \left(\boldsymbol{\underline{y}} - \boldsymbol{\underline{y}}^{o} \right)^{T} \boldsymbol{\underline{R}}^{-1} \left(\boldsymbol{\underline{y}} - \boldsymbol{\underline{y}}^{o} \right)$$

Model 4D $\underline{\mathbf{P}}$ directly, as localised ensemble covariance,

then model $\underline{\mathbf{C}}$ using transforms

4D localised linear combination of ensemble trajectories

concatenated control vectors

Transformed penalty function

4DEnVar: using an ensemble of 4D trajectories which samples background errors Extra Details

model $\underline{\mathbf{C}}$ using transforms

It is common to use a 3D \mathbf{C}

and persistence in time: \mathbf{I}

4D localised linear combination of ensemble trajectories

can be built from 3D localised perturbations and constant α_k .

Matrix notation: Buehner (2005) avoids \circ using $diag(\mathbf{x})$ for the diagonal matrix: $diag(\mathbf{x})_{i,i} = \mathbf{x}_i$

 $\underline{\mathbf{C}} = \underline{\mathbf{U}}^{\alpha} \underline{\mathbf{U}}^{\alpha T}$

 $\mathbf{C} = \mathbf{U}^{\alpha} \mathbf{U}^{\alpha T}$ $\underline{\mathbf{C}} = \underline{\mathbf{I}} \mathbf{C} \mathbf{I}^{T}$

$$\boldsymbol{\alpha}_{k} = \mathbf{U}^{\alpha} \mathbf{v}_{k}^{\alpha}$$
$$\delta \mathbf{x} (t) = \sum_{k=1}^{N} \boldsymbol{\alpha}_{k} \circ \mathbf{x}_{k}^{\prime} (t)$$

$$\delta \mathbf{x}\left(t\right) = \sum\nolimits_{k=1}^{N} diag\left(\mathbf{x}_{k}^{\prime}\left(t\right)\right) \boldsymbol{\alpha}_{k}$$



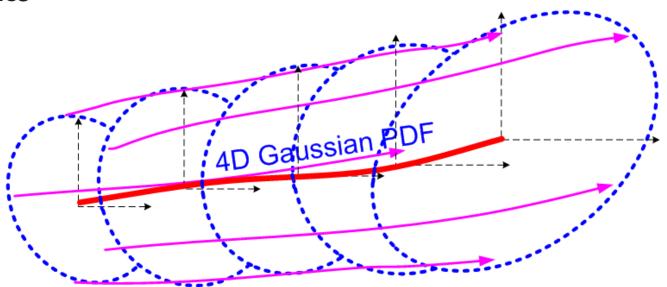
$$\delta \mathbf{x} = \sum_{k=1}^{N} \mathbf{\alpha}_k \circ \mathbf{x}'_k$$

$$\delta \mathbf{\underline{x}} = \sum_{k=1}^{N} \mathbf{\underline{\alpha}}_k \circ \mathbf{\underline{x}}'_k$$



Incremental 4D-Ensemble-Var

Met Office



Trajectories of perturbations from ensemble mean

Ensemble mean trajectory

Localised trajectories define 4D PDF of possible increments

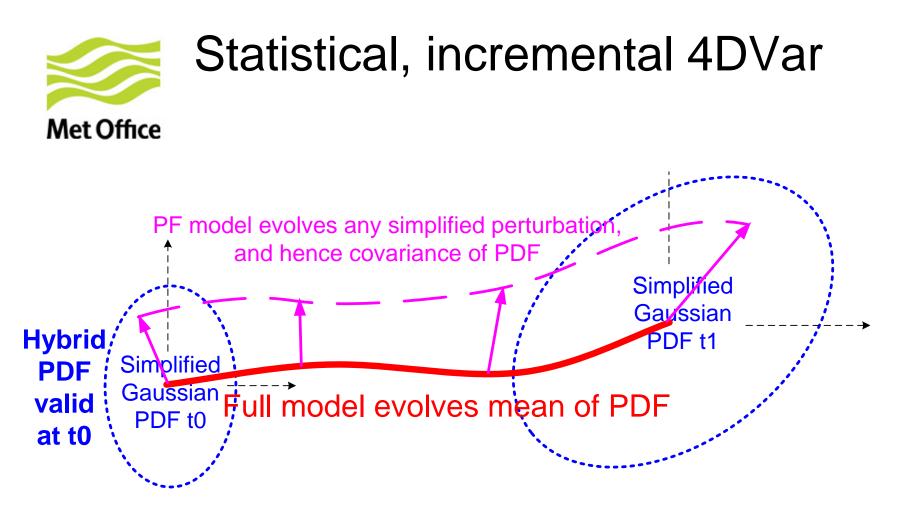
Statistical 4D-Var approximates entire PDF by a Gaussian.

4D analysis is a (localised) linear combination of nonlinear trajectories. It is not itself a trajectory.



Comparison with 4DVar

© Crown copyright Met Office



Statistical 4DVar approximates entire PDF by a 4D Gaussian defined by PF model.

