

CO₂ profiling by space-borne Raman lidar

ISTP 2019,
Toulouse,
20-24 May 19

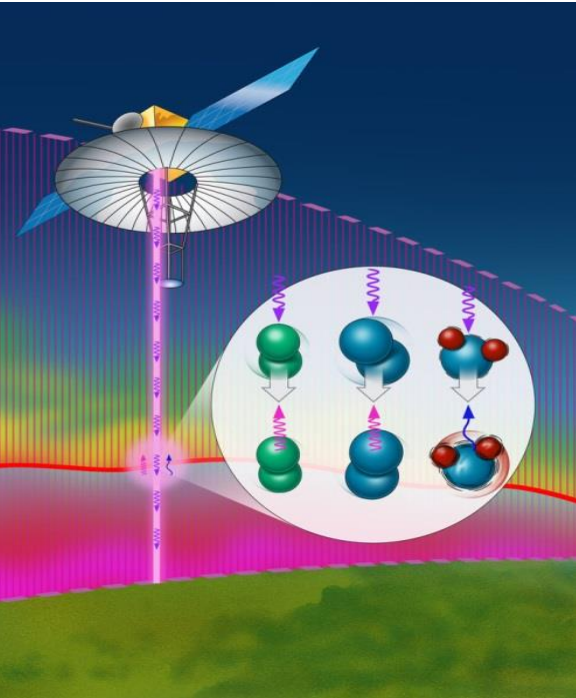
Paolo Di Girolamo¹, Volker Wulfmeyer², Andreas Behrendt²,
Carmine Serio¹, Davide Dionisi³

¹Scuola di Ingegneria, Università della Basilicata, Potenza, Italia

²Institut fuer Physik und Meteorologie, Universitaet Hohenheim, Stuttgart, Germany

³ISMAR-CNR, Roma, Italy

Motivation



- CO₂ mixing ratio in the atmosphere has substantially increased from values around 300 ppm in the fifties of previous century to a current value of 415-420 ppm, with an annual increasing rate of approximately 2 ppm.
- Approximately 50 % of CO₂ amount produced through fossil fuel combustion and other human activities is injected in the atmosphere and accumulates in it, while the remaining 50 % is absorbed by the oceans and the terrestrial biosphere.
- Forests cover ~30% of the Earth's global land area and store large amounts of carbon captured from the atmosphere.

• In order to properly quantify this sink mechanism and its contribution to the carbon cycle, accurate measurements of the CO₂ gradients between the forest floor and the top of the canopy, and their temporal variations, are urgently needed.

• This ultimately translates into the capability to perform accurate and high vertical and horizontal resolution measurements of CO₂ mixing ratio profiles.



Although the **space** and **ground network** for **CO₂ monitoring** has regularly expanded over the past 50 years, it **does not guarantee** the necessary **spatial** and **temporal resolution** needed for a **quantitative analysis** of **sources** and **sinks**.

Space sensors provide CO₂ measurements **above forest canopies**, which do not allow to properly estimate **Gross Primary Production (GPP)**.

These observational gaps could be addressed with an **active remote sensing system** in space based on the **roto-vibrational Raman lidar technique**.

Raman lidar technique for CO₂ profiling

Despite its important potential scientific impact, the CO₂ vibrational Raman lidar technique has received **limited attention** in the last **25 years** both at theoretical and experimental level, mainly because of the **limited precision** that this technique could guarantee, due to the **limited cross-sections** and to the **low concentration** of the **sounded species**.

However, in the last decade, considerable **technological progress** has been achieved in the **design and development** of **solid-state laser sources** with **high optical power**, **large-aperture telescopes** and **high-gain and quantum efficiency detectors**.

This **technological progress** allows today reaching a **new level of performance** in the **Raman lidar measurements** of CO₂.

In a **Raman system**, **CO₂ profile measurements** are possible, together with **simultaneous measurements** of the **temperature and water vapour mixing ratio profile** and a **variety of additional variables** (**aerosol backscatter and extinction profiles**, **PBL depth**, **cloud top and base heights**, **cloud optical depth**).

Observational requirements for CO₂ profiling

In order to properly **design** and **size** a **Space Raman lidar**, **observational requirements** have to be assessed and defined.

Vertical variability in CO₂

Seasonal and **annual mean** of **CO₂ vertical profiles** reflect the **combined influences** of **surface fluxes** and **atmospheric mixing**.

During the **summer** in the Northern Hemisphere, **atmospheric CO₂ concentrations** are generally **lower** near the surface than in the **free troposphere**, reflecting the greater impact of **terrestrial photosynthesis** over **industrial emissions** at this time (Stephens et al., 2007).

Conversely, during the **winter**, **respiration** and **fossil-fuel sources** lead to **high low-altitude** atmospheric **CO₂ concentrations** at northern locations. The **gradients** are **comparable** in magnitude in the **two seasons**, this being of the order of **10 ppm**.

CO₂ gradients between the **forest floor** and the **top of the canopy** are in the range **75-100 ppm** in **daytime** and **10-50 ppm** at **night**. Both ranges represented spring conditions when canopy leaf area development is not completed.

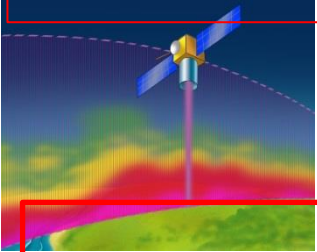
The **Space Raman lidar** has to be **designed** and **sized** with the goal of allowing **measurements** of these **gradients**.



- Vertical resolution: **200 m**
- Horizontal resolution : **50 km**
- Accuracy: **2-5 ppm**

Mission Characteristics

The **lidar setup** considered in the present research effort heavily relies on the **mission concept** of **ATLAS** proposed to the European Space Agency in response to the Call for Explorer-10 Mission Ideas. Simulations consider a **sun-synchronous low Earth orbit**, with an **orbiting height** and **speed** of **450 km** and **7 km/s**, respectively. A **dawn-dusk orbit** with **overpasses** at **6/18 h local time** has been selected for the simulations.



Instrumental concept benefits from **recent advances** in **solid-state laser**, **large-aperture telescope**, and **detector technologies**.

- **frequency-tripled diode-laser pumped Nd:YAG laser**
- average power of **250 W** at 355 nm
- **new generation** of **pump chambers** and **diode lasers** (demon. under dev. at UHOH)
- **electrical-to-optical efficiency** **> 5 %** (by improving power supply and diode laser efficiencies, reducing electrical losses, optimizing pump chambers and the laser conversion efficiency).
- **Inclusion** of **several amplification stages** (3-4), each one embedding **high-density stacks** of **pumping diodes**
- Use of **diode laser pumping** determines **radiative cooling** to be sufficient.
- **Far less complex** than, e.g., for **ADM** or **EarthCARE**. Does not require:
 - frequency stabilization unit,
 - beam shaping optics,
 - specific operational modes (repet. rates and bursts).

Transmitter



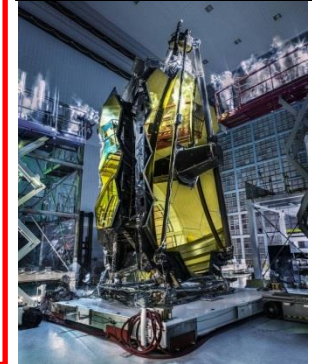
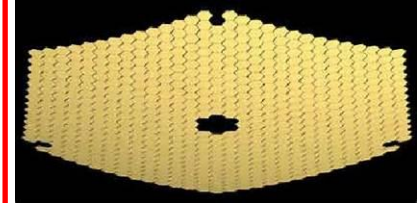
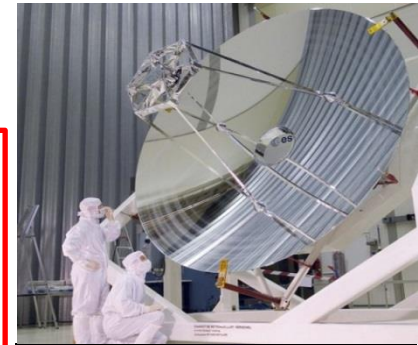
Slab amplifiers with single pulse energies of 12 J (5 times higher) were demonstrated in Liu et al. (2017). A pulse energy >100 J in Mason et al. (2017).

