

Weakly coupled assimilation of hydrographic profiles into isopycnal ocean models with ensemble data assimilation method

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Introduction

- Hydrographic profiles are crucial to initialise the vertical structure of ocean models (Oke and Schiller, 2007; Zhang et al., 2009; Karspeck et al., 2013; Brune et al., 2015)

Objective: complement NorCPM with assimilation of T&S

1. The 'best' implementation setting
 - Analysis scheme (innovation coordinate)
 - Localisation radius
 - Observation error estimation
2. Short-term reanalysis with this 'best' setting

1.1 Analysis scheme

*Is it better to assimilate observation in z-coordinate (**EnKF-z**) or to convert them in isopycnal coordinate before assimilation (**EnKF-p**)?*

1) **EnKF-z**:

$$\bar{\mathbf{X}}_a = \bar{\mathbf{X}}_f + \mathbf{P}_f \mathbf{H}^T (\mathbf{H} \mathbf{P}_f \mathbf{H}^T + \mathbf{R}_z)^{-1} (\mathbf{y} - \mathbf{H}(\bar{\mathbf{X}}_f)),$$

where H is an operator from model space to observation space.

2) **EnKF-p** (Xie and Zhu, (2010), Thacker and Esenkov, 2002):

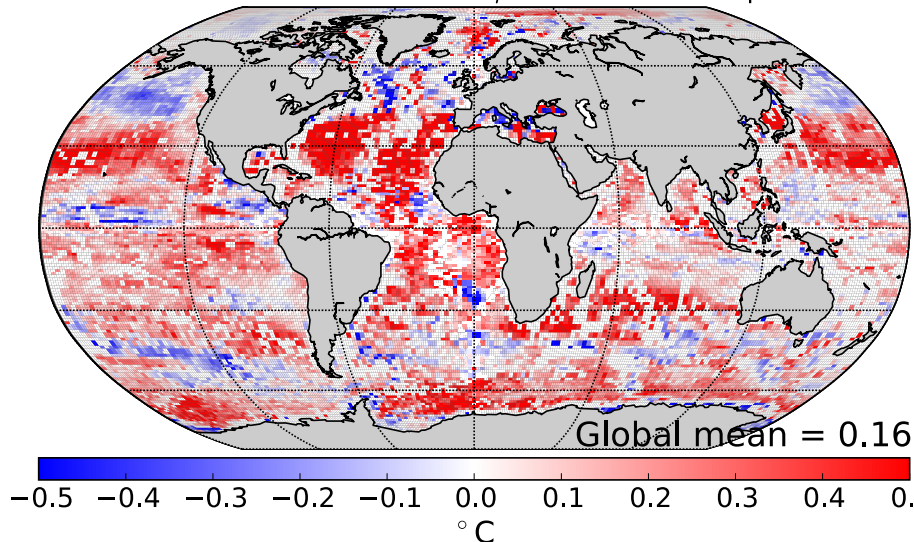
$$\bar{\mathbf{X}}_a = \bar{\mathbf{X}}_f + \mathbf{P}_f (\mathbf{P}_f + \mathbf{R}_r)^{-1} (F(\mathbf{y}) - \bar{\mathbf{X}}_f),$$

where F is a non-linear operator from observation space to model space.

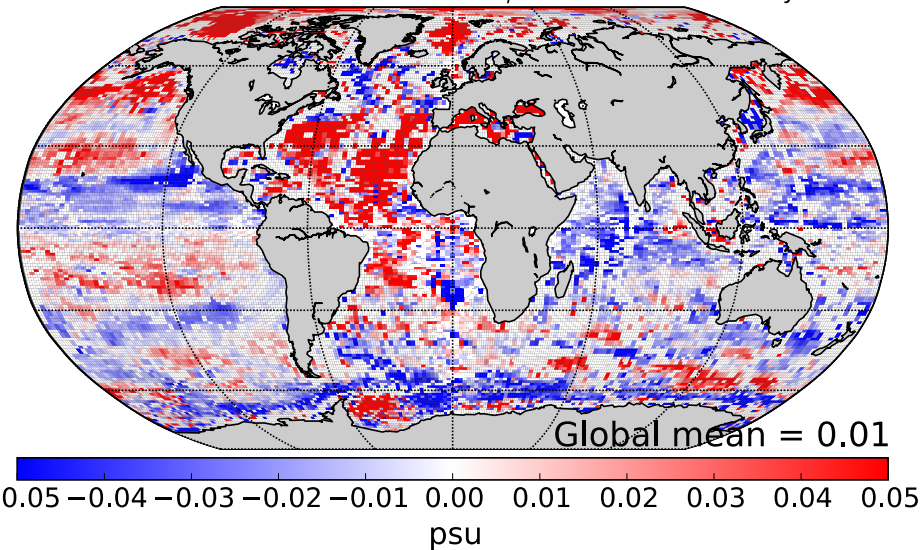
$$\mathbf{R}_r = \mathbf{F} \mathbf{R}_z \mathbf{F}^T$$

