



Multifrequency radar retrievals of rain microphysics

Evaluation of the rain representation in the WRF Model

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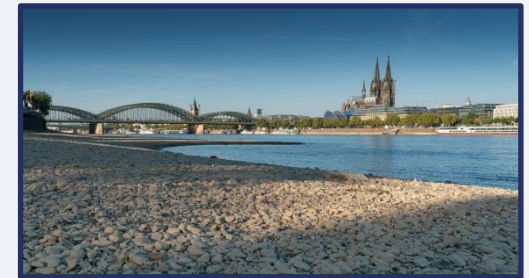
Introduction

► Why study precipitation?

- forecast at regional scale (flood, drought,...)
- quantify its impact at the climate scale via the latent heat associated with condensation and evaporation



Paris - January 2018



Cologne - November 2018

► Method?

- better represent the microphysics processes of precipitation: evaporation, breakup, self-collection...
- investigate the vertical profile of the Drop Size Distribution (DSD)

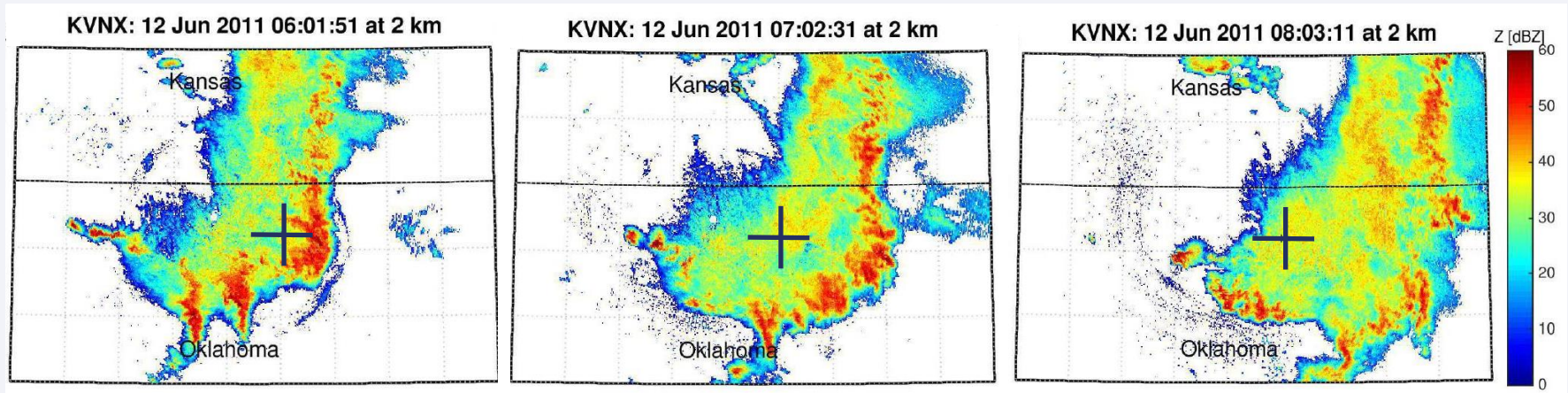
Artist's view of a Drop Size Distribution



Courtesy of NASA's Goddard Space Flight Center

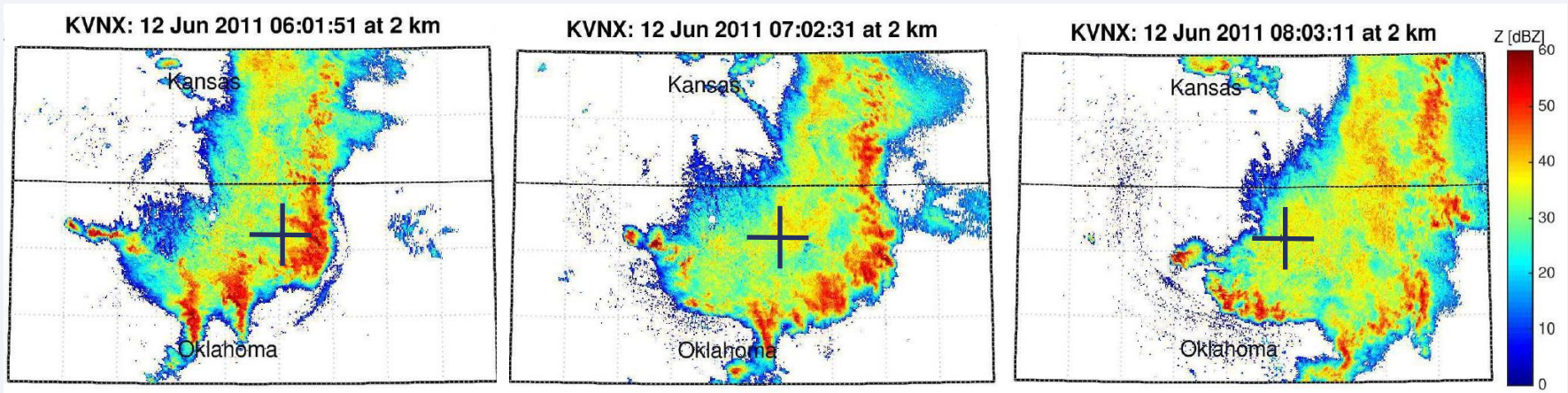
► **Objective: Evaluate the DSD representation in mesoscale models thanks to heavily instrumented observation site**

Squall line case – 12 June 2011 (Oklahoma)

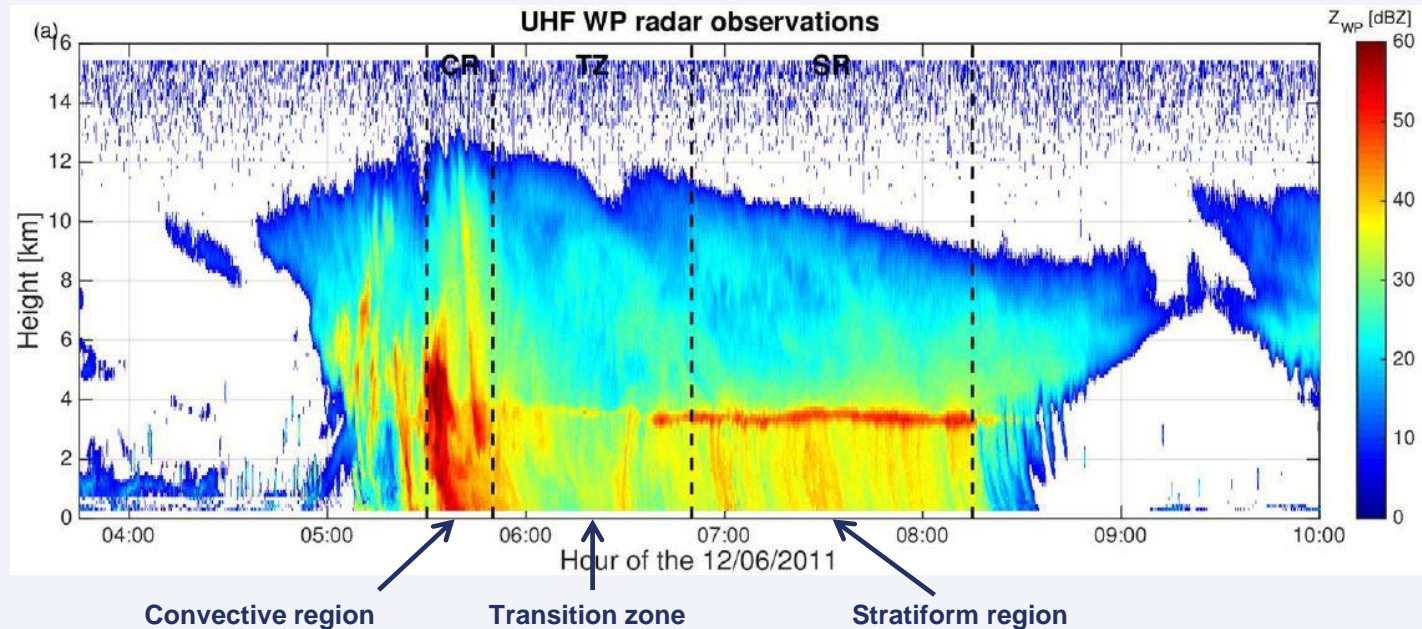


S-band
radar reflectivity
at the ARM SGP

Squall line case – 12 June 2011 (Oklahoma)