Mission Characteristics

Receiver



- large-aperture, lightweight telescope, 4 m diameter
- Different technological solution:
 - rigid primary mirror (single physical element or segmented optics),
 - Rigid central mirror with folded deployable outer sections
 - inflatable optics
- several glass materials (e.g. Zerodur, SiC, etc.) with appropriate low weight and thermal stability properties for this type of space application
- no astronomical quality needed (no diffraction-limited performances ($swe < \lambda/14$), surface wavefront error $< \lambda$ RMS)

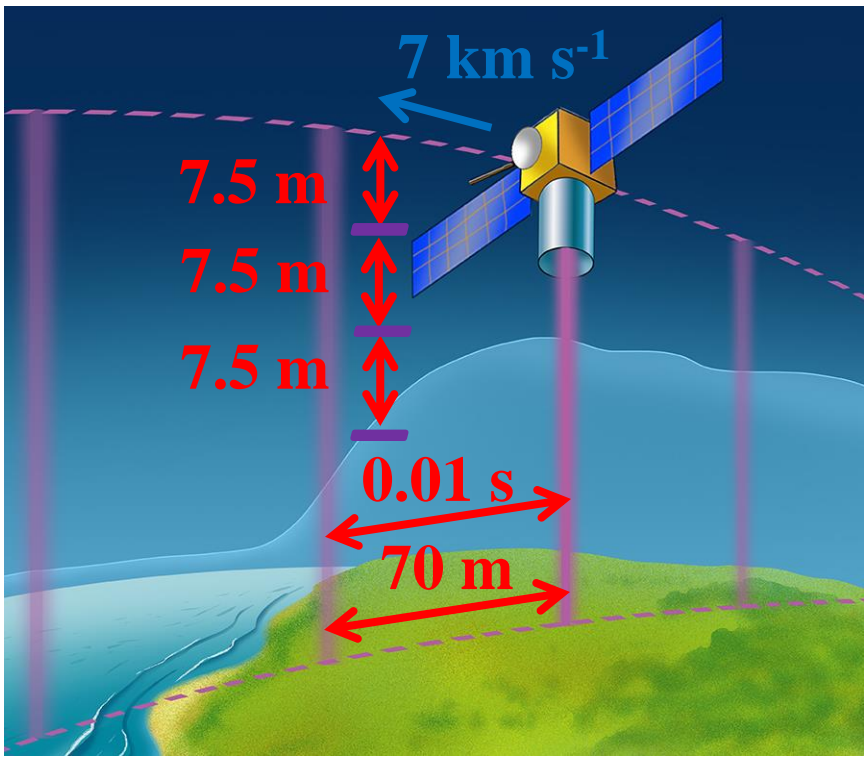


Large aperture primary mirrors (with a total surface of $\sim 10 \text{ m}^2$), with adequate rigidity, low weight (primary mirror areal density $\sim 15 \text{ kg/m}^2$), high wave-front quality ($< \lambda/3$) and sufficient temporal and thermal stability, have been demonstrated (Mazzinghi *et al.*, 2006) to be developable based on the use of segmented mirrors, including a very rigid carbon-fibre composite back-plane and a thin Zerodur glass shell, supported by a set of high efficiency electromagnetic, actively-controlled actuators [102].

- receiving field-of-view = $25 \mu\text{rad}$

• In the present mission concept, the Raman lidar collects **five primary lidar signals**:

- CO₂ vibrational Raman signal $P_{CO_2}(z)$
- water vapour roto-vibrational Raman signal $P_{H_2O}(z)$
- O₂-N₂ high- and low-quantum number rotational Raman signals $P_{loJ}(z) \& P_{HiJ}(z)$
- elastic backscatter signal @ λ_0 $P_{\lambda_0}(z)$



Horizontal and vertical signal averaging to reduce signal statistical uncertainties.

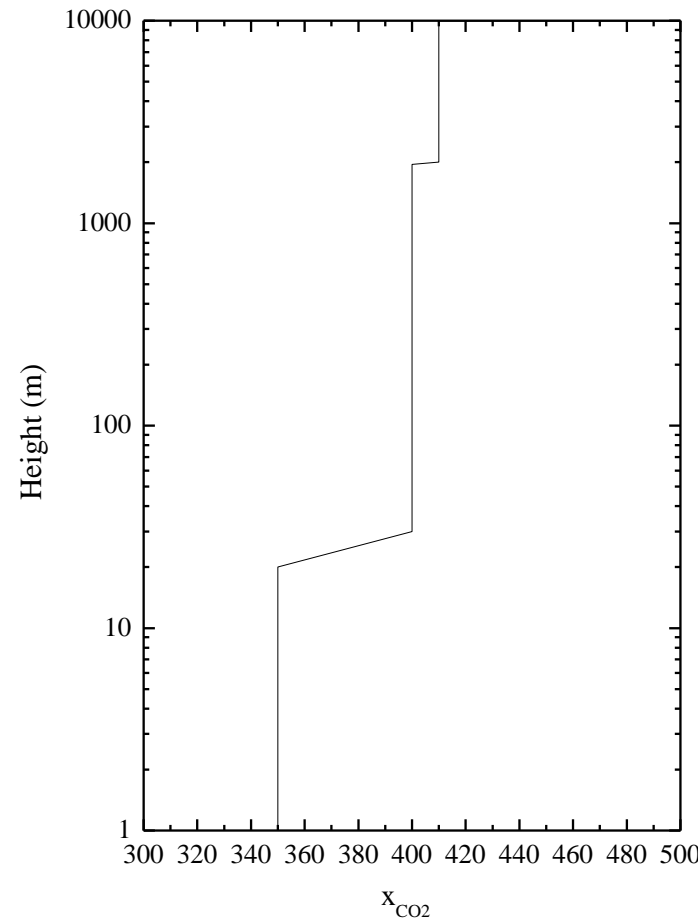
Vertical resolution $\rightarrow \Delta z = 200 \text{ m}$
 Horizontal resolution $\rightarrow \Delta h = 50 \text{ km}$

obtained with an time integration of 7.14 s

An Input CO₂ mixing ratio profile data for the Simulator

was generated which accounts **for different vertical gradients** associated with the conflicting contributions of **terrestrial photosynthesis** and **industrial emissions**, and **CO₂ capturing within forest canopy**. This input profile includes a :

- **10 ppm** increase at an **altitude of 2 km** from a value of **400 ppm** to one of **410 ppm**, introduced in order to simulate the daytime **CO₂ depletion** taking place within the **mixed layer**. In addition to this;
- **50 ppm** decrease is considered at an **altitude of 50 m above the surface level**, which is intended to represent **CO₂ capturing within forest canopy**.
- Other atmospheric quantities considered in the simulation include:
 - **vertical profiles** of **pressure, temperature, and humidity** from the **U.S. Standard Atmosphere 1976** atmospheric reference model;
 - **Aerosol optical properties** from the **median aerosol extinction data** from the **ESA ARMA Model**.



The **Space Raman lidar** has to be designed and sized with the goal of performing **measurements** allowing to **resolve** these **gradients**.