4D analysis increment is a trajectory of the PF model.

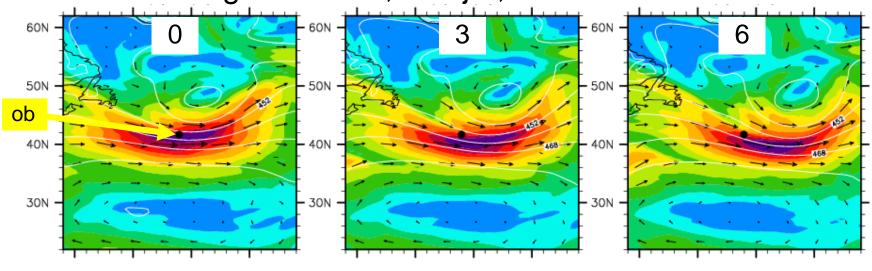
Lorenc & Payne 2007



En-4DVar analysis increment $\delta \mathbf{x} = \mathbf{M} \Sigma_{k=1}^{N} \boldsymbol{\alpha}_{k} \circ \mathbf{x}_{k}'$

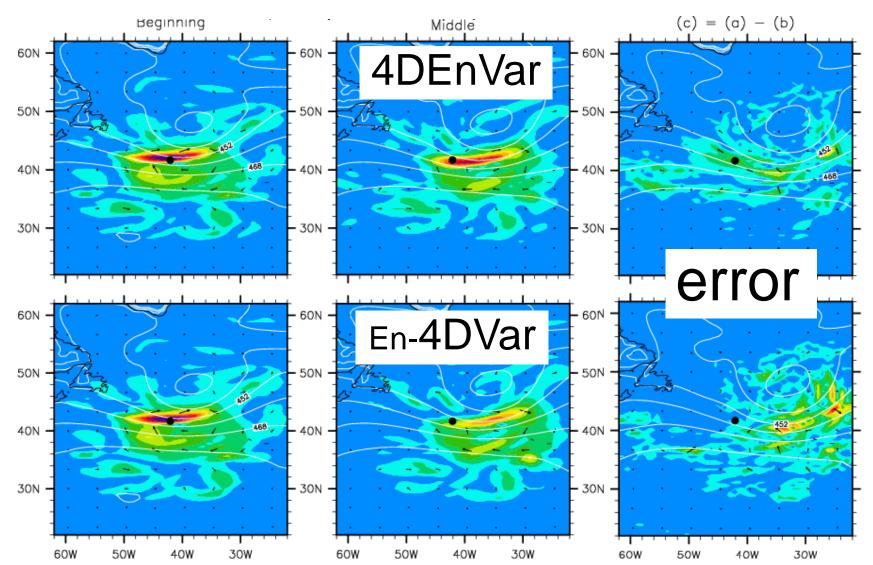
4DEnVar analysis increment $\delta \mathbf{x} = \Sigma_{k=1}^{N} \boldsymbol{\alpha}_{k} \circ \mathbf{x}_{k}^{\prime}$

Test with a single wind ob, in a jet, at the start of the window





100% ensemble 1200km localization scale





hybrid-4DEnVar

4D analysis increment
$$\delta \mathbf{x} = \beta_c \mathbf{I} \delta \mathbf{x}_0 + \beta_e \Sigma_{k=1}^N \boldsymbol{\alpha}_k \circ \mathbf{x}'_k$$

Localized 4D covariance $\mathbf{P} = \beta_c^2 \mathbf{IBI}^T + \beta_e^2 \mathbf{C} \circ \mathbf{X} \mathbf{X}^T$

hybrid-4DVar

4D analysis increment $\delta \mathbf{x} = \mathbf{M} \left(\beta_c \delta \mathbf{x}_0 + \beta_e \Sigma_{k=1}^N \boldsymbol{\alpha}_k \circ \mathbf{x}'_k \right)$

Localized 4D covariance $\mathbf{P} = \mathbf{M} \left(\beta_c^2 \mathbf{B} + \beta_e^2 \mathbf{C} \circ \mathbf{X} \mathbf{X}^T \right) \mathbf{M}^T$

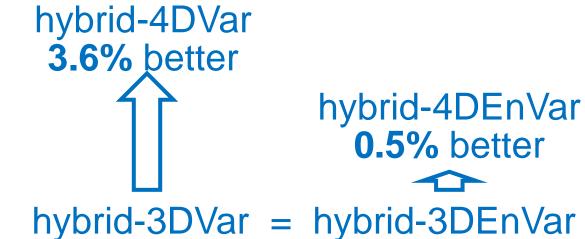


Met Office trial of 4DEnVar Lorenc et al. (2014)

Our first trial copied settings from the hybrid-4DVar:

- C with localisation scale 1200km,
- hybrid weights β_c^2 =0.8, β_e^2 =0.5

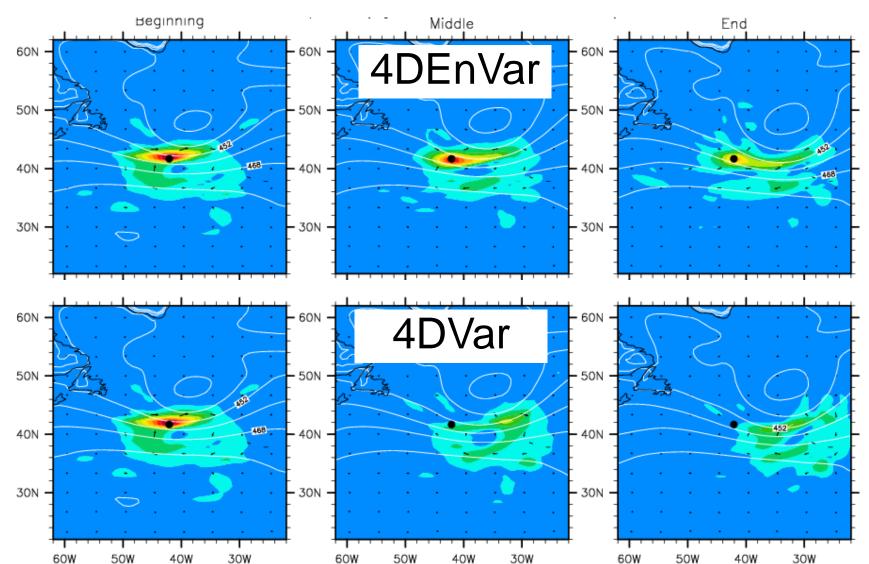
Results were disappointing:



The reason was the large weight given to the climatological covariance, which is treated like 3DVar in 4DEnVar



50-50% hybrid 1200km localization scale





Relative "Strong Constraint Errors"

We ran similar tests on a Hurricane Sandy case.

Here the ensemble covariances dominated, making hybrid-4DEnVar perform better.

	Jet case	Hurricane
1200km localization scale		Sandy
4DEnVar	<mark>(51%</mark>)	<mark>57%</mark>)
En-4DVar	54%	69%
Hybrid-4DEnVar	78%	(<mark>66%</mark>)
Hybrid-4DVar	66%	75%

When the ensemble covariances dominated the increments, and the horizontal localisation was not too severe, 4DEnVar had better consistency with the strong constraint than 4DVar. Runs with smallest deviation from model constraint

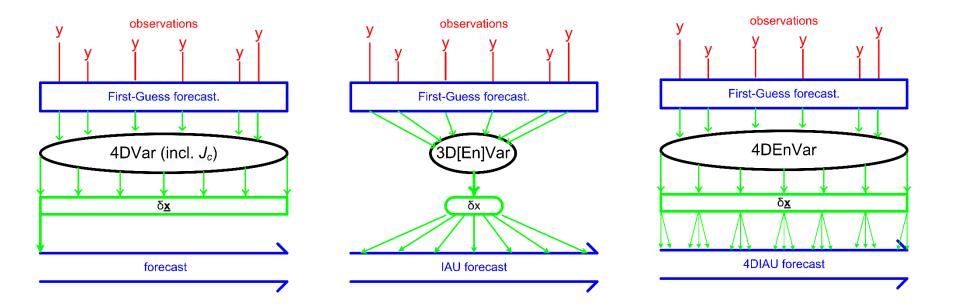


Conclusions from 4D analysis increment study

- The main error in our hybrid-4DEnVar (v hybrid-4DVar) is that the climatological covariance is used as in 3DVar.
- 2. 3D localisation not following the flow is not an important error for our 1200km localisation scale and 6hour window, but does become important for a 500km scale.

Initialization





4DVar's $J_c = \frac{1}{2} \left(\mathbf{F} \delta \underline{\mathbf{x}} \right)^T \mathbf{G}^{-1} \left(\mathbf{F} \delta \underline{\mathbf{x}} \right)$ produces balanced increments by penalizing gravity waves.

IAU applied a related to model.

4DIAU has less time-filter (Polavarapu *et al.*, time-filtering, but is effective at 2004) while adding increments cancelling noise in the increment trajectory.



4DIAU – Initialisation

Lorenc *et al.* (2014) studied the simple 4DIAU method – it works effectively for 4DEnVar.

4DIAU is used in operational 4DEnVar in

- EC (Canada) Buehner et al. (2015),
- NCEP (USA) (Kleist, personal communication).

The only problem is that it rules out an outer-loop.



The maintenance and running costs of hybrid-4DVar are larger, so there is an incentive to improve hybrid-4DEnVar.

We need to reduce the weight on climatological **B** relative to the ensemble covariance.

We must first improve the ensemble covariances:

- a bigger ensemble;
- better ensemble generation;
- better filtering of ensemble covariance, e.g. localization.

Encouraging progress has been made in all of these.