EnKF-z vs. EnKF- ρ in OSSE

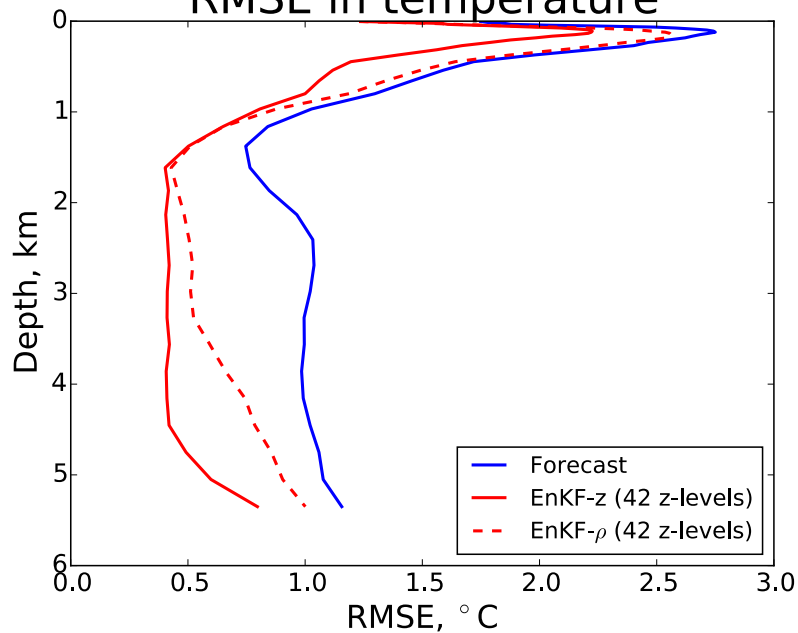
Difference of RMSEs in the EnKF- ρ and EnKF-z for temperature



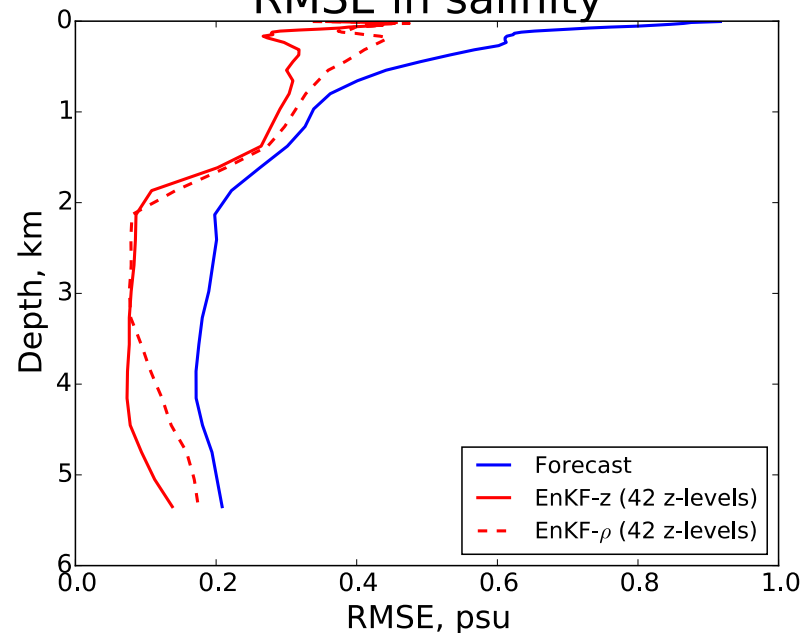
Difference of RMSEs in the EnKF- ρ and EnKF-z for salinity