- Vertical resolution: **200 m**
- Horizontal resolution : **50 km**
- Accuracy: **2-5 ppm**

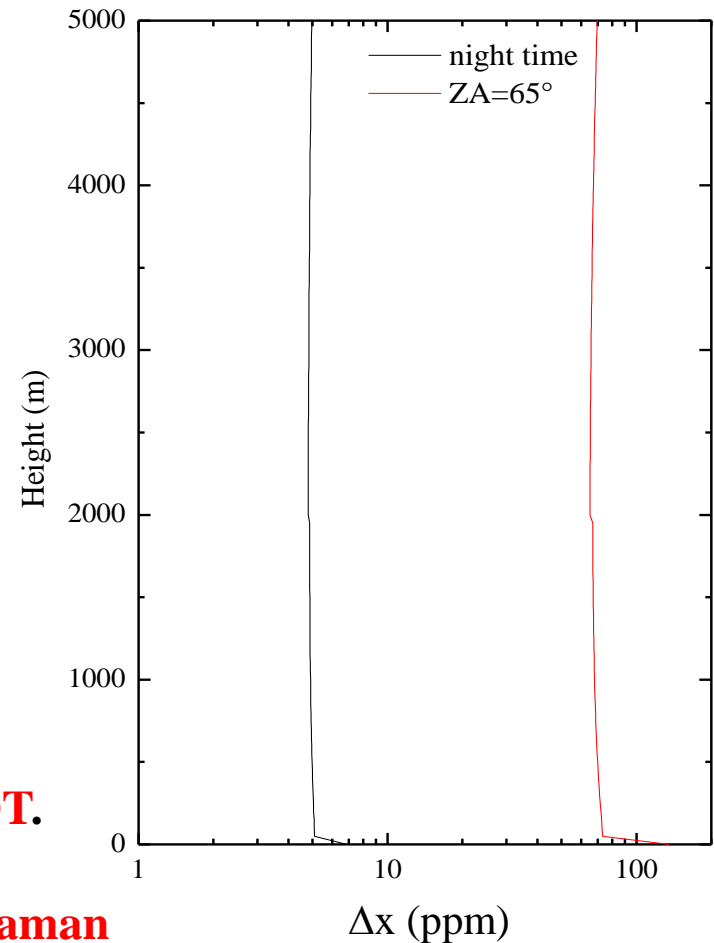
Simulation results

An **assessment** of the expected **system performance** has been performed based on the application of an **analytical simulation model** developed at **University of Basilicata**.

Considering a **vertical** and **horizontal resolution** of **200 m** and **50 km**, respectively, the **statistical uncertainty** (precision) affecting **CO₂ mixing ratio** profile measurements is not exceeding **5 ppm** at **night** and **60 ppm** in **daytime** from the **surface** up to an **altitude of 5 km**.

Night-time performance is acceptable, but **daytime** is **NOT**.

However, so far we have been exploiting only the Raman signal in the anti-Stokes $2\nu_2$ vibrational band.



CO₂ Spectroscopy

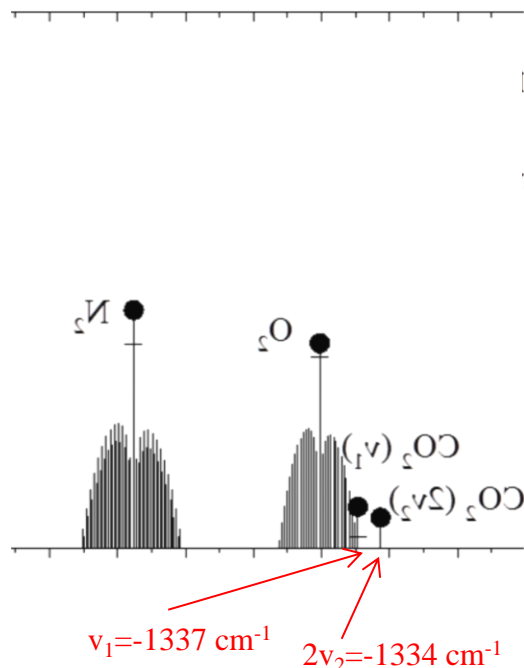
Fermi resonance results in the **splitting** of **two vibrational bands** that have **nearly the same energy** and **symmetry** in both IR and Raman spectroscopies. The two bands are usually a fundamental vibration and either an overtone or combination band.

As a result, **two strong bands** are observed in the **spectrum**, instead of the expected **strong** and **weak bands**. It is not possible to determine the contribution from each vibration because of the resulting mixed wave function.

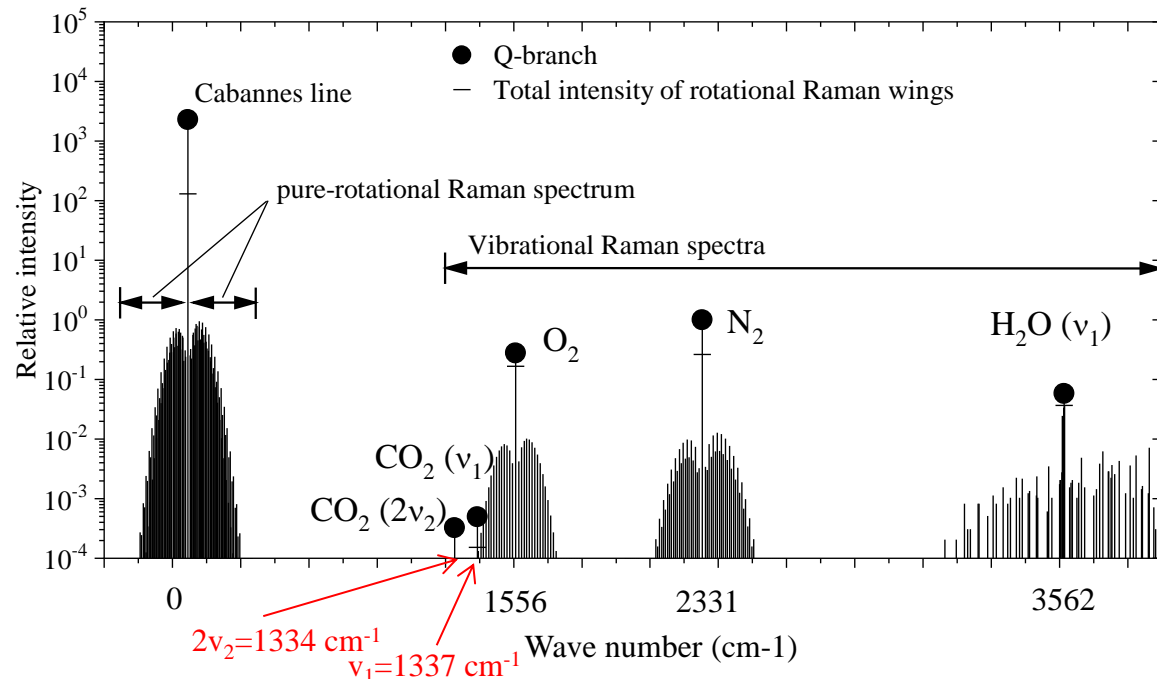
The **three fundamental vibrations** of CO₂ are $\nu_1=1337\text{ cm}^{-1}$, $\nu_2=667\text{ cm}^{-1}$, $\nu_3=2349\text{ cm}^{-1}$.

As a result of Fermi resonance, **two strong bands** are present, with frequencies $\nu_1=1337\text{ cm}^{-1}$ and $2\nu_2(2 \times 667\text{ cm}^{-1})=1334\text{ cm}^{-1}$.