S-band and UHF
radar reflectivity
at the ARM SGP



K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site

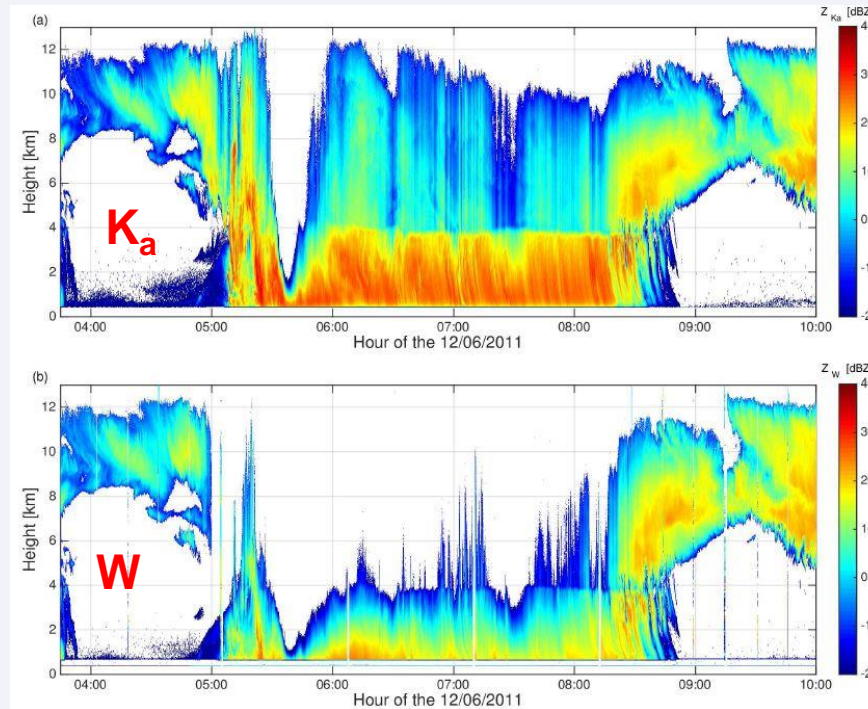


Photo courtesy of S. Collis

K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site



Photo courtesy of S. Collis



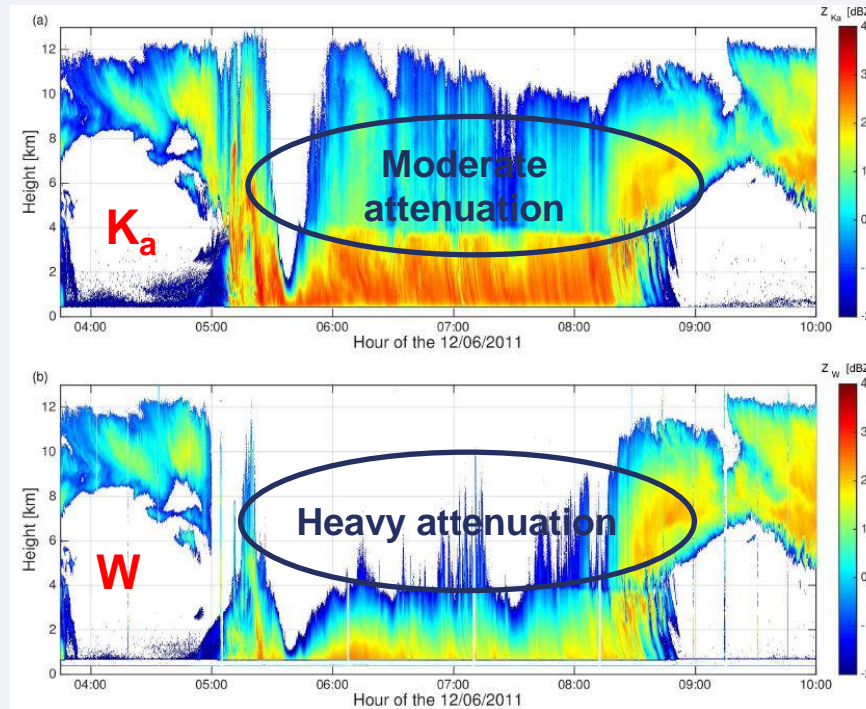
Radar reflectivity:

Z_{K_a} & Z_W

K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site



Photo courtesy of S. Collis



Radar reflectivity:

$$Z_{K_a} \text{ \& } Z_W$$

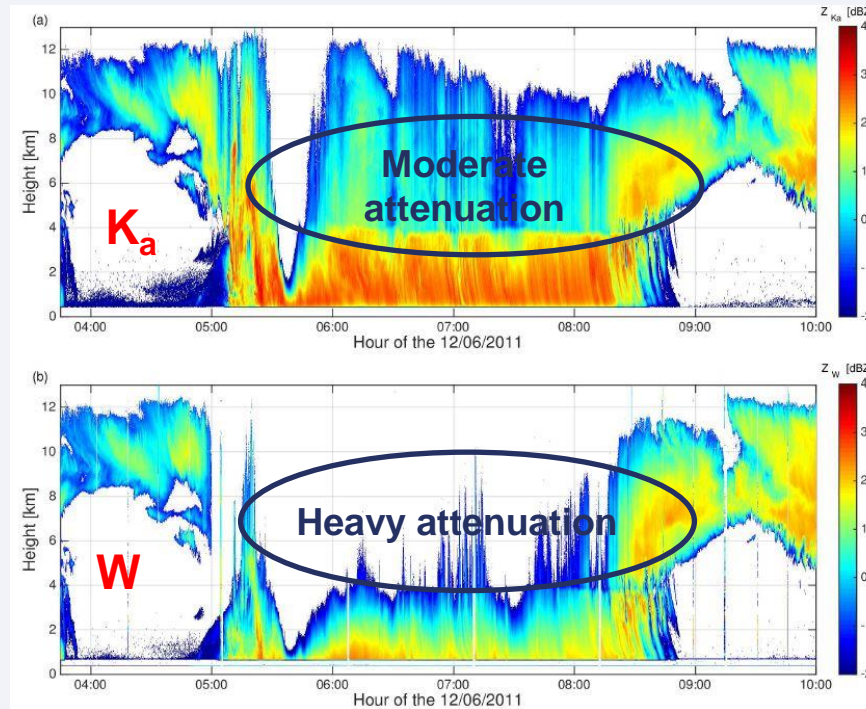
Differential atten.:

$$\Delta A$$

Diff. scattering:

$$\Delta \text{ non-Rayleigh}$$

K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site



Radar reflectivity:

$$Z_{K_a} \text{ \& \ } Z_W$$

Differential atten.: ΔA

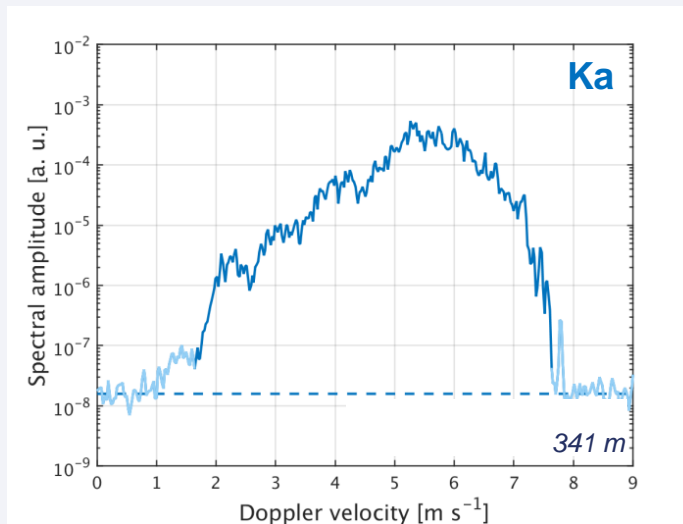
$$\Delta A$$

Diff. scattering:

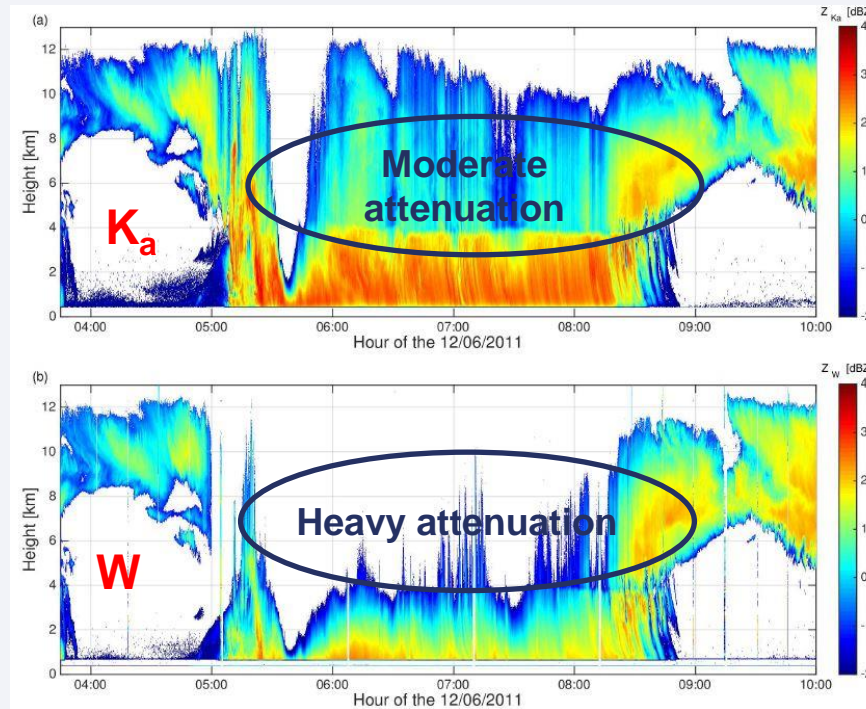
$$\Delta \text{ non-Rayleigh}$$

→ Doppler spectra:

$$S_{K_a} \text{ \& \ } S_W$$



K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site



Radar reflectivity:

$$Z_{K_a} \text{ \& \ } Z_W$$

Differential atten.: ΔA

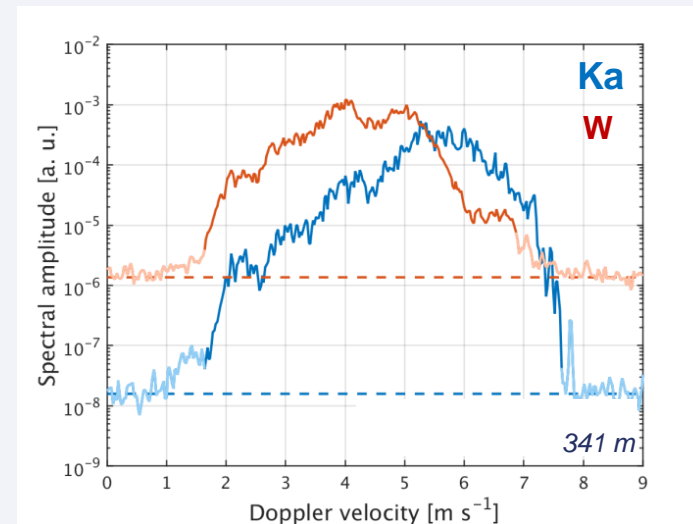
$$\Delta A$$

Diff. scattering:

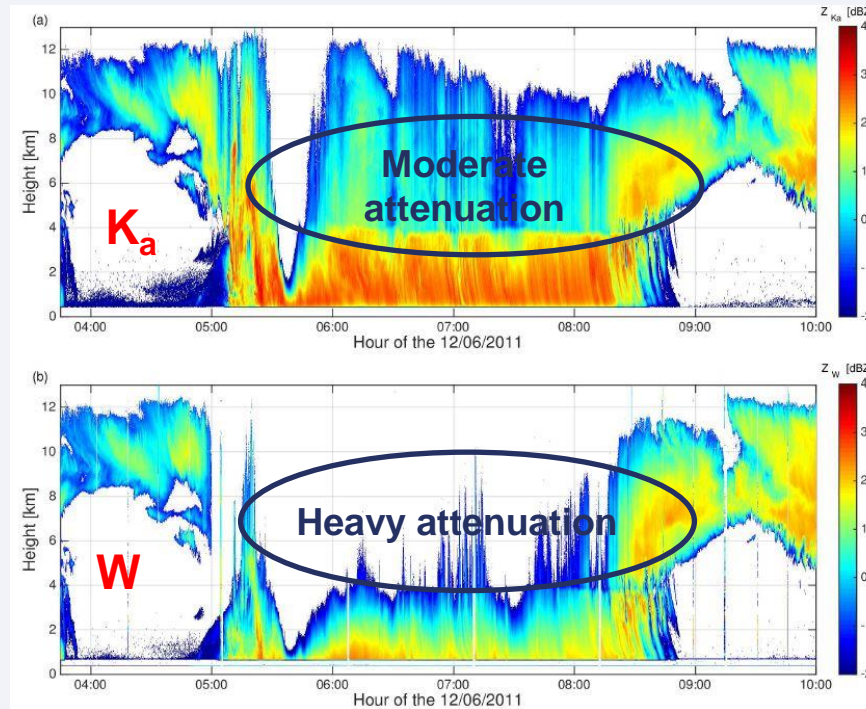
$$\Delta \text{ non-Rayleigh}$$

→ Doppler spectra:

$$S_{K_a} \text{ \& \ } S_W$$



K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site



Radar reflectivity:

$$Z_{Ka} \text{ \& \ } Z_W$$

Differential atten.: ΔA

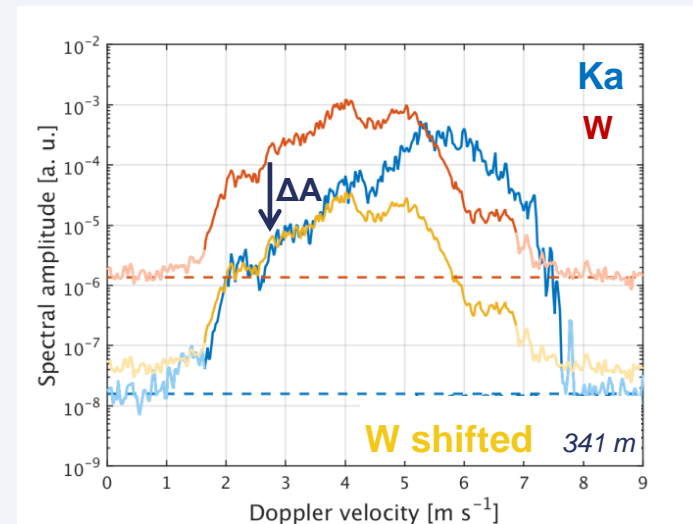
$$\Delta A$$

Diff. scattering:

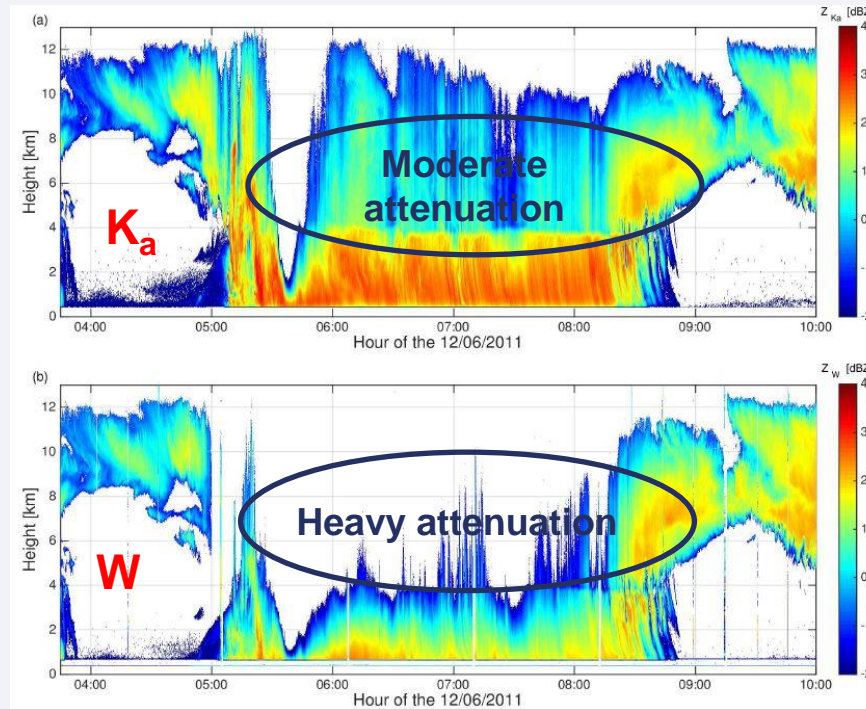
$$\Delta \text{ non-Rayleigh}$$

→ Doppler spectra:

$$S_{Ka} \text{ \& \ } S_W$$



K_a (35 GHz) and W-band (94 GHz) radar observations at ARM SGP site



Radar reflectivity:

$$Z_{K_a} \text{ \& \ } Z_W$$

Differential atten.: ΔA

$$\Delta A$$

Diff. scattering:

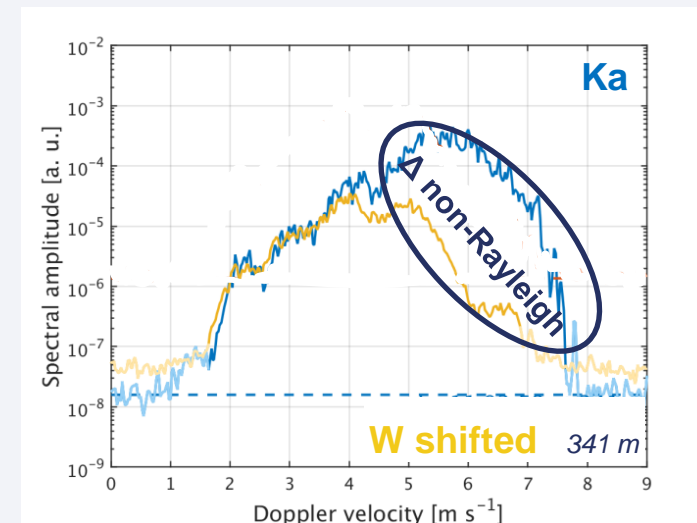
$$\Delta \text{ non-Rayleigh}$$

→ Doppler spectra:

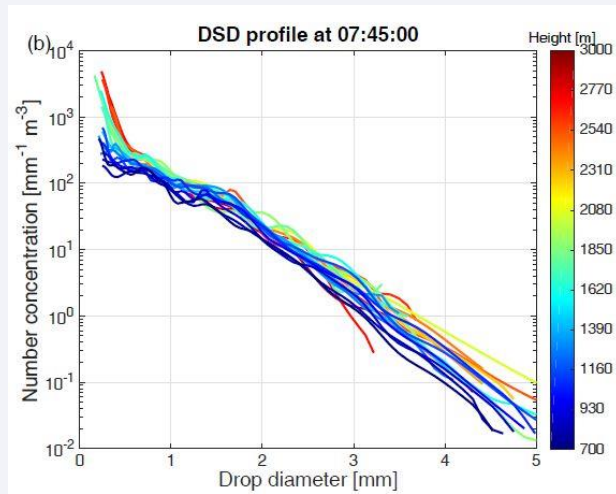
$$S_{K_a} \text{ \& \ } S_W$$

variational approach

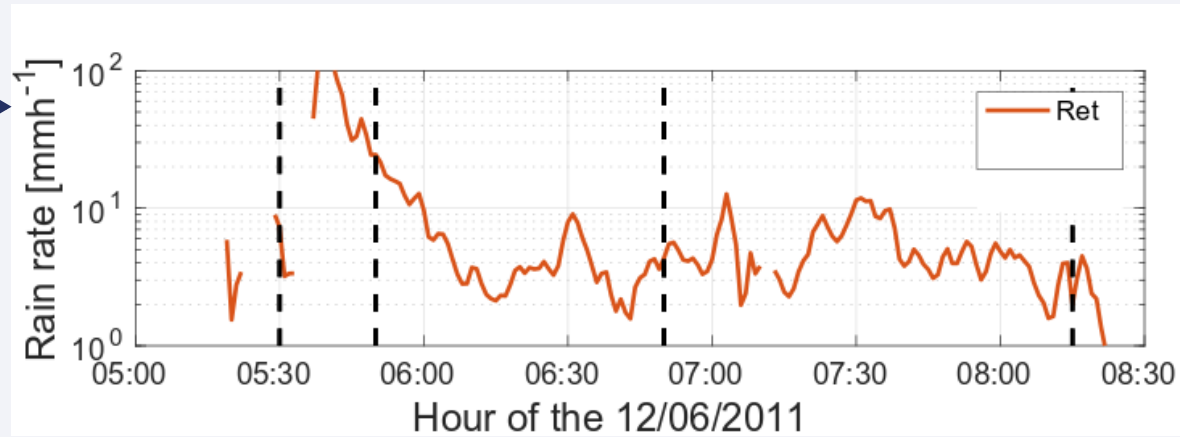
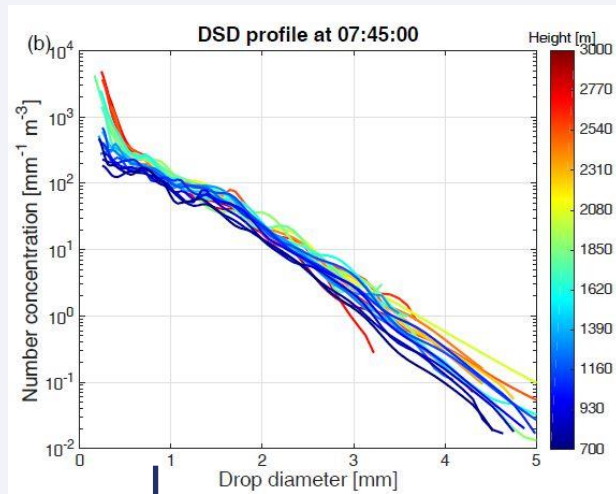
→ Retrieval of the vertical wind w and drop concentration per diameter bin (DSD) at high resolution (2 sec., 50 m)



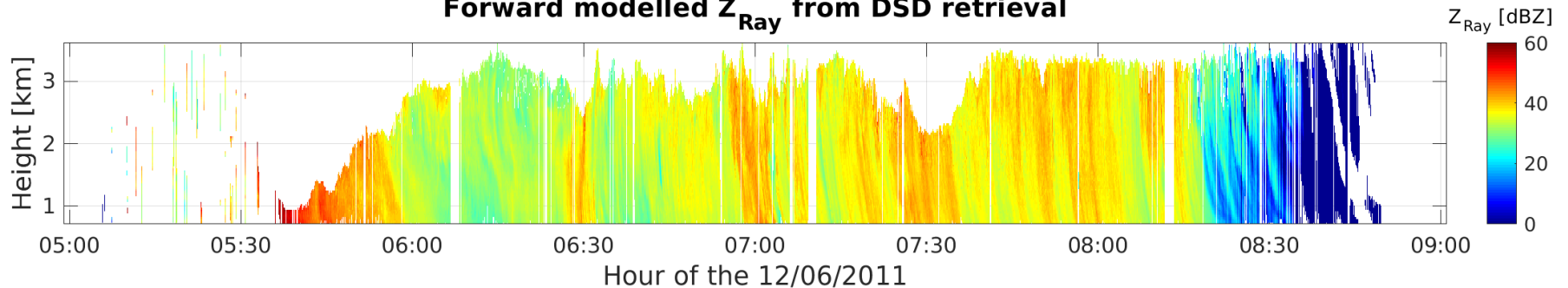
Ka and W-band radar retrieval of DSD profiles: Validation



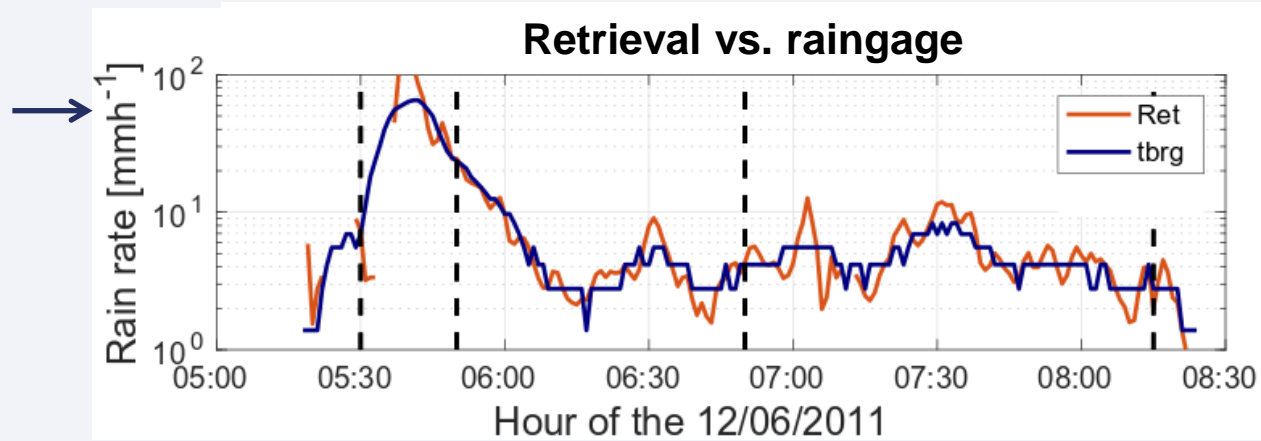
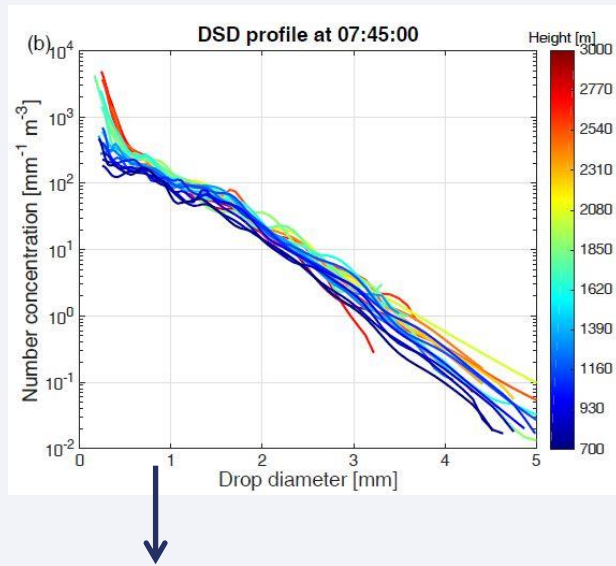
Ka and W-band radar retrieval of DSD profiles: Validation



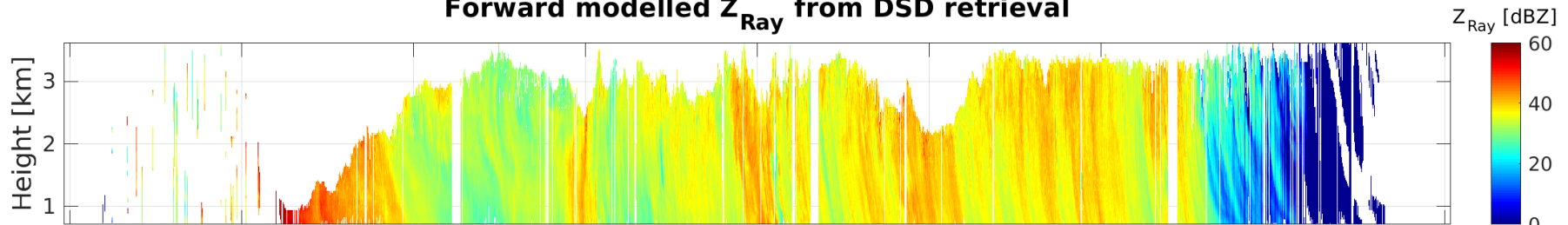
Forward modelled Z_{Ray} from DSD retrieval