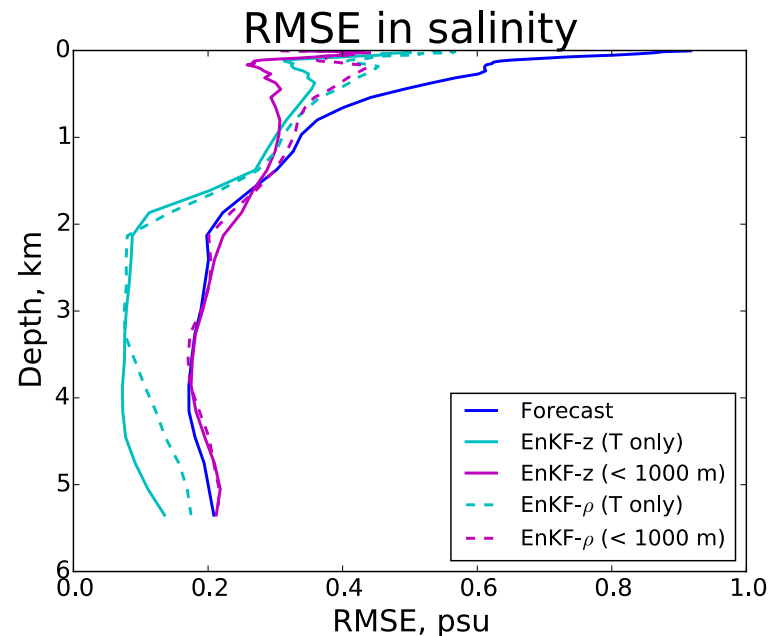
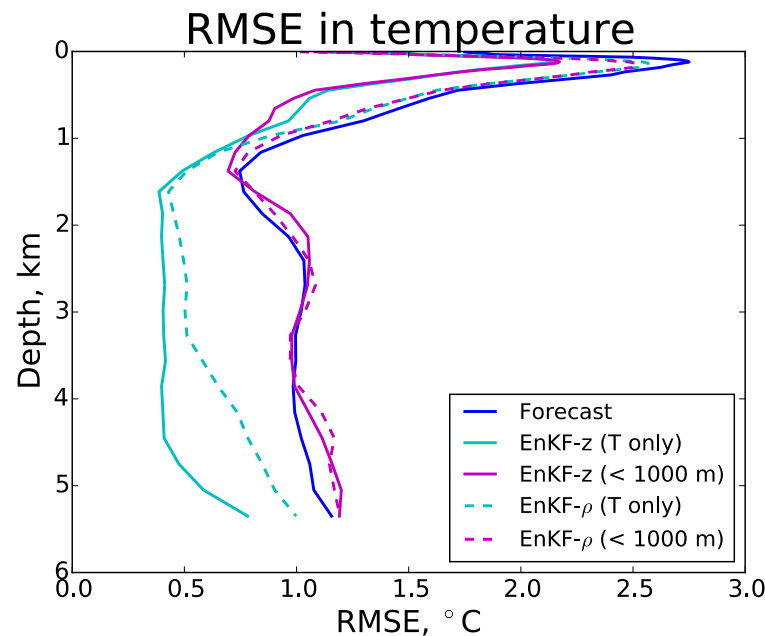
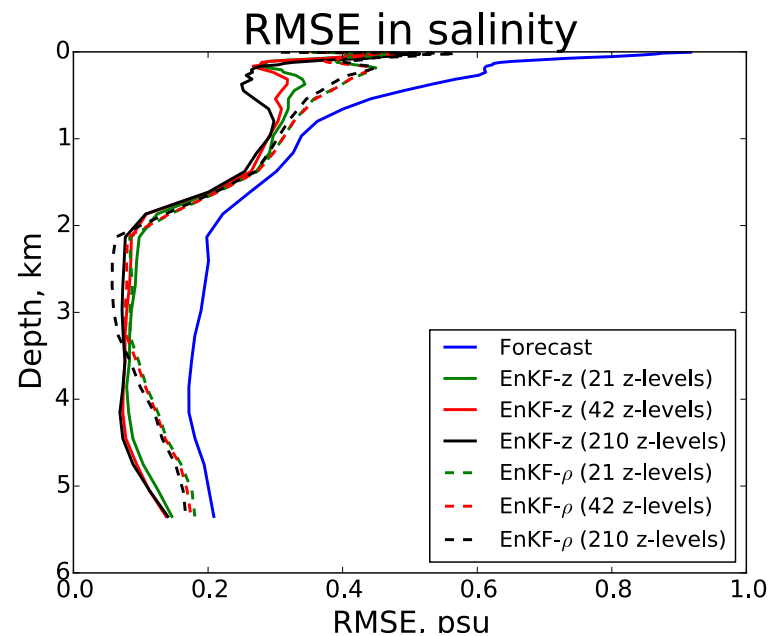
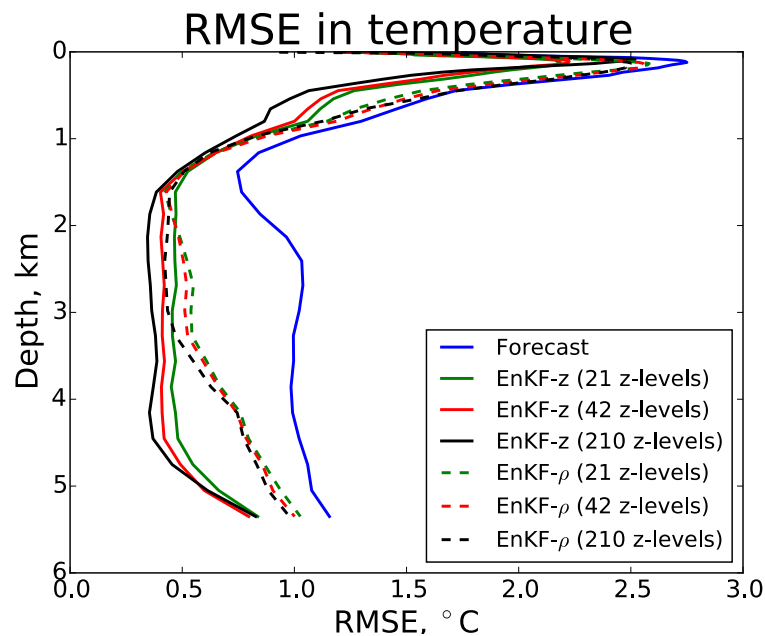
RMSE in temperature