Localisation

© Crown copyright Met Office

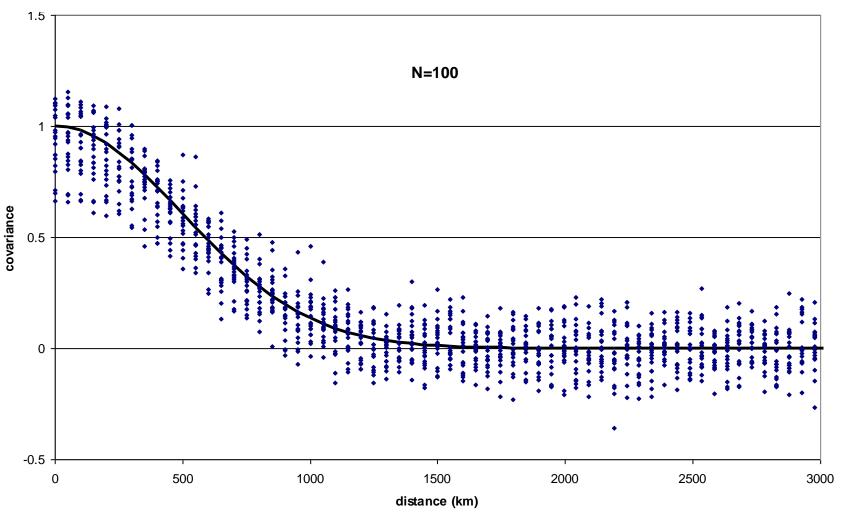


Ensemble covariance filtering

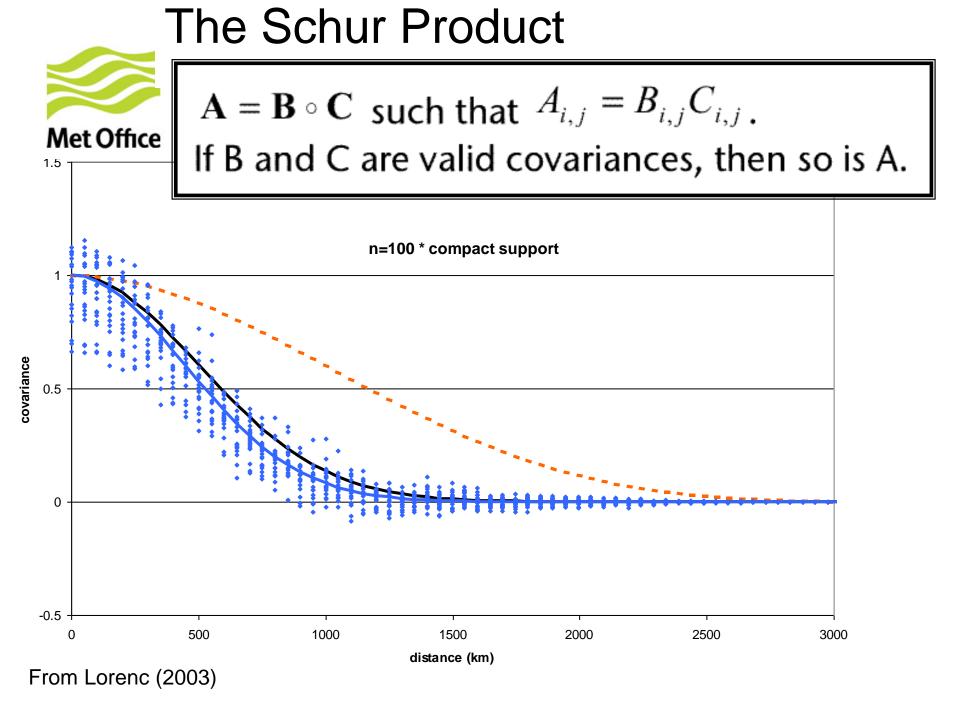
Covariance \mathbf{B} is big! We need a large ensemble PLUS clever filtering to reduce sampling noise, based on 2 ideas:

- Assume some correlations are near zero, & localise: horizontal, vertical, spectral, transformed variables;
- Assume local homogeneity apply smoothing: horizontal, rotational, and time.





From Lorenc (2003)





Localisation – time-dimension

3D: $\mathbf{P} = \mathbf{C} \circ \mathbf{X}\mathbf{X}^T$ 4D: $\underline{\mathbf{P}} = \underline{\mathbf{C}} \circ \underline{\mathbf{X}}\mathbf{X}^T$