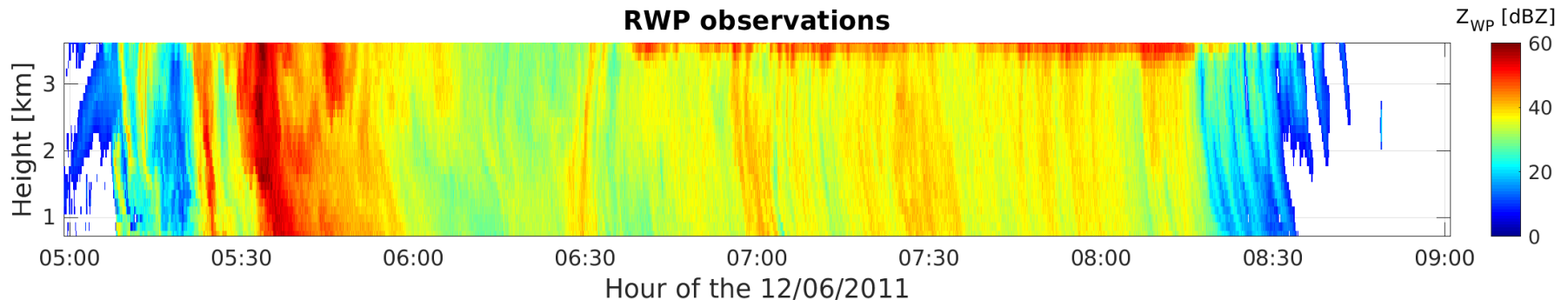
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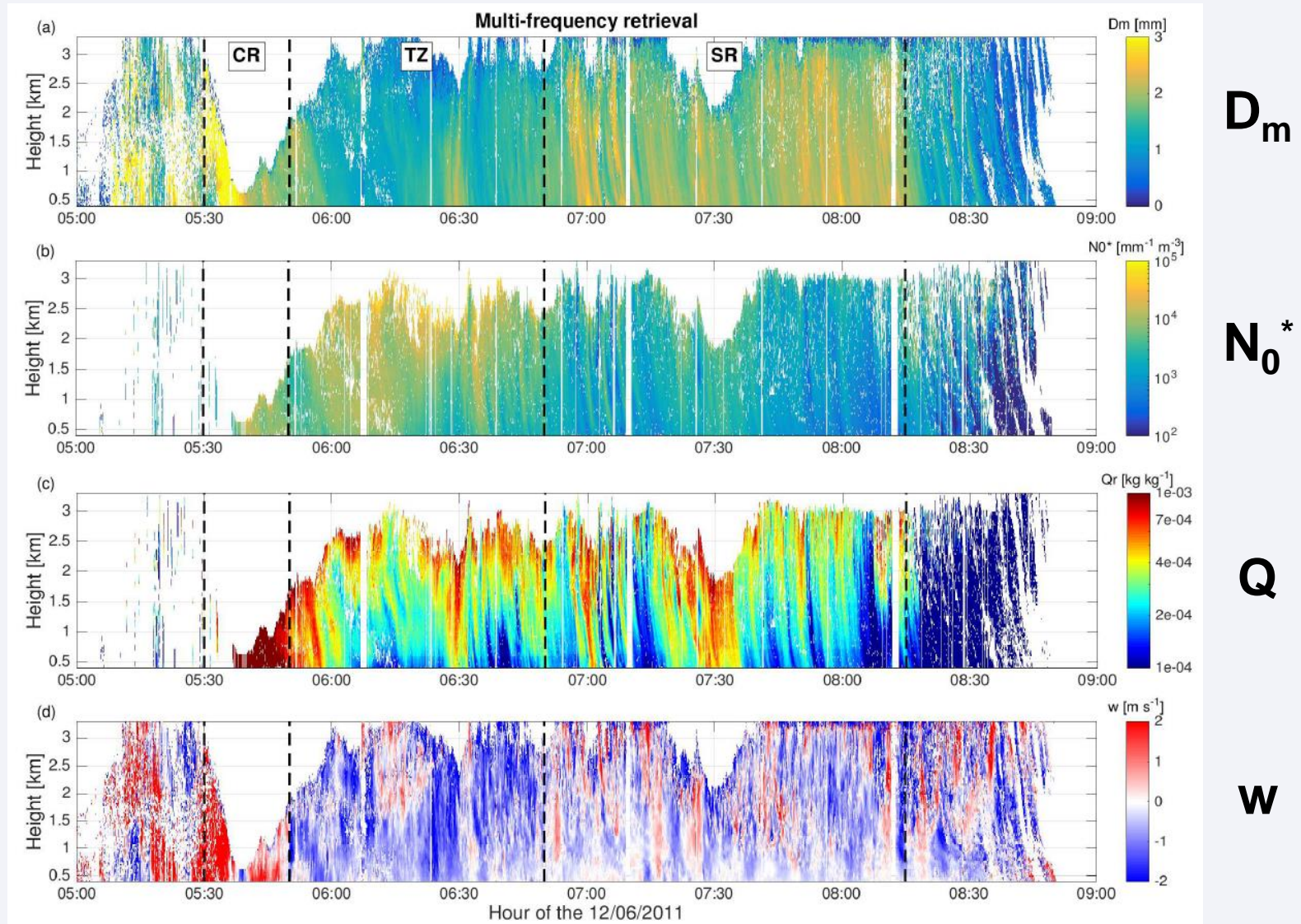
Forward modelled Z_{Ray} from DSD retrieval



RWP observations



K_a and W-band radar retrieval of DSD profiles: DSD moments



Homogeneous properties within the SR

→ Investigate the vertical variability and comparison with models possible

Model setup and initiation



THE WEATHER RESEARCH & FORECASTING MODEL

► WRF simulation

- 3 nested domains:

d01: 136 x 50 pts & $\Delta x = \Delta y = 12$ km

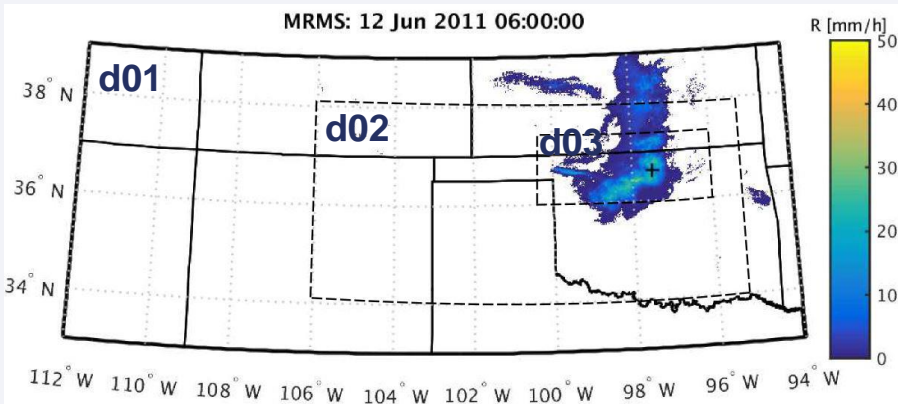
d02: 252 x 114 pts & $\Delta x = \Delta y = 4$ km

d03: 384 x 152 pts & $\Delta x = \Delta y = 1$ km

- 72 levels in the vertical coordinate:

$\Delta z \approx 250$ m

- Initiation: ERA-Interim reanalyzes data (ECMWF) at 00:00 UTC on 11 Jun. 2011



► Simulations performed using two different microphysics schemes:

Morrison et al. (2009)

Thompson et al. (2008)

Thompson scheme	Morrison scheme
Q_c, Q_r, Q_i, Q_s, Q_g	Q_c, Q_r, Q_i, Q_s, Q_g
N_r, N_i	N_r, N_i, N_s, N_g