RMSE in salinity

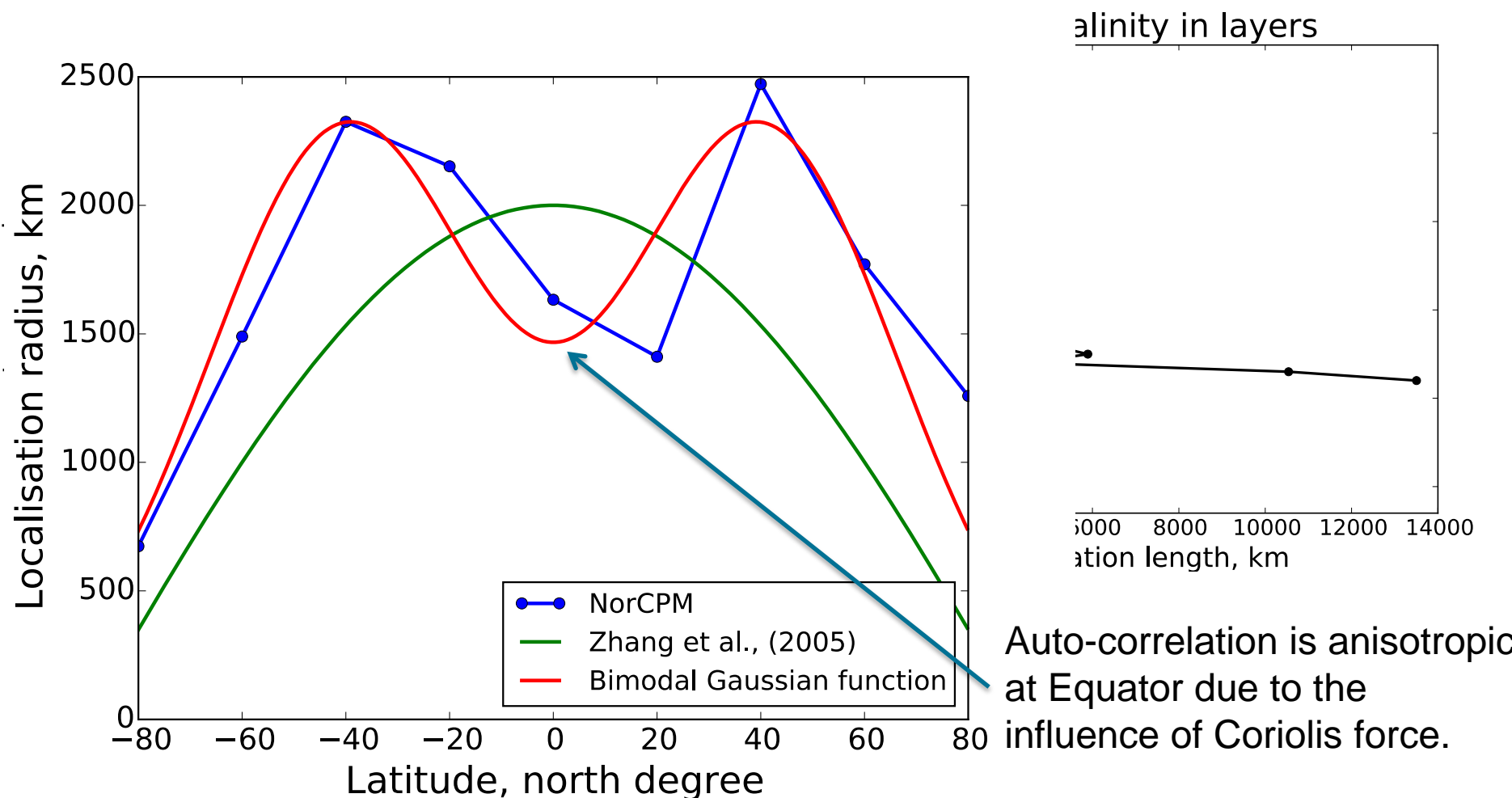


Sensitivity on other datasets

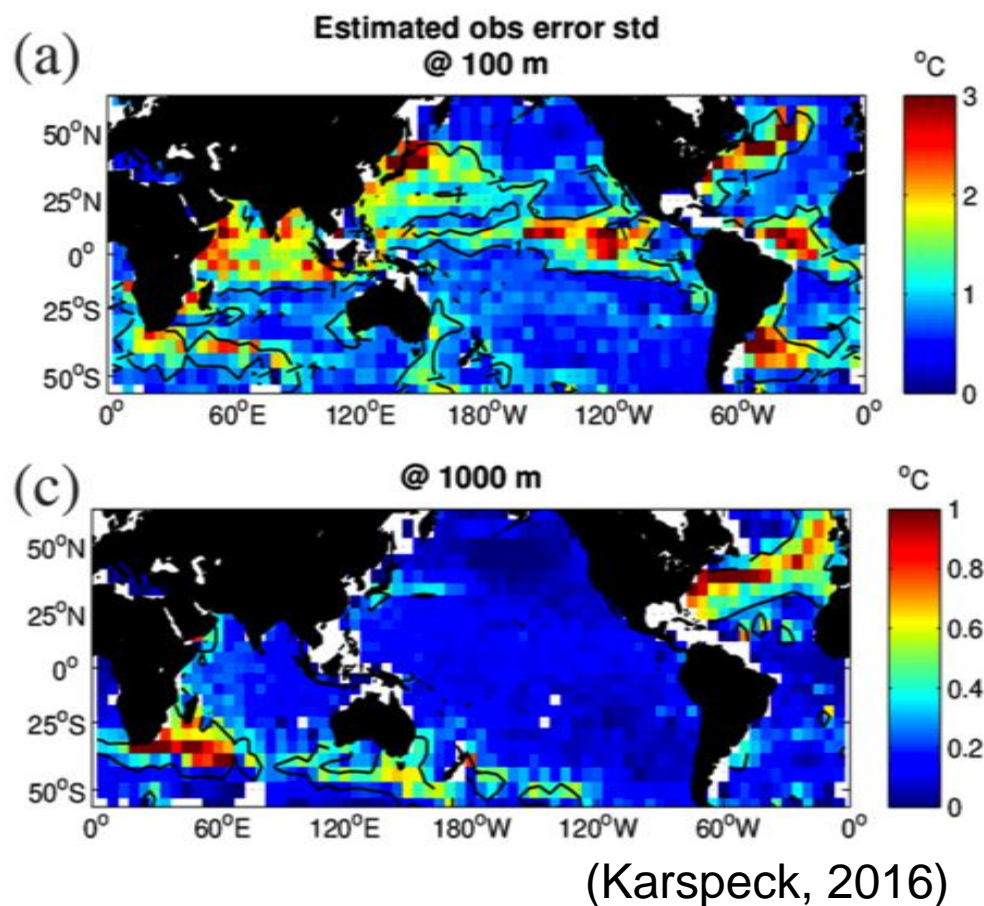
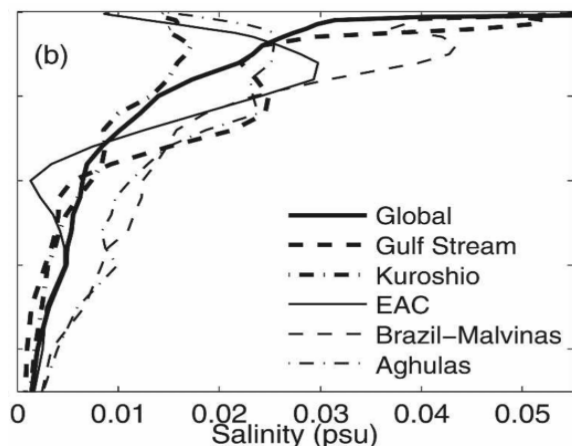
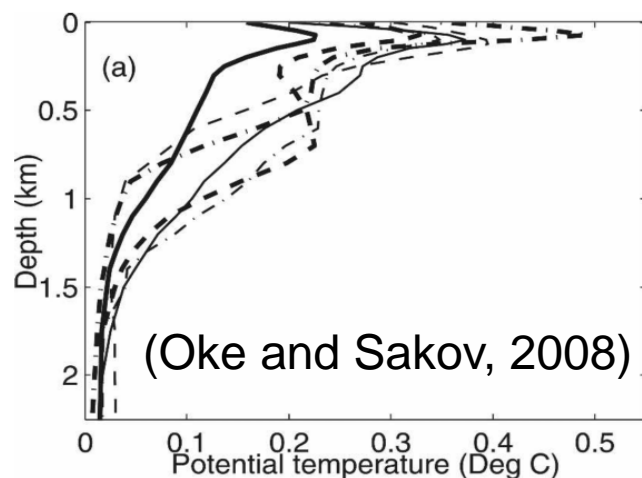


1.2 Localization in ensemble DA

- Ensemble size is too small to span the whole model space (spurious correlations).
- We approximate here the horizontal localisation radius by the auto-correlation length scale that best fits the Gaspari and Cohn function.