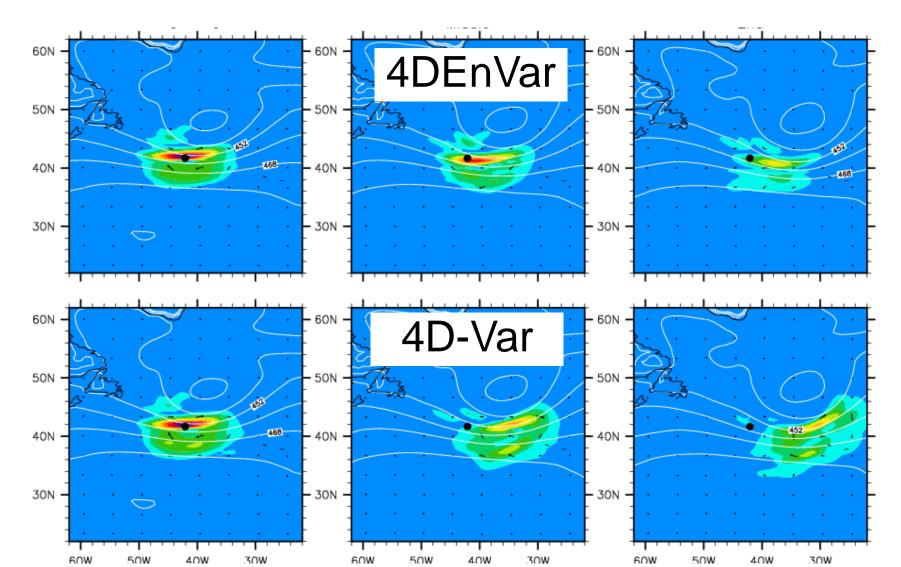
 \mathbf{C} is a correlation matrix, near 0 where the real covariance is small, otherwise near 1. Nearly all 4DEnVar implementations have no time-localisation, so $\mathbf{C} = \mathbf{I}\mathbf{C}\mathbf{I}^{T}$.

• Research ideas for time-varying localisation exist,

e.g. Bishop and Hodyss (2011); Ota et al. (2013).

- The need for flow-following localisation depends on the time-window.
- *WARNING*: Time-localisation might give optimal estimate at one time, but NOT an optimal ongoing filter!

Met Office 100% ensemble 500km localization scale





Localisation – transformed variables

Any linear transform of \mathbf{X} gives an equally valid covariance. So we can equally well localise using $\mathbf{P} = \mathbf{T}^{-1} \left(\mathbf{C} \circ \mathbf{T} \mathbf{X} \mathbf{X}^T \mathbf{T}^T \right) \mathbf{T}^{-T}$.

The Met Office EnVar localises transformed variables:

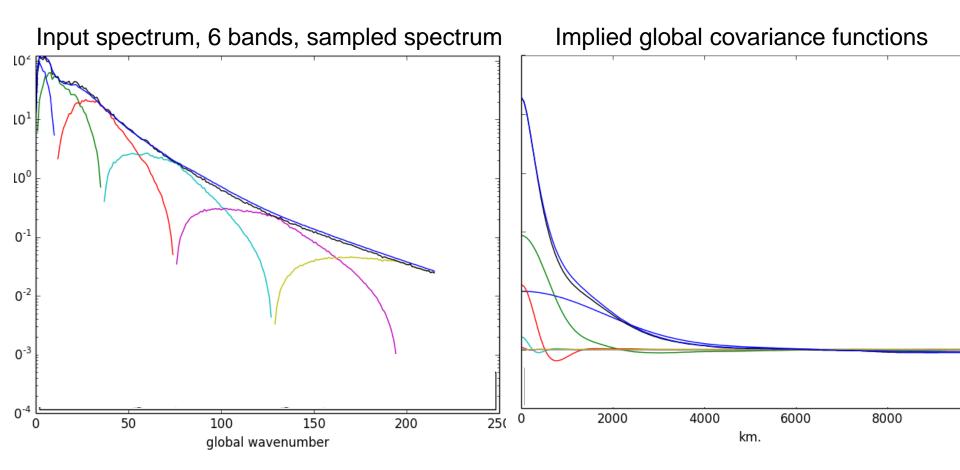
- this avoids the imbalance normally caused by localisation.

There is the option of localising between transformed variables - this imposes balance relationships implicit in the transforms.

Buehner and Charron (2007); Buehner (2012) suggested spectral localisation;

- this has the effect of smoothing in space;
- the wavebands allow different horizontal localisations for each scale.

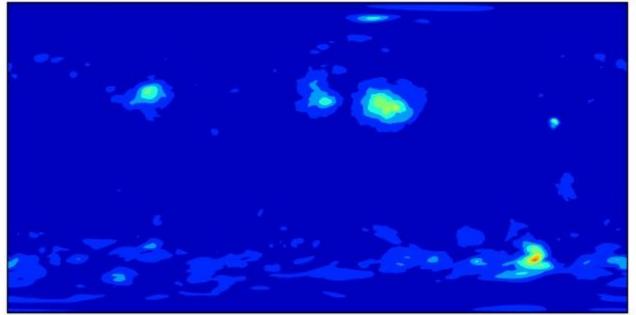
Spectral [waveband] localisation, Met Office of pressure at 4km

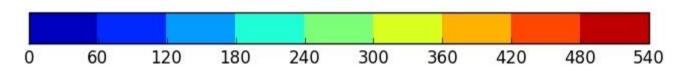


Ensemble spread in pressure at level 24 (~4km)