Stokes branch



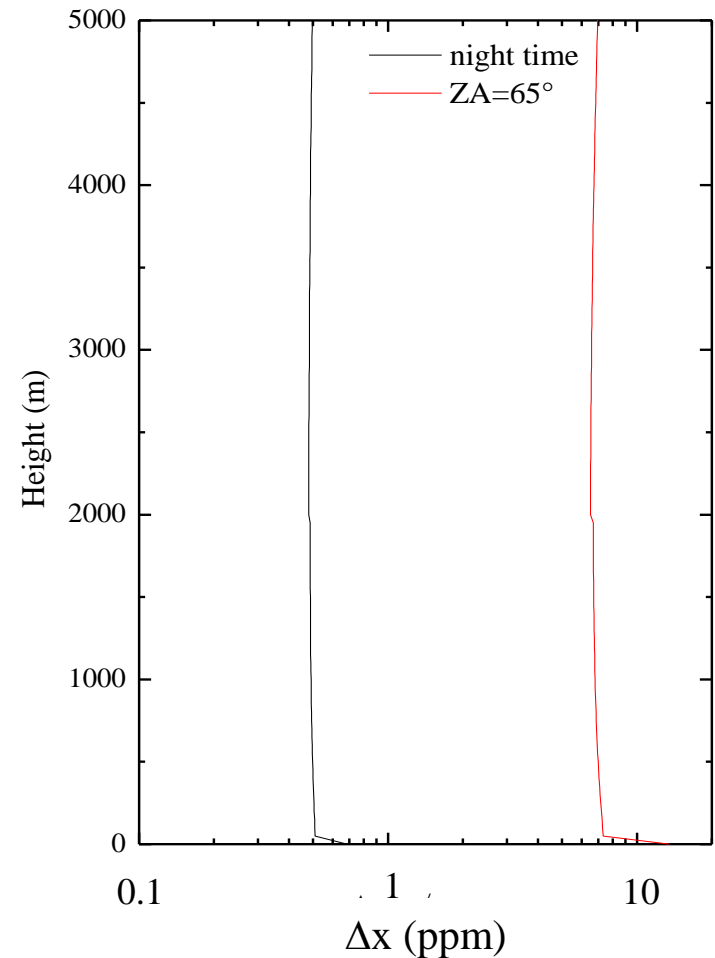
Anti-Stokes branch



Thus, measuring the **Raman lidar echoes** from the **two bands ν_1 and $2\nu_2$** both in the **Stokes** and **anti-Stokes branches**, the intensity of the **CO₂ Raman lidar signal** can be increased by a factor of **4 to 5**.

This solution can be implemented in combination with a reduction of the **receiving field-of-view** from **25 μrad** to **10 μrad** .

This translates into a **statistical uncertainty (precision)** affecting **CO₂ mixing ratio** profile measurements not exceeding **0.5 ppm** at **night** and **5 ppm** in **daytime** from the **surface** up to an **altitude of 5 km**.



Measured parameters/expected performance

Clear-sky conditions

Accuracy estimated with experimentally-validated performance model (Di Girolamo et al., 2006)

Variable	Resolution (vertical/horizontal)	Precision
Water vapor mixing ratio	200 m / 50 km	2-20 %
Temperature	200 m / 50 km	0.4-0.75 K
Relative humidity	200 m / 50 km	2-20 %
Planetary boundary layer height	100 m / 5 km	100 m
Particle load, water vapor extinction	50-100 m / 10-50 km	1-3 % & 3-20 %
	50-100 m / 5 km	50-100 m
	50 km	5 %

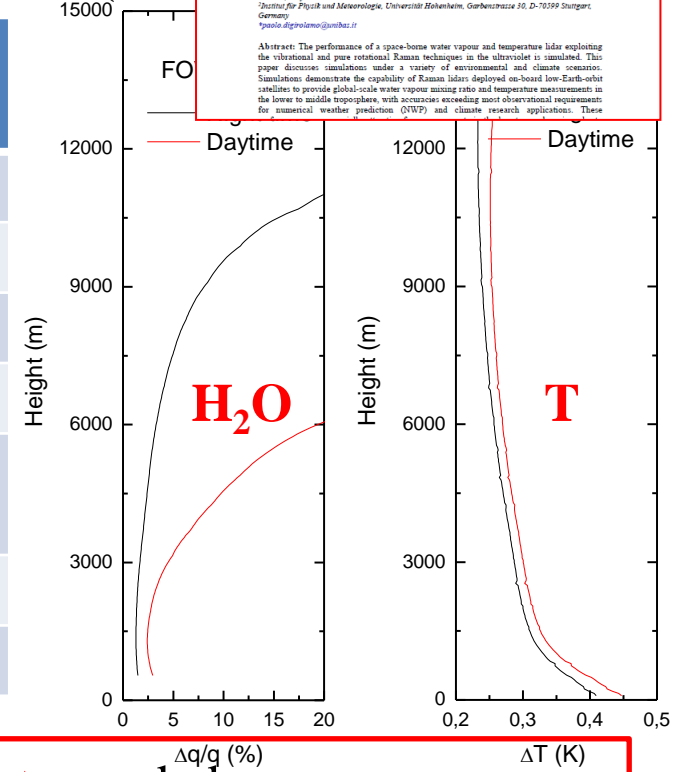
Research Article
Optics EXPRESS
VOL. 26, NO. 7, 2 APR 2018 | OPTICS EXPRESS 8128

Space-borne profiling of atmospheric thermodynamic variables with Raman lidar: performance simulations

PAOLO DI GIROLAMO,^{1*} ANDREAS BEHRENDT,² VOLKER WULFMEYER³

¹Scienze di Ingegneria, Università degli Studi della Basilicata, P.zza dell'Ateneo Lucano n. 10, 85100 Potenza, Italy
²Institut für Physik und Meteorologie, Universität Hohenheim, Garbenstrasse 30, D-70559 Stuttgart, Germany
³paolo.digirolamo@unibas.it

Abstract: The performance of a space-borne water vapour and temperature lidar exploiting the vibrational and pure rotational Raman techniques in the ultraviolet is simulated. This paper discusses simulations under a variety of environmental and climate scenarios. Simulations demonstrate the capability of Raman lidars deployed on-board low-Earth-orbit satellites to provide global-scale water vapour mixing ratio and temperature measurements in the lower to middle troposphere, with accuracies exceeding most observational requirements for essential weather prediction (EWP) and climate research applications. These



Spaceborne profiling of atmospheric temperature and particle extinction with pure rotational Raman lidar and of relative humidity in combination with differential absorption lidar: performance simulations

Paolo Di Girolamo, Andreas Behrendt, and Volker Wulfmeyer

The performance of a spaceborne temperature lidar based on the pure rotational Raman (RR) technique in the UV has been simulated. Results show that such a system deployed onboard a low-Earth-orbit satellite would provide global-scale clear-sky temperature measurements in the troposphere and lower atmosphere with precisions that satisfy World Meteorological Organization (WMO) threshold observational requirements for essential weather prediction and climate research applications. Furthermore, nighttime temperature measurements would still be within the WMO threshold observational requirements in the presence of several cloud structures. The performance of aerosol extinction measurements from space, which can be carried out simultaneously with temperature measurements by RR lidar, is also assessed. Furthermore, we discuss simulations of relative humidity measurements from space obtained from RR temperature measurements and water-vapour data measured with the differential absorption lidar (DIAL) technique. © 2006 Optical Society of America
 OCSZ code: 010.1210, 120.0120, 380.3040.