1.3 Observation error



The error estimation of Karspeck (2016):

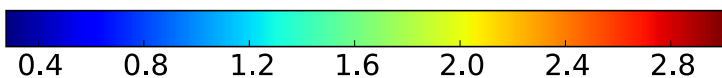
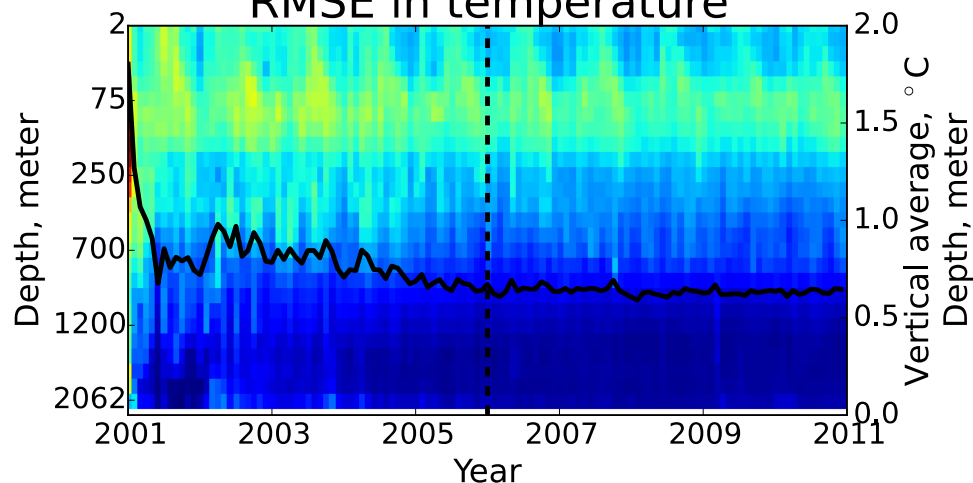
$$\text{ObsVar} = \text{RMSE}^2 - \text{BIAS}^2 - \text{EnsVar}$$

2. Verification of the stability of the system

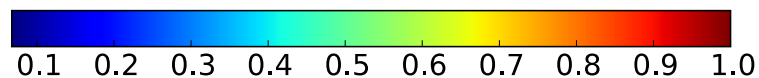
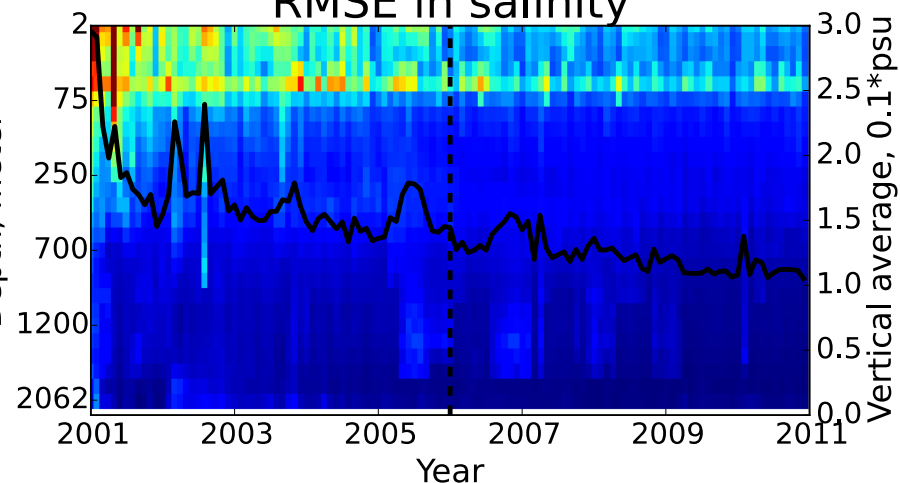
- Carry out a 10-year reanalysis (2001-2010)
 - Observations: EN4 dataset
 - Model: NorCPM-L (2° atmosphere, ocean, sea ice)
 - 30 ensemble members
 - Monthly assimilation

Accuracy: RMSE and bias

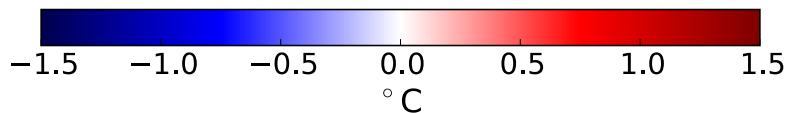
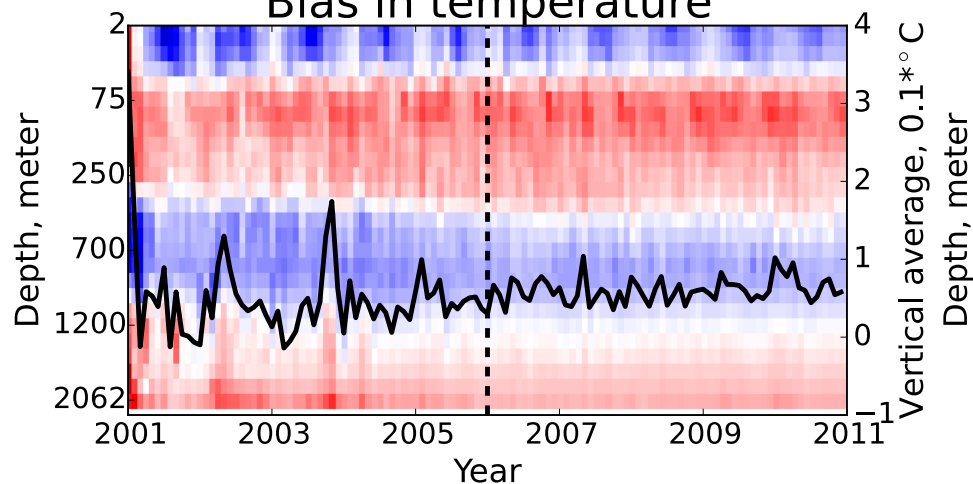
RMSE in temperature