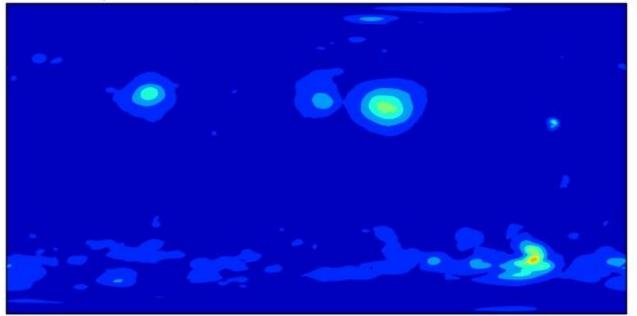
covsampleen sampled background sigma in exptac632eTR200 p24 3996m, min=5.29 mean= 35.4 max= 436

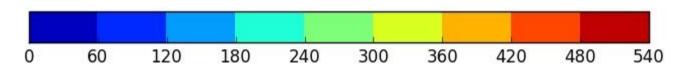




Spread, after localisation using 6 wavebands (~4km)









Localisation – different implementations

The EnVar schemes covered so far use an " α control variable" (Lorenc, 2003) localisation to compensate for a small ensemble.

Another approach is to make extra members using eigenvectors of \mathbf{C} :

- This was used in early 4DEnVar (Liu et al., 2009)
- and in spectral localisation (Buehner, 2012).

At what stage we localise makes a big difference:

- at start of time-window, before model En-4DVar (Clayton et al., 2013);
- EnKF schemes typically localise in ob-space after H; this is not as good (Campbell *et al.*, 2010) and cannot handle wavebands;
- the LETKF (Hunt *et al.*, 2007; Harlim and Hunt, 2007) localises by selecting obs outside the solution of the KF equation.



© Crown copyright Met Office

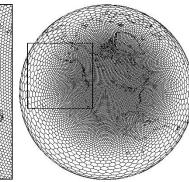


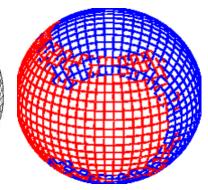
Hybrid-4DEnVar	Hybrid-4DVar
4D via use of ensemble trajectories. Only 3D use of climatological B .	4D via use of <u>M</u> & <u>M</u> ⁷ .
No forecast model inside algorithm, easy to add variables.	Needs special software for <u>M</u> & <u>M</u> ⁷ , effort needed to add variables.
Needs memory to store trajectories.	<u>M</u> & <u>M</u> ^{T} have to be run each iteration.
Needs large ensemble and good localisation.	Needs good covariance model.

- The Met Office's operational hybrid-4DVar works well on current & next computer. But <u>M</u> & <u>M</u>^T are struggling with model changes and resolution increases.
- hybrid-4DEnVar will be the first method coded for use with the planned GungHo model on the massively parallel computer expected next decade.



Parallelisation





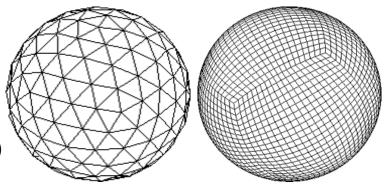
4DVar has several potential problems looming in the next decade - their timing for each centres will depend on their computers and models:

1. Need new design to use millions of parallel threads, especially in sequential runs of linear (PF) and Adjoint models.

2. Forecast models are being redesigned to address this – a maintenance issue for the PF and Adjoint models.

4DEnVar is a simple solution, using the ensemble trajectories, pre-calculated in parallel, instead of the models inside 4DVar.

If Fourier filters and Poisson solvers are not available then the LETKF is an easier approach.







Hybrid-4DEnVar	EnKF
Needs transforms and filters of model fields, each iteration.	Works locally, using ensemble predictions of observations $H(\underline{\mathbf{x}}^{b}_{k})$.
	Easy to build & cheap to run; cost small compared to ensemble forecasts.
Using wavebands, better for analysing a wide range of scales, including large scales from global ensemble. Can use hybrid with climatological B .	

- We cannot yet run a big enough ensemble of good enough forecasts to effectively sample background error covariance. Priority is on improving model & computer.
- Meanwhile Met Office is running 3DVar and developing 4DVar, with a small forecast ensemble via downscaling of the global ensemble.
- We will start experimenting with En-DA with a simple EnKF (concentrating on model forecasts and infrastructure). 4DEnVar is not yet planned.



How to create the ensemble?

• From a separate EnKF system. E.g.:

Canada uses an independent EnKF (Houtekamer et al. 2014)

➢Met Office MOGREPS (Bowler et al., 2008; Flowerdew and Bowler, 2011) uses a Localised ETKF re-centred on a deterministic analysis.

• From an Ensemble of 4DEnVar (EDA: Bonavita et al., 2012)

We are experimenting with this, including a MeanPert algorithm to reduce cost.