1. Introduction
 Atmospheric temperature is a key parameter characterizing the state of the atmosphere at any given time and location. The atmospheric temperature field shows a large variability on both the horizontal and the vertical scale, as well as on the time scale, and its variability is difficult to predict because of the highly variable distribution of radiative forcing factors. Among these are clouds and water vapor.
 An adequate comprehension of meteorological processes and climate phenomena requires accurate measurements of atmospheric temperature with global coverage and high temporal and spatial resolution. For climate research, accurate measurements of temperature trends are essential not only at the Earth's surface but also in the entire troposphere and stratosphere. For weather research, accurate temperature profiling is crucial for investigating atmospheric stability and for the detection of regions with high convective available potential energy (CAPE) where initiation of convection can take place. Observational requirements to be fulfilled by networks of satellite remote sensors have been defined by the World Meteorological Organization (WMO)¹ and by the World Climate Research Program (WCRP), and are listed in the Committee on Earth Observation Satellites (CEOS WMO) on-line database² (Table 1). Threshold requirements for temperature observations specified by global and regional numerical weather prediction (NWP), synoptic meteorology, and nowcasting imply globally distributed measurements of atmospheric temperature throughout the troposphere with an accuracy of 2 K, with an observing cycle of 1 h and a vertical and horizontal resolution of 1 and 200 km, respectively. Additionally, threshold requirements identified by SPARC (Stratospheric Processes and their Role in Climate) and global climate modeling imply atmospheric temperature measurements up to the lower stratosphere with an accuracy of 1 K, with an observing cycle of 1 h and a vertical and horizontal resolution of 2 and 500 km, respectively. Global observations fulfilling these requirements will have a positive effect on atmospheric

i.e.,
/day
200

- No ancillary data needed
- Errors provided with each retrieved profile
- Clear air, above clouds, through broken clouds & through/below overcast thin clouds

Summary

- An accurate quantitative assessment of the various components of the carbon cycle requires accurate measurements of the different sources and sinks.
- For the purpose of estimating forests' carbon capturing capabilities, accurate measurements of CO₂ gradients between the forest floor and the top of the canopy are needed, which ultimately translates into the capability to measure CO₂ concentration profiles.
- Simulations reveal that a space-borne Raman lidar based on the roto-vibrational technique, if properly conceived (i.e. exploiting both ν_1 and $2\nu_2$ vibrational bands in both the Stokes and anti-Stokes branches), may provide CO₂ mixing ratio profile measurements with the level of accuracy needed to quantitatively assess the different sources and sinks
- Ground-based demonstrators of the present instrument concept are under development both at Univ. of Basilicata and Univ. of Hohenheim.

carbon dioxide measurements based on the application of the Raman lidar technique have been carried out since the early nineties (Riebesell, 1990; Ansmann et al., 1992). The technique is based on the measurement of anelastically retro-diffused laser radiation (retro-diffusion Raman roto-vibrational) from the carbon dioxide molecules present in the atmosphere.

Despite its important potential scientific applications, the Raman lidar technique for the measurement of carbon dioxide has received little attention in the last 25 years both at theoretical and experimental level, mainly due to the limited precision that characterizes this measure, due to the limited section d impact of the scattering phenomenon on which the technique is based and the low concentration of the species of interest.

Misure a carattere sperimentale del contenuto atmosferico dell' anidride carbonica basate sull'impiego della tecnica lidar Raman sono state realizzate sin dai primi anni novanta (Riebesell, 1990; Ansmann *et al.*, 1992). La tecnica si basa sulla misura della radiazione laser retro-diffusa anelasticamente (retro-diffusione Raman roto-vibrazionale) dalle molecole di anidride carbonica presenti in atmosfera. Nonostante le sue importanti potenziali applicazioni scientifiche, la tecnica lidar Raman per la misura dell' anidride carbonica ha ricevuto negli ultimi 25 anni poca attenzione sia a livello teorico che sperimentale, principalmente a causa della limitata precisione che caratterizza questa misura, riconducibile alla limitata sezione d'urto del fenomeno di scattering su cui la tecnica si basa ed alla bassa concentrazione della specie d'interesse.

Ad ogni modo, nell'ultimo decennio sono stati conseguiti enormi progressi nella progettazione e nello sviluppo sperimentale di sorgenti laser a stato solido con elevata potenza, di telescopi a grande apertura e di rivelatori ad elevato guadagno ed efficienza quantica, progressi che consentono di poter raggiungere oggi un nuovo livello di prestazioni nelle misure lidar Raman dell' anidride carbonica.

Per stimare le prestazioni potenziali di un sistema lidar è necessario far uso di un simulatore. A tale scopo, la precisione delle misure lidar Raman del rapporto di mescolamento dell'anidride carbonica, sia in orario diurne che notturno, è stata stimata mediante l'impiego di un simulatore sviluppato presso la Scuola di Ingegneria dell'Università della Basilicata (Di Girolamo *et al.*, 2006). La grandezza misurata dal lidar Raman, sia nel caso del vapore acqueo (Whiteman *et al.*, 1992) che della CO_2 (Ansmann *et al.* 1992), è il rapporto di mescolamento rispetto all'aria secca. Per poter quantificare questa grandezza si fa uso della misura simultanea della retrodiffusione Raman da parte dell' azoto molecolare per normalizzare il segnale Raman retro-diffuso dalle molecole di vapore acqueo o di anidride carbonica.

Le simulazioni sono state effettuate assumendo una integrazione temporale di misura (risoluzione temporale) di 1 ora ed una risoluzione verticale di 75 m sotto 1.25 km, di 150 m nella regione di quote 1.25-2.0 km, di 250 m nella regione di quote 2.0-2.5 km, di 400 m nella regione di quote 2.5-3.0 km e di 600 m sopra 3.0 km.

I dati di ingresso del modello includono un profilo del rapporto di mescolamento dell'anidride carbonica caratterizzato da un aumento di 10 ppm ad un'altezza di 2.2 km (da un valore di 350 ad uno di 360 ppm), introdotto per poter simulare la diminuzione della CO₂ che si verifica al tetto dello strato misto durante il corso del giorno. Questo profilo simula una possibile condizione poco dopo il tramonto. Il carico aerosolico è stato simulato assumendo la presenza di uno strato con un coefficiente di estinzione costante e pari a 0.05 km⁻¹ nei primi 2 km di quota. Le simulazioni considerano un sistema lidar, con una sorgente laser di tipo Nd:YAG in grado di emettere una potenza ottica nell'UV (354.7 nm) di 30 W (energia di singolo impulso=3 J, frequenza di ripetizione=100 Hz) ed un telescopio per la raccolta dei segnali lidar con uno specchio primario di diametro pari a 0.6 m. Le simulazioni considerano inoltre che la selezione spettrale della banda Raman del CO₂ (Q-branch della banda Raman roto-vibrazionale u₂ della CO₂, che risulta spostata rispetto alla lunghezza d'onda laser di 1285 cm⁻¹) venga effettuata mediante l'impiego di un filtro interferenziale, con le seguenti specifiche: lunghezza d'onda centrale (CWL) = 371.71 nm, larghezza spettrale di banda (BW) = 0.1 nm, trasmissione a CWL = 40%, reiezione fuori banda = OD6 @ 200-1200 nm, OD12 @ 354.7 nm e OD7 @ 375-387 nm. Le simulazioni considerano infine una altezza di volo dell'aereo ospitante il sistema lidar di 4 km.

Le simulazioni evidenziano come la diminuzione del rapporto di mescolamento della CO₂ che si verifica al tetto dello strato misto sia ben risolto utilizzando un sistema lidar Raman dimensionato come sopra. Più specificamente, utilizzando le risoluzioni temporali e verticali sopramenzionate, la precisione della misurazione risulta pari a ~ 2 ppm fino a circa 3 km di quota.

La lunghezza d'onda centrale (371,71 nm) è quasi coincidente con la ventiduesima riga spettrale della banda anti-Stokes dello spettro roto-vibrazionale dell'ossigeno molecolare, specie che rappresenta una potenziale fonte di contaminazione per la misura della CO₂. Riebesell (1990) e Ansmann *et al.* (1992) giunsero alla conclusione che le misure lidar Raman di anidride carbonica fossero difficilmente realizzabili a causa dell'impossibilità di poter stimare in modo accurato l'entità dell'interferenza da parte delle righe rotazionali dell'ossigeno molecolare. Questi autori ipotizzarono inoltre che misure lidar con precisioni dell'ordine del ppm potessero essere contaminate dall'eventuale presenza di fluorescenza generata dalle ottiche del ricevitore o dalle particelle atmosferiche. Tuttavia, queste ricerche erano state condotte utilizzando un laser ad eccimeri (miscela Xe:Cl), caratterizzato da uno spettro di emissione che si estende per circa 0.4 nm. L'emissione di radiazione laser su questo ampio spettro di lunghezze d'onda rende la separazione tra lo scattering Raman di O₂ e CO₂ più difficile rispetto a quella realizzabile mediante l'uso di filtri interferenziali a banda stretta, quali quelli di cui si ipotizza sull'uso in questo prototipo, e di una sorgente laser del tipo Nd: YAG, con larghezza spettrale dell'impulso emesso di ~ 0.02 nm.