THOM-CTL

MORR-CTL

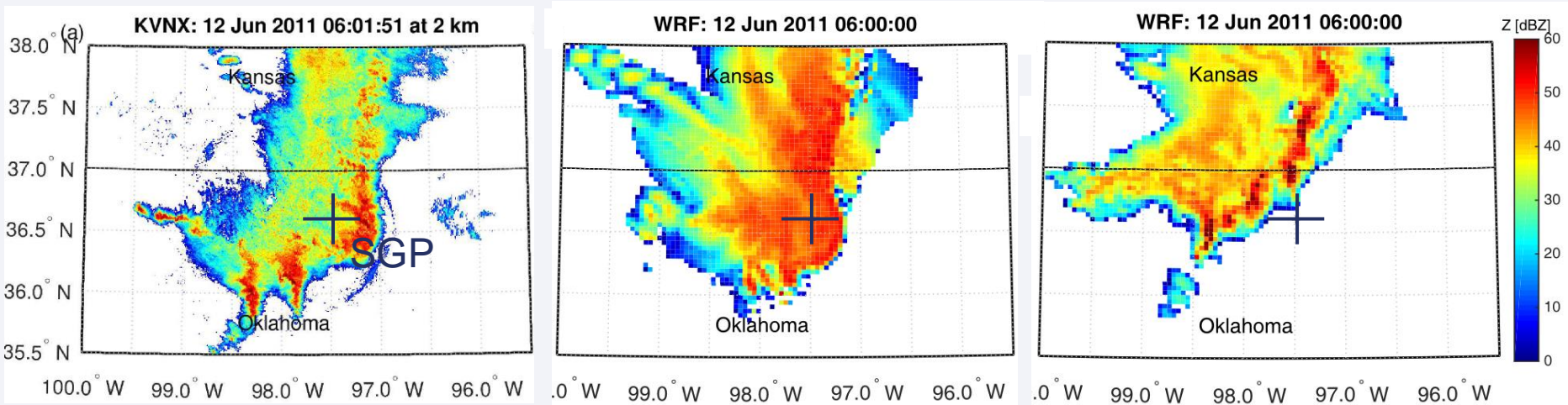
Simulations vs. observations

Radar reflectivity at 06:00 UTC

Observations

WRF-MORR

WRF-THOM

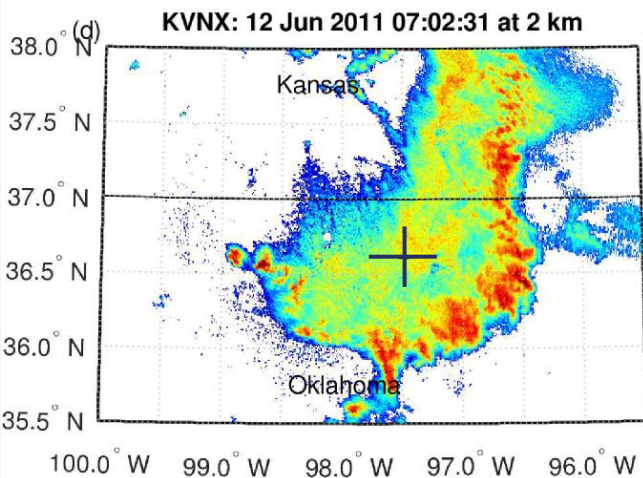


- ▶ Reasonable agreement for the location between the measured and simulated radar reflectivity.
- ▶ Simulated radar reflectivity is over-estimated using both schemes.

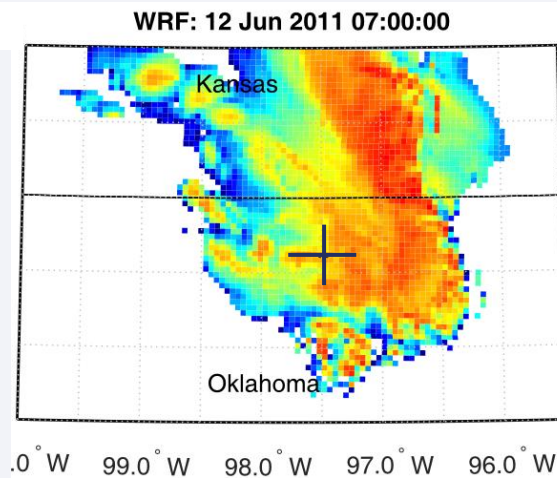
Simulations vs. observations

Radar reflectivity at 07:00 UTC

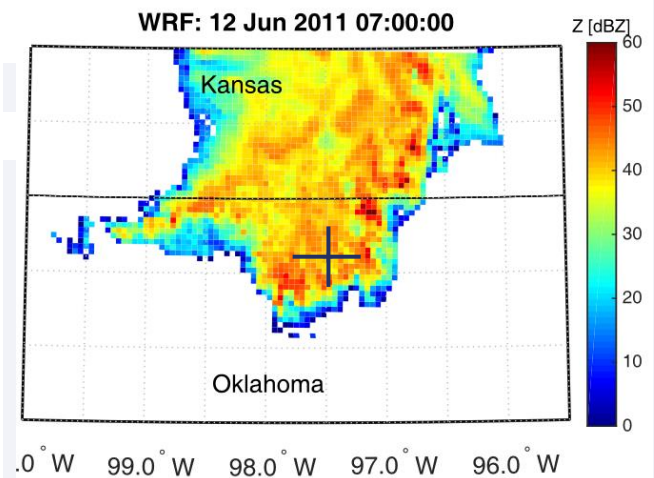
Observations



WRF-MORR



WRF-THOM



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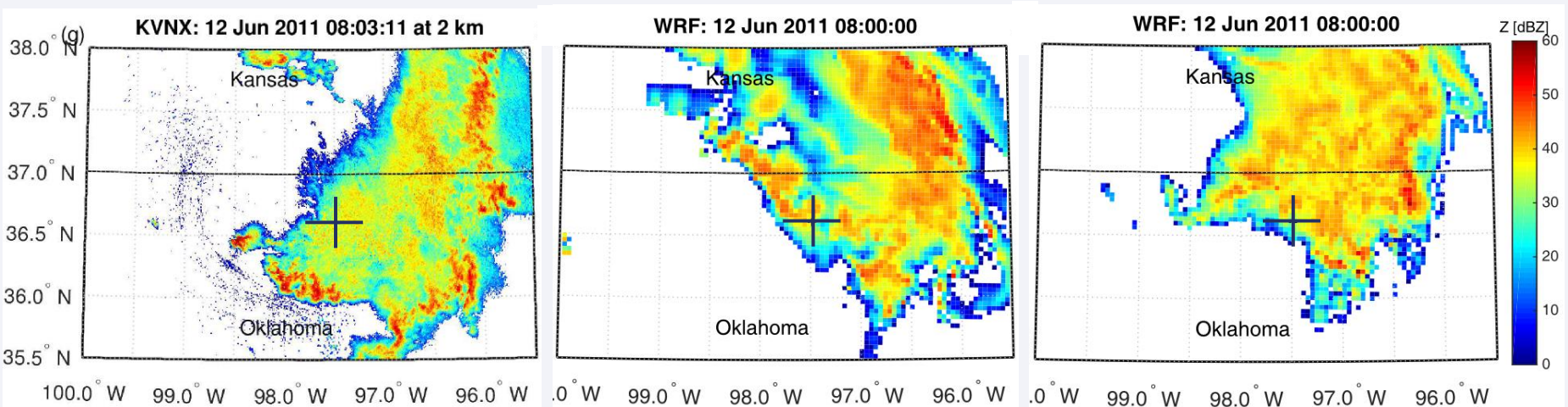
Simulations vs. observations

Radar reflectivity at 08:00 UTC

Observations

WRF-MORR

WRF-THOM



- ▶ Reasonable agreement for the location between the measured and simulated radar reflectivity.
- ▶ Simulated radar reflectivity is over-estimated using both schemes.
- ▶ Model reproduces quite well the evolution of the squall line system.
- ▶ But, both schemes fail to reproduce the Transition Zone.