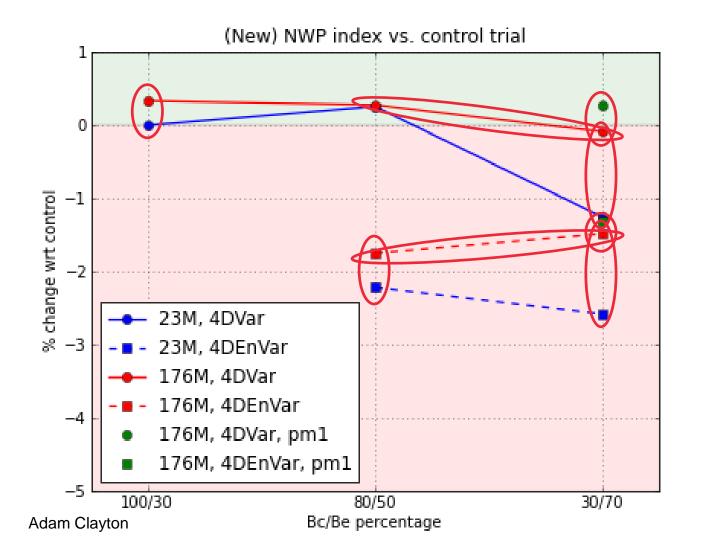
From a single EnVar's Hessian information: EVIL (Auligné, 2012)

• Downscaling from a global ensemble

>Met Office experience with regional MOGREPS \Rightarrow this is key process >Canada's regional 4DEnVar uses the global ensemble (Caron et al., 2015)



Trials of increased ensemble size and weight



Modest improvement when increasing ensemble size

Much larger improvement when ensemble weight is high

4DVar performs worse with high ensemble weight, 4DEnVar performs better

Using ensemble modes from the wrong time brings a small benefit



4DEnVar: Summary of Talk

• Design

4D using ensemble \Rightarrow hybrid-4DVar / hybrid-4DEnVar / EnKF

Comparison with 4DVar

4DEnVar time propagation OK except in hybrid.

• Localisation & filtering of ensemble covariance

Space, transformed, spectral (wavebands), time, how.

Plans

Method (global & convective), MPP, getting ensemble, size



Questions and answers

© Crown copyright Met Office

LETKF solves KF equations separately for each grid-point, with local observations.

Use the matrix of ensemble model-ob perturbations calculated using nonlinear H:

ETKF for mean analysis

The ensemble-space matrix inversion is solved directly

SQRT-filter for the analysis pertubations

Each point's ${f w}$ is a row of matrix whose columns are ${f lpha}$

 $\underline{\mathbf{y}}_{k}^{\prime} = \frac{1}{\sqrt{N-1}} \left(\underline{H} \left(\underline{\mathbf{x}}_{k}^{b} \right) - \overline{\underline{H} \left(\underline{\mathbf{x}}^{b} \right)} \right)$ $\underline{\mathbf{Y}}^b = \left[\begin{array}{ccc} \underline{\mathbf{y}}_1' & \cdots & \mathbf{y}_N' \end{array}
ight]$ $\delta \mathbf{x} = \mathbf{X}^b \mathbf{w}$ $\mathbf{w} = \tilde{\mathbf{P}}^{a} \left(\underline{\mathbf{Y}}^{b} \right)^{T} \mathbf{R}^{-1} \left(\mathbf{y}^{o} - \underline{H} \left(\underline{\mathbf{x}}^{b} \right) \right)$ $\widetilde{\mathbf{P}}^{a} = \left[\mathbf{I} + \left(\underline{\mathbf{Y}}^{b}
ight)^{T}\mathbf{R}^{-1}\underline{\mathbf{Y}}^{b}
ight]^{-1}$ $\underline{\mathbf{X}}^{a}=\left(ilde{\mathbf{P}}^{a}
ight)^{1/2}\underline{\mathbf{X}}^{b}$ $\mathbf{A} = \left[oldsymbol{lpha}_1' \ \cdots \ oldsymbol{lpha}_N'
ight] arpropto \left[egin{array}{c} \mathbf{w}_1^T \ arpropto \ arpropto \ \mathbf{w}_N' \end{array}
ight] arpropto$

Hunt *et al.* (2007); Harlim and Hunt (2007) apply the factor $1/\sqrt{N-1}$ to $\mathbf{w} \& \boldsymbol{\alpha}$ rather than to \mathbf{X}^{b} .