Come accennato in precedenza, il Q-branch dello spettro di Raman u2 della CO₂ è quasi coincidente con la ventiduesima riga anti-Stokes dello spettro roto-vibrazionale dell'ossigeno molecolare. Calcoli effettuati sulla base del valore della larghezza di banda del filtro interferenziale considerato per il sistema proposto (0.1 nm) indicano che il contributo di questa riga anti-Stokes dell'O₂ al segnale Raman della CO₂ è inferiore all'1 % (-3-4 ppm). Una opportuna modellazione della variabilità dell'intensità di questa riga rotazionale in funzione della temperatura può consentire di stimare in modo accurato l'ampiezza dell'interferenza in modo che questa possa essere sottratta dal segnale Raman dell'anidride carbonica (Whiteman *et al.*, 2001). Whiteman *et al.* (2001) hanno determinato che l'applicazione di questo approccio può consentire di ridurre a 0.3 ppm l'incertezza della misura del CO₂ causata da questa interferenza.

Nell'ambito di questo progetto si renderà necessario un accurato studio della fluorescenza generata dalle ottiche del ricevitore o dalle particelle atmosferiche. Tuttavia, misure preliminari ottenute utilizzando uno spettrometro accoppiato ad un ricevitore Raman (Whiteman *et al.*, 2001) non hanno evidenziato alcun significativo contributo della fluorescenza nella regione spettrale di pertinenza della banda Raman del CO₂, anche se la fluorescenza dovuta agli aerosol è stata osservata a lunghezze d'onda più lunghe.

Il sistema lidar progettato e sviluppato in forma prototipale nell'ambito di questo progetto sarà in grado di misurare oltre che i profili verticali del rapporto di mescolamento della CO₂, anche i profili verticali del rapporto di mescolamento del vapor acqueo, della temperatura e delle proprietà ottiche del particolato atmosferico (coefficiente di backscattering ed estinzione a 355 nm).

Il segnale lidar Raman generato dalle molecole di anidride carbonica è molto più debole rispetto a quello generato dalle molecole di vapor acqueo, azoto ed ossigeno molecolare (questi ultimi due usati anche per la misura della temperatura atmosferica) e del segnale lidar elastico generato dal particolato atmosferico. Questo fa sì che l'incertezza statistica che caratterizza le misure di queste ulteriori grandezze atmosferiche sia sensibilmente inferiore rispetto a quella che caratterizza la misura dell'anidride carbonica, rendendo quindi possibili misure accurate di questi parametri con tempi di integrazione molto più contenuti.

Referenze

Riebesell, M., 1990: Raman lidar for the remote sensing of the water vapor and carbon dioxide profile in the troposphere (in German). Ph.D. thesis, GKSS document 901/F/13, University of Hamburg, 127 pp.

Ansmann, A., M. Riebesell, C. Weitkamp, E. Voss, W. Lahmann, and W. Michaelis, 1992b: Combined Raman Elastic-backscatter lidar for vertical profiling of moisture, aerosol extinction, backscatter, and lidar ratio. *Appl. Phys.*, 55B, 18-28.

Whiteman, D. N., S. H. Melfi, and R. A. Ferrara, 1992: Raman lidar system for the measurement of water vapor and aerosols in the earth's atmosphere. *Appl. Opt.*, 31, 3068-3082.

Whiteman, D. N., G. Schwemmer, T. Berkoff, H1, Plotkin, L. Ramos-Izquierdo, and G. Pappalardo, 2001: Performance modeling of an airborne Raman water vapor lidar, *Appl. Opt.*, 40,375-390.

atmosfera del rapporto di
mescolamento dell' anidride carbonica.
Nell'ambito dell'OR4 si intende altresì
avviare lo sviluppo prototipale di
questo sistema, nonché realizzare studi
per l'ingegnerizzazione e la qualifica
per volo aereo del sistema.

A tal riguardo si specifica che, nonostante le sue importanti potenziali applicazioni scientifiche la tecnica lidar Raman roto-vibrazionale per la misura dell' anidride carbonica ha ricevuto negli ultimi 25 anni poca attenzione sia a livello teorico che sperimentale, principalmente a causa della limitata precisione che questa tecnica era in grado di garantire, riconducibile alla limitata sezione d'urto del fenomeno di scattering Raman roto-vibrazionale ed alla bassa concentrazione della specie d'interesse. Nell'ultimo decennio, però, notevoli progressi sono stati conseguiti nella progettazione e sviluppo sperimentale di sorgenti laser a stato solido con elevata potenza ottica, di telescopi a grande apertura e di rivelatori ad elevato guadagno ed efficienza quantica, progressi che consentono oggi di poter raggiungere un nuovo livello di prestazioni nelle misure lidar Raman dell' anidride carbonica. Fruttando questi progressi, nell'ambito del progetto si intende verificare l'effettiva fattibilità e quindi progettare un sistema lidar basato su questa tecnica,

By exploiting these advances, the project intends to verify the actual feasibility and therefore to design a lidar system based on this technique.

Il cambiamento climatico in atto (e.g., *Working Group I Contribution to the IPCC fifth Assessment Report, Climate Change 2013, The Physical Science Basis*, <http://www.ipcc.ch/report/ar5/wg1/>) ed il relativo riscaldamento globale sono conseguenza di un'alterazione dell'effetto serra naturale del pianeta. L'effetto serra naturale è dominato per oltre il 75% dal vapore acqueo e per il restante 25% principalmente dalla CO₂. Mentre le attività antropiche aggiungono solo *poche gocce* al ciclo idrologico della Terra, esse hanno sostanzialmente alterato quello del carbonio: la concentrazione della CO₂ in atmosfera è passata dai circa 300 ppmv degli anni '50 dello scorso secolo, agli attuali 405-410 ppmv del 2017, con un incremento annuo che sta superando la stima di 2 ppmv.

La combustione di combustibili fossili ed altre attività umane immettono nell'atmosfera circa **40 miliardi tonnellate di CO₂ ogni anno**. Le osservazioni della CO₂ e della sua concentrazione in atmosfera sono principalmente eseguite attraverso una rete globale di stazioni a terra. Le misure indicano che **meno della metà della CO₂ immessa si accumula nell'atmosfera**. La restante parte è evidentemente assorbita dagli **oceani** e dalla **biosfera terrestre** (e.g. Le Quere et al 2009, 2013). Sebbene la rete di monitoraggio al suolo si sia regolarmente espansa negli ultimi 50 anni, consentendo di avere un quadro accurato della variabilità superficiale della concentrazione della CO₂ a scala globale, essa **non garantisce la dovuta risoluzione spaziale e temporale necessaria per analisi quantitative di sorgenti (sources) e pozzi (sinks)** ed, in ultima analisi, non è in grado di pervenire al bilancio biogenico connesso con l'attività di fotosintesi clorofilliana della vegetazione (Gross Primary Productivity o GPP).

Allo stato attuale esiste una carenza notevole in relazione alle osservazioni di come l'attività di fotosintesi clorofilliana e la respirazione della vegetazione risponde, a scala locale, all'incremento di temperatura e della CO₂, e ai crescenti periodi di siccità, come quelli verificatisi, e.g., nell'anno 2017. La GPP è direttamente connessa con la salute delle vegetazione, il suo stato di stress, ed in generale tutti quegli aspetti legati alla fenologia delle piante.

Questa problematica è di particolare rilevanza al nostro Paese, la cui superficie boschiva occupa circa il 35% del territorio con incendi e fenomeni franosi sempre più frequenti.