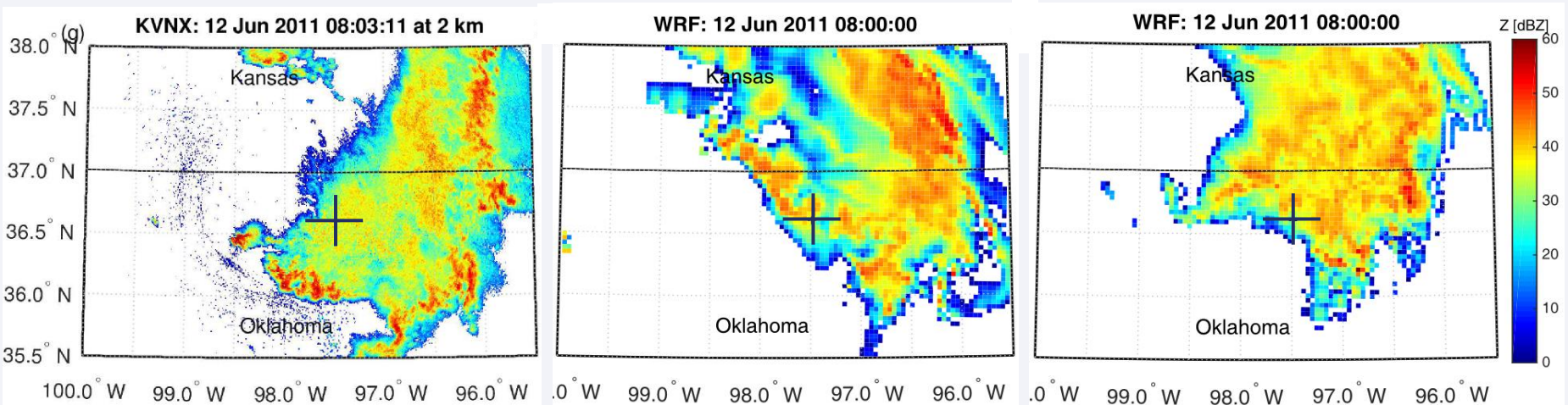
Simulations vs. observations

Radar reflectivity at 08:00 UTC

Observations

WRF-MORR

WRF-THOM



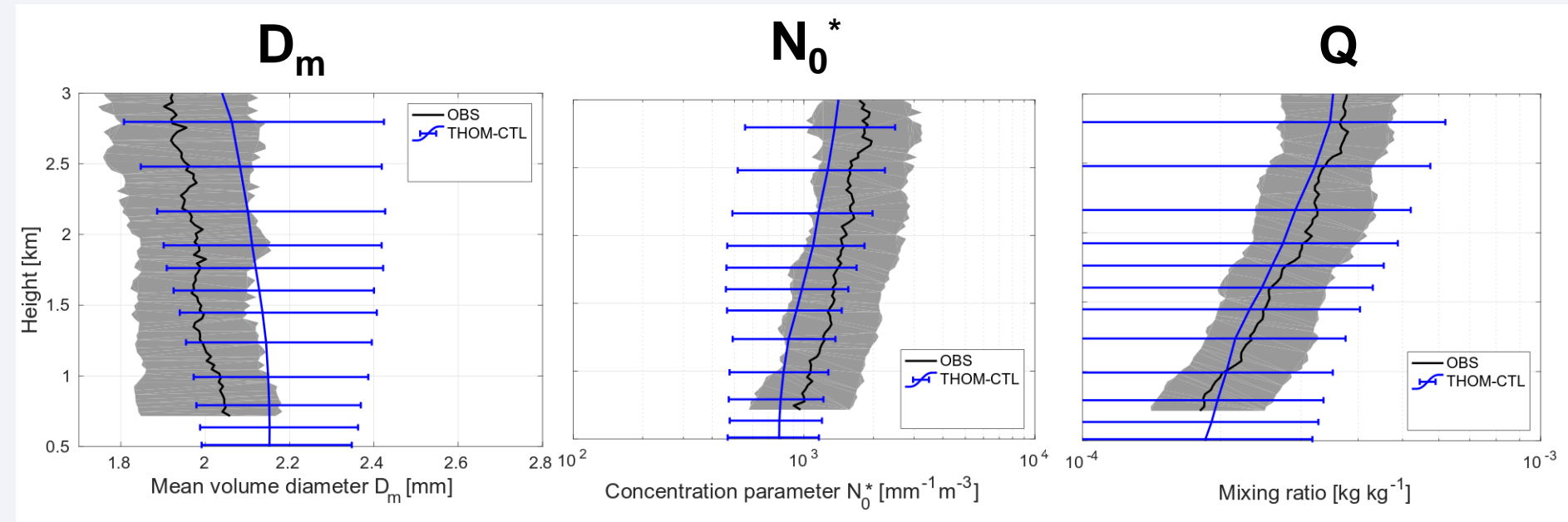
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▶ Statistical approach for profiles comparison

DSD properties

THOM-CTL

within the Stratiform Region

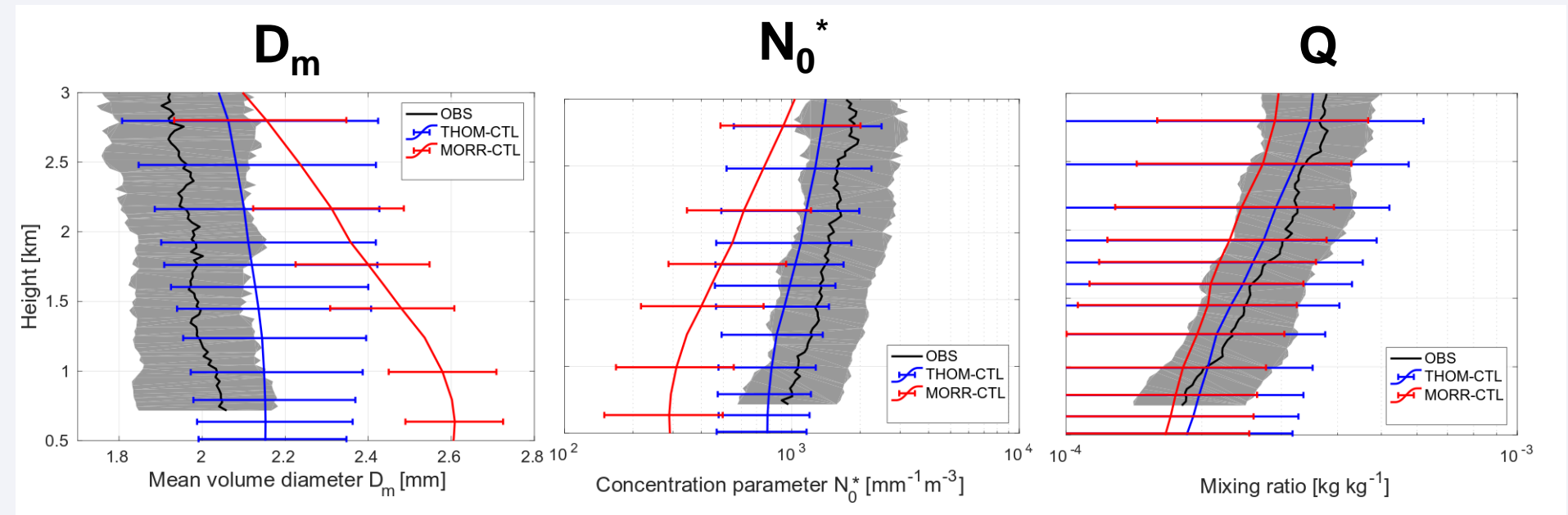


► Reasonable agreement between DSD profiles and THOM-CTL simulation

DSD properties

within the Stratiform Region

THOM-CTL
MORR-CTL



- ▶ Reasonable agreement between DSD profiles and THOM-CTL simulation
- ▶ Excessive size-sorting in MORR-CTL simulation

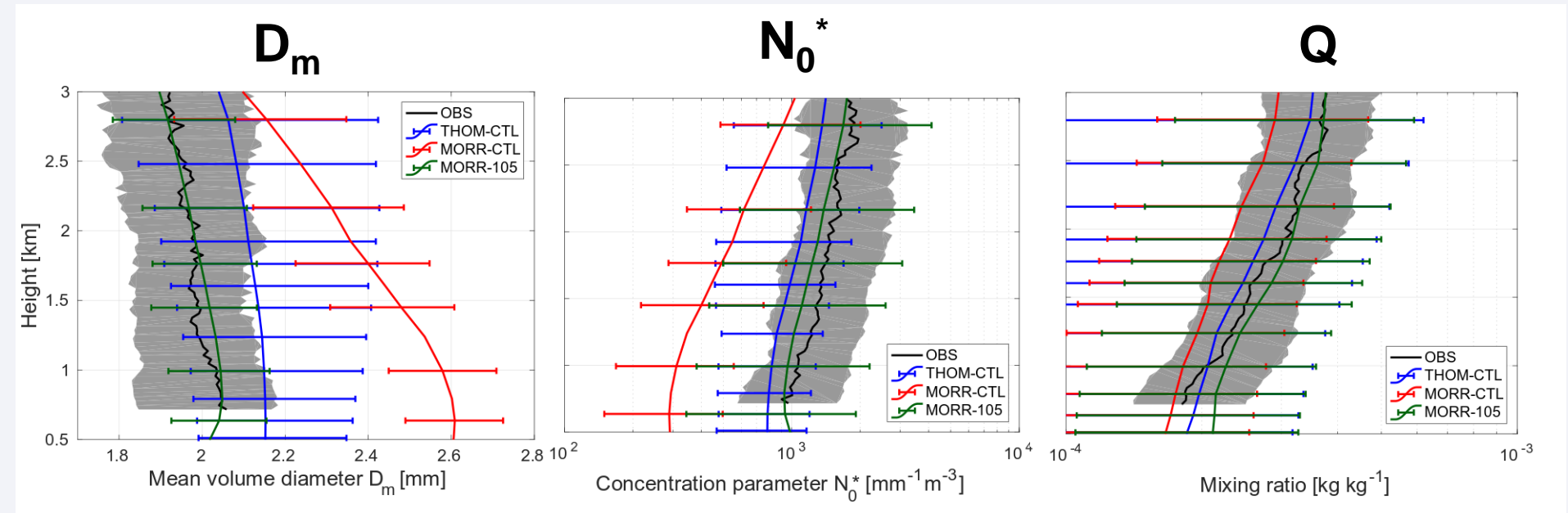
DSD properties

within the Stratiform Region

THOM-CTL

MORR-CTL

MORR-105



- ▶ Reasonable agreement between DSD profiles and THOM-CTL simulation
- ▶ Excessive size-sorting in MORR-CTL simulation
- ▶ Better agreement when the breakup efficiency is increased (as in Morrison et al., 2012)