References

- Auligné T. 2012. An integrated ensemble/variational hybrid data assimilation. Presented at International Conference on Ensemble Methods in Geophysical Sciences Toulouse, France. 12-16 November 2012.
- Bishop CH, Hodyss D. 2011. Adaptive ensemble covariance localization in ensemble 4D-Var state estimation. Mon. Weather Rev. 139(4): 1241-1255.
- Bonavita M, Isaksen L, Holm E. 2012. On the use of EDA background error variances in the ECMWF 4D-Var. Q. J. R. Meteorol. Soc. 138: 1540–1559, doi: http://dx.doi.org/10.1002/qj.1899.
- Bowler NE, Arribas A, Mylne KR, Robertson KB, Beare SE. 2008. The MOGREPS short-range ensemble prediction system. Q. J. R. Meteorol. Soc. 134: 703–722, doi:http://dx.doi.org/10.1002/qj.234.
- Buehner M. 2005. Ensemble-derived stationary and flow-dependent background-error covariances: Evaluation in a quasi-operational NWP setting. *Q. J. R. Meteorol.* Soc. 131: 1013–1043, doi:http://dx.doi.org/10.1256/qj.04.15.
- Buehner M. 2012. Evaluation of a spatial/spectral covariance localization approach for atmospheric data assimilation. *Mon. Weather Rev.* 140: 617–636, doi:http: //dx.doi.org/10.1175/MWR-D-10-05052.1.
- Buehner M, Charron M. 2007. Spectral and spatial localization of background-error correlations for data assimilation. Q. J. R. Meteorol. Soc. 133: 615–630, doi: http://dx.doi.org/10.1002/qj.50.
- Buehner M, McTaggart-Cowan R, Beaulne A, Charette C, Garand L, Heilliette S, Lapalme E, Laroche S, Macpherson SR, Morneau J, Zadra A. 2015. Implementation of deterministic weather forecasting systems based on ensemble-variational data assimilation at environment canada. part i: The global system. *Mon. Weather Rev.* submitted.

Campbell WF, Bishop CH, Hodyss D. 2010. Vertical Covariance Localization for Satellite Radiances in Ensemble Kalman Filters. Mon. Weather Rev. 138(1): 282-290.

- Caron JF, Milewski T, Buehner M, Fillion L, Reszka M, Macpherson S, St-James J. 2015. Implementation of deterministic forecasting systems based on ensemblevariational data assimilation at environment canada. part ii: The regional system. monitoring the observation impact on the short-range forecast. *Mon. Weather Rev.* **submitted**.
- Clayton AM, Lorenc AC, Barker DM. 2013. Operational implementation of a hybrid ensemble/4D-Var global data assimilation system at the Met Office. Q. J. R. Meteorol. Soc. 139: 1445–1461, doi:http://dx.doi.org/10.1002/qj.2054.



- Flowerdew J, Bowler NE. 2011. Improving the use of observations to calibrate ensemble spread. *Q. J. R. Meteorol. Soc.* **137**: 467–482, doi:http://dx.doi.org/10.1002/ qj.744.
- Harlim J, Hunt BR. 2007. Four-dimensional local ensemble transform Kalman filter: numerical experiments with a global circulation model. *Tellus A* **59**(5): 731–748, doi:http://dx.doi.org/10.1111/j.1600-0870.2007.00255.x.
- Houtekamer P, Deng X, Mitchell HL, Baek SJ, Gagnon N. 2014. Higher resolution in an operational ensemble Kalman filter. Mon. Weather Rev. 142(3): 1143–1162.
- Hunt BR, Kostelich EJ, Szunyogh I. 2007. Efficient data assimilation for spatiotemporal chaos: A local ensemble transform Kalman filter. *Physica D* 230: 112–126, doi:10.1016/j.physd.2006.11.008.
- Liu C, Xiao Q, Wang B. 2009. An ensemble-based four-dimensional variational data assimilation scheme. part ii: Observing System Simulation Experiments with advanced research WRF (ARW). *Mon. Weather Rev.* **137**: 1687–1704, doi:http://dx.doi.org/10.1175/2008MWR2699.1.
- Lorenc AC. 2003. The potential of the ensemble Kalman filter for NWP a comparison with 4D-Var. Q. J. R. Meteorol. Soc. 129: 3183–3203, doi:http://dx.doi.org/ 10.1256/qj.02.132.
- Lorenc AC, Bowler NE, Clayton AM, Fairbairn D, Pring SR. 2014. Comparison of hybrid-4DEnVar and hybrid-4DVar data assimilation methods for global NWP. *Mon. Weather Rev.* accepted: 19pp, doi:10.1175/MWR-D-14-00195.1.
- Ota Y, Derber JC, Kalnay E, Miyoshi T. 2013. Ensemble-based observation impact estimates using the NCEP GFS. *Tellus A* **65**, doi:http://dx.doi.org/10.3402/tellusa. v65i0.20038.
- Polavarapu S, Ren S, Clayton A, Sankey D, Rochon Y. 2004. On the relationship between incremental analysis updating and incremental digital filtering. *Mon. Weather Rev.* **132**: 2495–2502, doi:10.1175/1520-0493(2004)132<2495:OTRBIA>2.0.CO;2.

My related talks:

Comparing ensemble-variational assimilation methods for NWP: Can hybrid-4DEnVar match hybrid-4DVar? WWOSC, Montreal, August 2014. <u>https://www.wmo.int/pages/prog/arep/wwrp/new/wwosc/documents/Lorenc_4DEnVar_4DVar.pdf</u>

Advances in data assimilation techniques and their relevance to satellite data assimilation. ECMWF Seminar on Use of Satellite Observations in NWP,8-12 September 2014. <u>http://www.ecmwf.int/sites/default/files/Lorenc_AdvancesDAsatellites.pdf</u>