Our current knowledge about atmospheric CO₂ concentrations and surface fluxes at regional scales over the globe comes primarily from ground-based in situ measurements of air sampling networks and tall towers (Reuter et al., *Atmos. Meas. Tech.*, 5, 1349–1357, 2012).

These measurements are used by assimilation systems like NOAA's (National Oceanic and Atmospheric Administration) CarbonTracker (Peters et al., 2007, 2010), modeling global distributions of atmospheric CO₂ mixing ratios and surface fluxes. Therefore, within this publication, we consider CT2010 (Carbon-Tracker version 2010) as current knowledge and reasonable a priori estimate for atmospheric CO₂ concentrations (Reuter et al., *Atmos. Meas. Tech.*, 5, 1349–1357, 2012). However, due to the sparseness of measurements, there are still large uncertainties especially on the surface fluxes (Stephens et al., 2007).

Coupling

Coupling between the Amazonian carbon cycle, global climate and global greenhouse gas burdens of CO₂, CH₄ and N₂O (Gatti et al., *Tellus* (2010), 62B, 581–594).

Previous lidar measurements

A 1.6 m differential absorption Lidar (DIAL) system for measurement of vertical CO₂ mixing ratio profiles was been developed. (Yasukuni Shibata et al., *Sensors* 2018, 18, 4064; doi:10.3390/s18114064).

CO₂ vertical profiles in the 5–25 km altitude range with an accuracy of about 2 ppm based on Limb measurements

Major limitations of our present knowledge of the global distribution of CO₂ in the atmosphere are the uncertainty in atmospheric transport and the sparseness of in situ concentration measurements. Limb viewing space-borne sounders such as the Atmospheric Chemistry Experiment Fourier transform spectrometer (ACE-FTS) offer a vertical resolution of a few km for profiles, which is much better than currently flying or planned nadir sounding instruments can achieve (P. Y. Foucher, *Atmos. Chem. Phys.*, 11, 2455–2470, 2011).

Seasonal swing ranged between 20 and 35 ppm for CO₂ mixing (CHUN-TA LAI, Tellus, 58B, 523–536, 2006).

Studies about current spatial and temporal variations of CO₂ in forest canopies provide critical information about how well a forest is coupled to the convective boundary layer (CBL) above the canopy, and therefore how susceptible this ecosystem is to increased atmospheric CO₂ (Buchmann et al., Global Change Biology, **2**, 421-432, 1996).

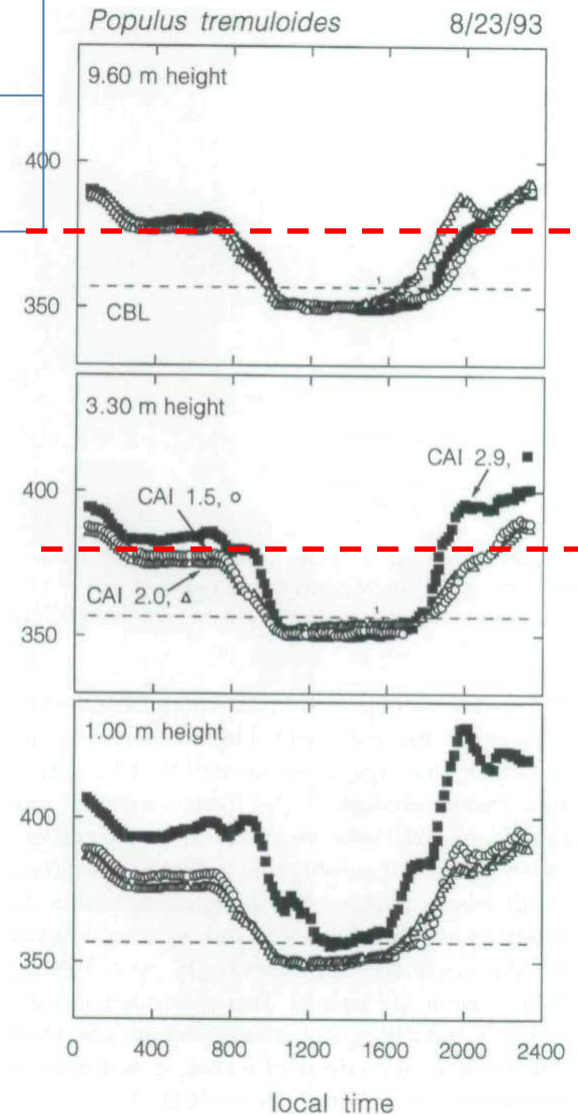
CO₂ gradients between the forest floor and the top of the canopy ranged between 75 and 100 ppmv at night, compared to 10-50 ppmv during the day. Both ranges represented spring conditions when canopy leaf area development had not been completed (Buchmann et al., Global Change Biology, **2**, 421-432, 1996).

Despite the lack of tree foliage and understorey vegetation, it is found a 5 ppmv canopy gradient during the day, but an average 23 ppmv gradient at night. (Buchmann et al., Global Change Biology, **2**, 421-432, 1996)

During June and July, [CO₂] in the upper and lower canopy (3.30-14.0 m) dropped below those of the CBL, depleting canopy air by 1-2 ppmv in June, and by 4-11 ppmv in July. Lowest canopy CO₂ were reached in July, before they increased again in fall (Buchmann et al., Global Change Biology, **2**, 421-432, 1996).

	Pinus contorta	Populus Tremuloides	Red Butte Canyon Acer spp
Canopy height	13-15 m	9-13 m	13-15 m

(Buchmann et al., Global Change Biology, 2, 421-432, 1996)



As clearly reported in the IPCC fifth Assessment Report, CO₂ emissions are already producing destructive effects to the plant ecosystem through the alteration of soil-atmosphere interaction mechanisms.

Although the space and ground network for CO₂ monitoring has regularly expanded over the past 50 years, it does not guarantee the necessary spatial and temporal resolution needed for a quantitative analysis of sources and sinks.

For the purpose of estimating forests' carbon capturing capabilities, accurate measurements of CO₂ gradients between the forest floor and the top of the canopy, which ultimately translates into the capability to measure CO₂ concentration profiles. Space sensors provide CO₂ measurements above forest canopies, which do not allow to properly estimate Gross Primary Production (GPP). These observational gaps could be addressed with an active remote sensing system in space based on the roto-vibrational Raman lidar technique. CO₂ profile measurements are possible, together with simultaneous measurements of the temperature and water vapour mixing ratio profile and a variety of additional variables (aerosol backscatter profile, aerosol extinction profile, PBL depth, cloud top and base heights, cloud optical depth). An assessment of the expected performance of the system has been performed based on the application of an analytical simulation model developed at University of Basilicata.

Space sensors provide CO₂ measurements above forest canopies, which do not allow to properly estimate the Gross Primary Production (GPP).

- The estimated amount of carbon dioxide (CO₂) in the atmosphere is equivalent to an amount of carbon of 8×10^{14} kg, while the amount stored in terrestrial biomass is 5×10^{14} kg, 60% of which (3×10^{14} kg) being stored in forest systems.

N is the number of atoms

number vibrations.

The same holds true for linear molecules, however the equation $3N-5$ is used, because a linear molecule has one less rotational degree of freedom. (For a more detailed explanation see: Normal Modes). Figure 1 shows a diagram for a vibrating diatomic molecule. The levels denoted by vibrational quantum numbers v represent the potential energy for the harmonic (quadratic) oscillator. The transition $0 \rightarrow 1$ is fundamental, transitions $0 \rightarrow n$ ($n > 1$) are called overtones, and transitions $1 \rightarrow n$ ($n > 1$) are called hot transitions (hot bands).

a **observational requirements** have to be properly assessed.