Evaporation rate

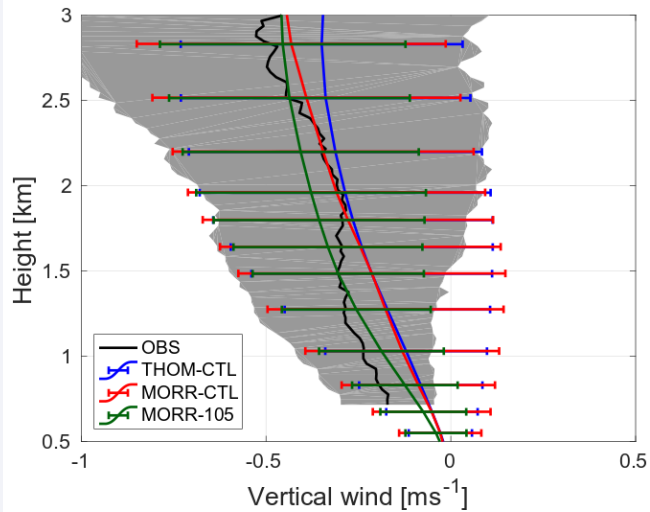
within the Stratiform Region

THOM-CTL

MORR-CTL

MORR-105

W



OBS from radar retrieval

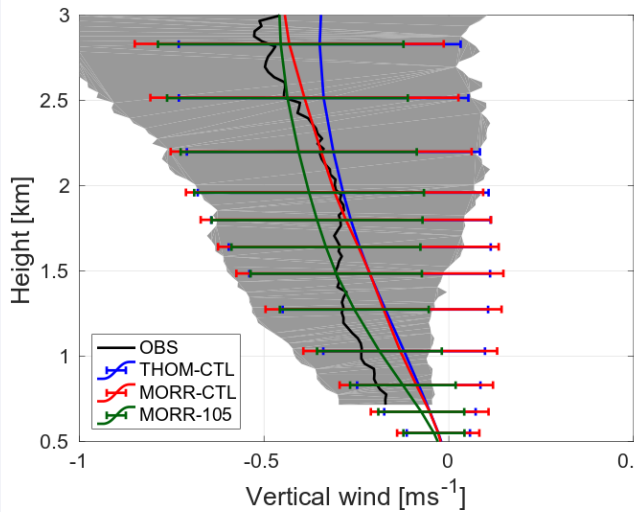
- Simulations reasonably reproduce the meso-scale downdraft associated with evaporation

Evaporation rate

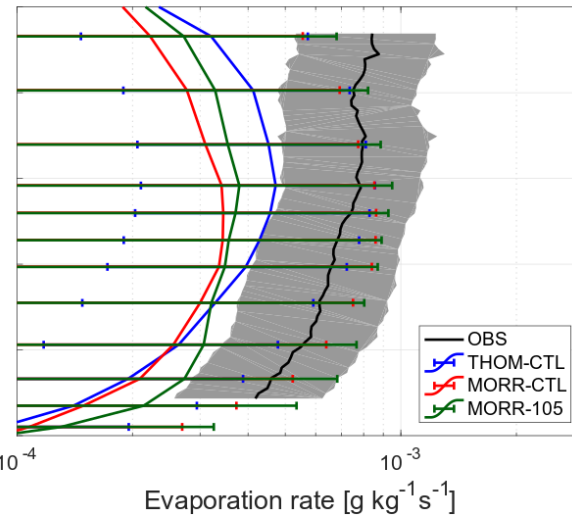
within the Stratiform Region

THOM-CTL
MORR-CTL
MORR-105

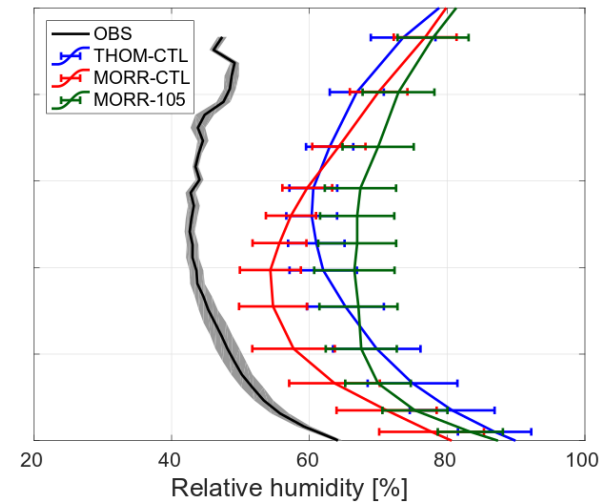
W



Evap. rate



RH



OBS from radar retrieval

OBS from Raman lidar
(Turner et al., 2002)

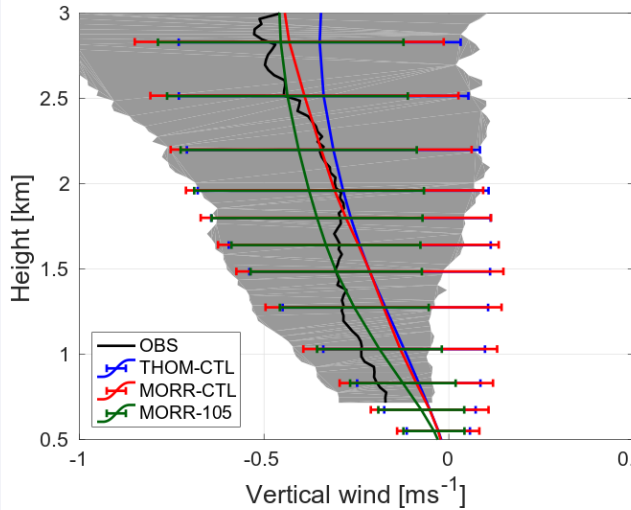
- ▶ Simulations reasonably reproduce the meso-scale downdraft associated with evaporation
- ▶ However, even if DSD profiles are well reproduced, evaporation is underestimated because RH is overestimated in the model

Evaporation rate

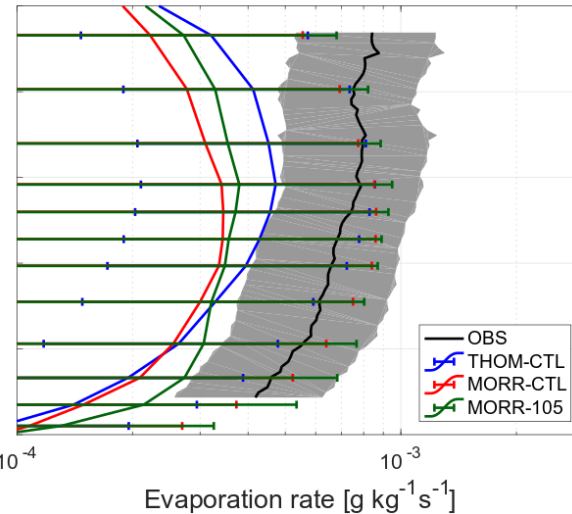
within the Stratiform Region

THOM-CTL
MORR-CTL
MORR-105

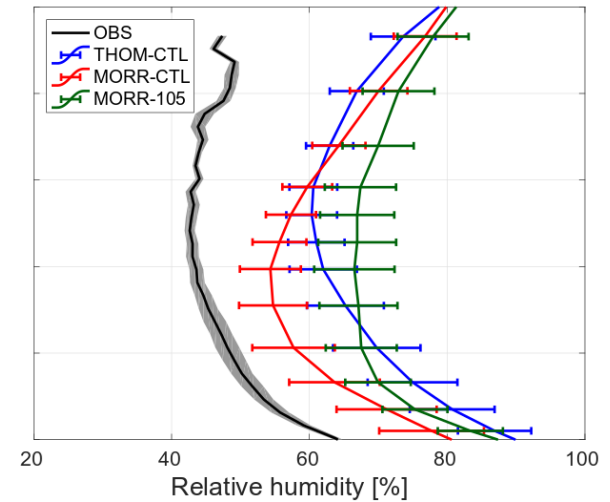
W



Evap. rate



RH



OBS from radar retrieval

OBS from Raman lidar
(Turner et al., 2002)

- ▶ Simulations reasonably reproduce the meso-scale downdraft associated with evaporation
- ▶ However, even if DSD profiles are well reproduced, evaporation is underestimated because RH is overestimated in the model
- ▶ Could impact the buoyancy of the atmosphere, cold pool intensity or the dynamics of the system
→ future work

Conclusions

Tridon et al., 2019, MWR
Planche et al., 2019, MWR

- ▶ Difficulties to reproduce the transition zone with WRF using either the Morrison or the Thompson scheme
- ▶ Recently developed retrieval technique providing both N_0^* and D_m at high resolution
- overall reasonable agreement in the DSD profiles
- ▶ in the Morrison scheme, the breakup efficiency had to be increased in order to get better agreement with the observations for this case study
- ▶ Despite the small DSD discrepancies, the evaporation rate is significantly underestimated in the model because it cannot reproduce the especially low observed RH
- How can this impact the buoyancy of the atmosphere, cold pool intensity or the dynamics of the system?
- ▶ We need more, and more persistent profiling observations with multi-frequency cloud radars

Thanks for your attention

Questions?