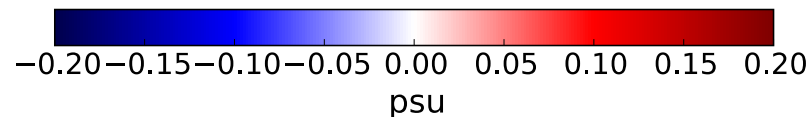
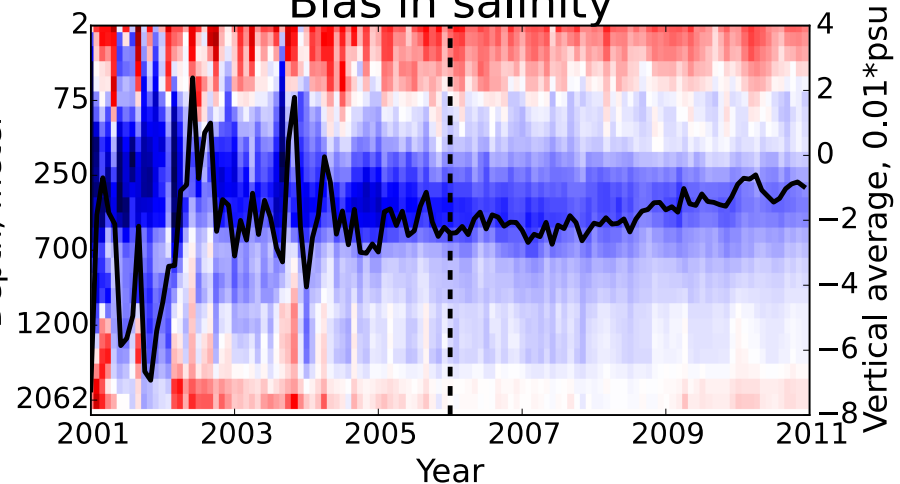
RMSE in salinity