CO₂ vertical variability

The seasonal and annual mean of CO₂ vertical profile reflect the combined influences of surface fluxes and atmospheric mixing. In the Northern Hemisphere, during the summer season, midday atmospheric CO₂ concentrations are generally lower near the surface than in the free troposphere, reflecting the higher impact of terrestrial photosynthesis over industrial emissions [5]. Conversely, during the winter, respiration and fossil-fuel sources lead to elevated low-altitude atmospheric CO₂ concentrations. The gradients are comparable in magnitude in both seasons, but the positive gradients persist for a larger portion of the year, leading to annual-mean gradients also showing higher atmospheric CO₂ concentrations near the surface than aloft [5]. The average midday differences in the Northern Hemisphere between 1 and 4 km, expressed in terms of CO₂ mixing ratio, is -2.2 ppm in summer, +2.6 ppm in winter, while instantaneous differences may be as large as 5-10 % [9]. The Southern Hemisphere locations show relatively constant CO₂ profiles in all seasons, with slightly higher values in the free troposphere [5]. Similar gradients are found in the NOAA's Carbon Tracker version 2010 [4]. CO₂ gradients between the forest floor and the top of the canopy range between 75 and 100 ppm at night, compared to 10-50 ppm during the day. Both ranges represented spring conditions when canopy leaf area development is not completed [10].

The central wavelength of the CO₂ band is almost coincident with the twenty-first anti-Stokes roto-vibrational O₂ line, which represents a potential source of contamination for CO₂ Raman lidar measurements [11]. Riebesell [6] and Ansmann et al. [7] came to the conclusion that CO₂ Raman lidar measurements were hardly feasible due to the difficulties in accurately estimating the magnitude of the interference by O₂ rotational lines. These authors also argued that fluorescence generated by either the receiver optics or atmospheric particles could potentially prevent from reaching a measurement accuracy at the 1 ppm level.

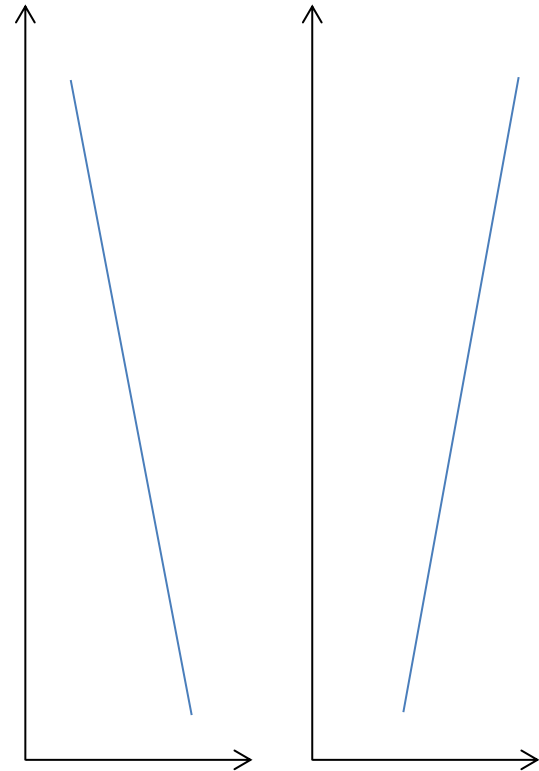
However, these early research efforts were conducted using excimer laser sources (Xe:Cl mixture), which have an emission spectrum typically spanning over a spectral interval of ~0.4 nm. Such broad spectrum makes separation of CO₂ and O₂ lines more difficult with respect to what is presently achievable based on the use of narrowband interference filters (0.1 nm) and injection-seeded Nd:YAG laser sources (typical spectral width of ~0.01 nm). Whiteman et al. [11] demonstrated that the bias affecting CO₂ measurements as a result of the contamination by the 21st O₂ rotational line is not exceeding 1 % if a 0.1-nm wide IF and a narrow-band Nd:YAG laser source (<0.02 nm) are used. Based on an appropriate rotational strength modeling as a function of temperature and relying on independent temperature profile measurements provided by the same Raman lidar system, the magnitude of the interference can be properly estimated and subtracted from the CO₂ signal so that a residual systematic uncertainty of the order of 0.1-0.2 ppm is left in CO₂ mixing ratio measurements.

Observational requirements for CO₂ profiling - continuation

The average Northern Hemisphere midday differences between altitudes of 1 and 4 km is -2.2 ppm in summer, $+2.6$ ppm in winter, and $+0.7$ ppm in annual mean.

The Southern Hemisphere locations show relatively constant CO₂ profiles in all seasons, with slightly higher values in the free troposphere.

Similar gradients are found in the NOAA's (National Oceanic and Atmospheric Administration) Carbon Tracker version 2010 (CT2010).

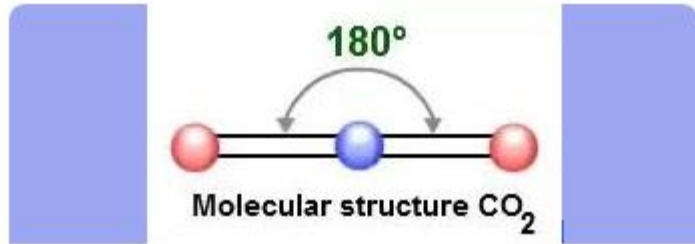


An assessment of the expected performance of the system has been performed based on the application of an analytical simulation model developed at University of Basilicata.

Fermi resonance results in the splitting of two vibrational bands that have nearly the same energy and symmetry in both IR and Raman spectroscopies. The two bands are usually a fundamental vibration and either an overtone or combination band.

As a result, two strong bands are observed in the spectrum, instead of the expected strong and weak bands. It is not possible to determine the contribution from each vibration because of the resulting mixed wave function.

**Sistema lidar Raman per la misura
dei profili verticali in atmosfera
del rapporto di mescolamento dell'
anidride carbonica**



For linear molecules, however the equation $3N-5$ is used, because a linear molecule has one less rotational degrees of freedom.

The Q branches of the Raman bands associated to the $2\nu_1:4\nu_2:\nu_1+2\nu_2$ Fermi resonance of $^{12}\text{C}^{16}\text{O}_2$ have been observed in the gas phase at 2543, 2671, and 2797 cm^{-1} and, in addition, at 2514 cm^{-1} ,

- Alternative option: Alexandrite laser source (BeAl_2O_4 doped with active chromium ions, Cr^{3+})
- Main advantages:
 - increase of laser gain with temperature which considerably simplifies laser cooling demands.
 - Tuneability between 730-780 nm so that desired wavelength can be reached by frequency-doubling (2-times larger efficiency than frequency-tripling)
 - pumped by diode lasers operating in the 500-650-nm region which are now commercially available with high power.

Although the **space** and **ground network** for **CO₂ monitoring** has regularly expanded over the past 50 years, it **does not guarantee** the necessary **spatial** and **temporal resolution** needed for **quantitative analysis** of **sources** and **sinks**.

- The **missing balance** between the **carbon released** in the atmosphere through the combustion of fossil fuels and deforestation on one side, and the **uptake by sinks** in **oceanic** and **terrestrial systems** on the other side is **responsible** for the **2 ppm** increasing rate in atmospheric **CO₂ concentration**.

A proper **definition** of the **CO₂ observational requirements** impose the **assessment** of **CO₂ time** and **space** (vertical and horizontal) variability



to ultimately define:

- vertical and horizontal resolution of measurements,
- measurement precision (RMS) and accuracy (bias)

but the **positive gradients** persist for a **greater portion of the year** and the **annual-mean gradients** also show **higher atmospheric CO₂ concentrations near the surface than aloft** (Stephens et al., 2007),

, with the only exception of the **vertical region from the forest floor to the top of the canopy** where uncertainty is **7 % (or 25 ppm)** and **135 % (or 470 ppm)**, respectively

The **three fundamental vibrations** of CO₂ are $v_1=1337\text{ cm}^{-1}$, $v_2=667\text{ cm}^{-1}$, $v_3=2349\text{ cm}^{-1}$.