Bias in temperature

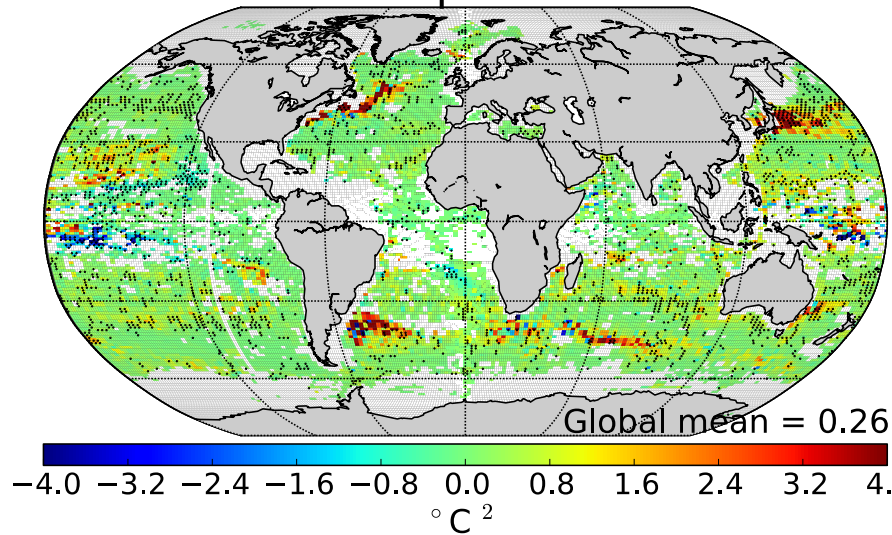


Bias in salinity

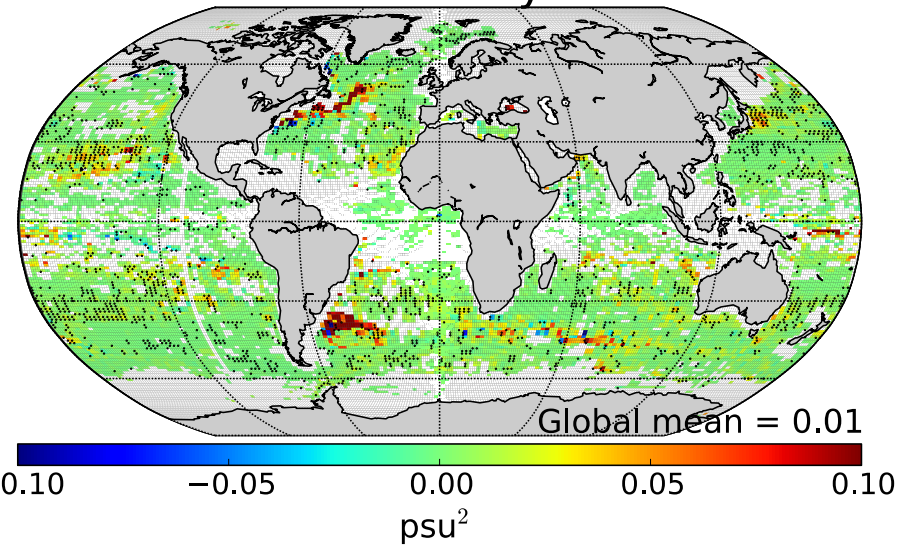


Reliability budget (Rodwell et al. 2016)

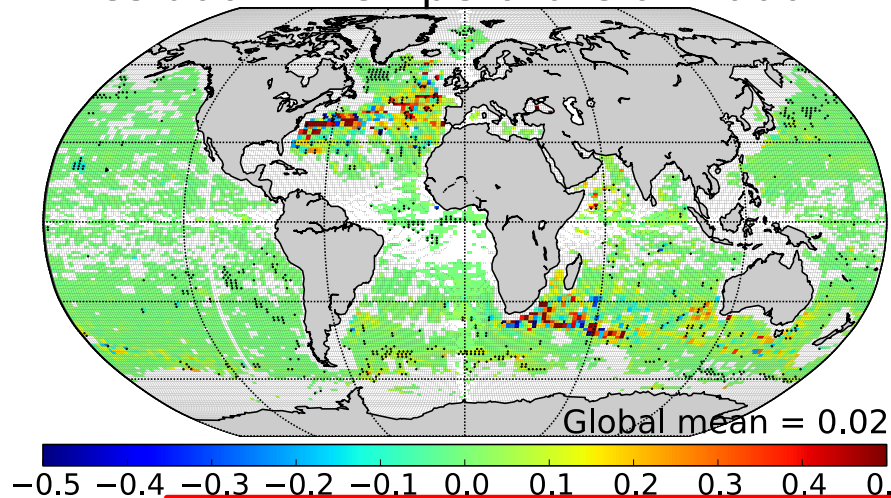
Residual in temperature at 200 m



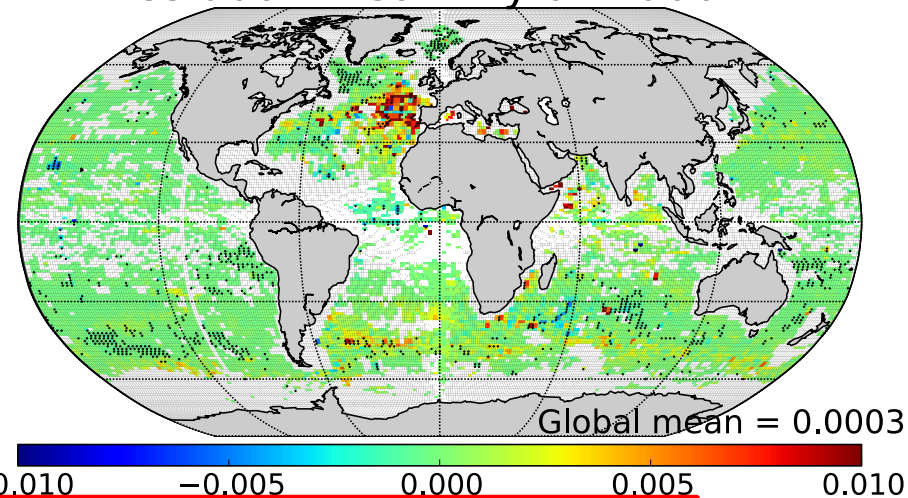
Residual in salinity at 200 m



Residual in temperature at 1000 m



Residual in salinity at 1000 m



$$r = \text{RMSE}^2 - \text{BIAS}^2 - \text{Ens Var} - \text{Obs Var}$$

Conclusions

- › **Implementation setting of assimilation of T-S profiles:**
 - It is preferable to assimilate observations in their original coordinate (z-coordinate).
 - The localization radius varies with latitude (a bimodal Gaussian).
 - Observation errors are estimated in regions.
- › **Short-term reanalysis**
 - DA reduces efficiently RMSE and bias.
 - The system requires 5-year spin-up assim to converge to a stable performance.
 - The system is statistically reliable in most regions.

Wang et al.: optimising assimilation of hydrographic profiles into isopycnal ocean models with ensemble data assimilation, in preparation.

Thanks

Questions?